

# Spatial and seasonal heterogeneity of sea ice microbial communities in the first-year ice of Terre Adélie area (Antarctica)

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**ABSTRACT:** Spatial and temporal changes in sea ice microbial communities were investigated at 4 stations located along a south-north transect on the land-fast ice of Dumont d'Urville station area (Adélie Land, 66° 40' S, 141° 01' E) during the ice coverage period (April to December). A seasonal pattern was observed in microalgae, bacteria and protozoan abundance distribution. A maximum chlorophyll *a* concentration occurred during fall ice formation in the surface layer with the highest values at the near-shore station (100 to 215 µg l<sup>-1</sup>). A second maximum was observed before ice breaking in the bottom ice (50 to 90 µg l<sup>-1</sup>). Microalgal communities were dominated by diatoms (>86% of the total cells), mainly represented by *Fragilariopsis*, *Nitzschia*, *Navicula* and *Pseudonitzschia* species. *Fragilariopsis curta* was the dominant species during the first bloom whereas *Fragilariopsis cylindrus*, *Nitzschia longissima* and *Tropidoneis* sp. were the main contributors during the second bloom at the bottom ice core. Maximum protozoan abundance was recorded during the fall bloom in the surface layer with dominance of ciliates, which contributed more than 75% of total cell numbers. During this period, the maximum ciliate abundance was associated with the maximum bacteria and diatom numbers and microalgal biomass. The dramatic decrease of the ice algal biomass from south to north paralleled that of the underlying water phytoplankton available for new ice incorporation. The spatial algal biomass decrease could explain the parallel decrease in the abundance of bacteria and heterotrophic ciliates through trophic interactions.

**KEY WORDS:** Antarctica · Land-fast ice · Microbial communities · Chlorophyll *a* · Nutrients

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## INTRODUCTION

Sea ice is recognized to be a dynamic and complex habitat for microbial communities that show high spatial and temporal variabilities. In the Southern Ocean, the circumpolar extent of sea ice varies from 5 to 20 million km<sup>2</sup> seasonally (Maykut 1985, Legendre et al. 1992). Antarctic sea ice represents one of the largest and most dynamic ecosystems of the world ocean. During initial ice growth, microbial communities are incorporated into several distinct microhabitats of sea ice (Horner et al. 1988). These so-called sympagic organisms consist of both autotrophic and heterotrophic assemblages. They are characterized by seasonal and spatial variations with regard to their composition, dis-

tribution and abundance (Horner et al. 1992, Ackley & Sullivan 1994, Archer et al. 1996). The spatial and temporal variations of sea ice communities suggest an active food web within the ice, channeling algal production to bacteria and protozoa (Stoecker et al. 1998, 2000, Delille et al. 2002). Sea ice algae are estimated to contribute ca. 24% of the total biogenic carbon produced in the ice-covered Southern Ocean (Legendre et al. 1992). The highest biomass of algae occurs near the bottom of fast ice and in platelet ice (Ackley et al. 1979). In fast ice, which occupies 1 to 5% of the total ice cover area around Antarctica, the algal standing crop is 3 orders of magnitude higher than in pack ice (Ackley & Sullivan 1994). Diatoms are generally the dominant group, with the highest microalgal biomass

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(Watanabe et al. 1990, Garrison & Buck 1991, Archer et al. 1996, McMinn 1996, Moro et al. 2000, Riaux-Gobin et al. 2000). However, phytoflagellates are often dominant during spring in the upper sea ice (Stoecker et al. 1997, 1998). Sea ice contains a wide variety of bacterial assemblages. Most bacteria isolated from sea ice have been found to be pigmented and highly cold-adapted, with both free-living and epiphytic bacteria present (Grossi et al. 1984). Most taxa isolated from sea ice belong to the  $\gamma$ -proteobacteria and the *Cytophaga-Flavobacterium-Bacteroides* division (Bowman et al. 1997, Gozink et al. 1998, Junge et al. 1998, Nichols et al. 1999, Reddy et al. 2002). 16S rDNA clone library analysis corroborated these culture data (Brown & Bowman 2001). Heterotrophic protozoa are also present in sea ice (Garrison & Buck 1991, Stoecker et al. 1993, Delille et al. 2002, Song & Wilbert 2002) and are mainly composed of ciliates, nanoflagellates and dinoflagellates (Archer et al. 1996, Garrison et al. 2005). They play a major role in the removal of algal and bacterial biomass in Antarctic waters (Becquevort et al. 2000, Caron et al. 2000, Vaqué et al. 2004). Despite the lack of data on the consumption rate of bacterial and algal production by heterotrophic protozoa, some indirect measurements or extrapolations suggest that protozoan grazing plays an important role in sea ice (Garrison & Buck 1991, Garrison et al. 1993, Archer et al. 1996). Few studies address the abundance and biomass of the 3 components of the microbial loop in the sea ice (Archer et al. 1996). Generally,

autotrophic biomass is dominant in sea ice. In the fast ice of Adélie Land it contributed on average 80.6% of the total biomass whereas bacterial and protozoan biomass accounted for only 16.4 and 3%, respectively (Delille et al. 2002).

Due to the difficulties of accessing the ice environment during ice formation and in the polar winter night, little is known about the spatial and seasonal distribution and composition patterns of the microbial assemblage in sea ice. Most of the previous studies on ice microbial communities distribution have been focused on short-term observations in limited near-shore regions (Watanabe et al. 1990, Archer et al. 1996). The purpose of the present work was to examine spatial and temporal heterogeneity of sea ice microbial communities inhabiting first-year ice at 4 stations located along a south-north transect in the Dumont d'Urville station area.

## MATERIALS AND METHODS

The study was conducted during the sea ice period from April to December 1998 in the Pointe Géologie Archipelago close to the French Antarctic station of Dumont d'Urville (Adélie Land, 66° 40' S, 141° 01' E). Sampling was carried out at 4 stations chosen for their accessibility throughout the ice season (Fig. 1). They were located along a south-north transect. Stn A corresponded to the zone previously investigated (Delille et al. 2002). It was located 500 m offshore in the middle of the channel between the main islands of the archipelago (20 m depth). The other intermediate stations (Stns B and C) were chosen in order to achieve a regular distribution along the study transect. Sampling was carried out from April to mid-December at Stn A (16 samples) and Stn B (13 samples), from the end of April to the beginning of December at Stn C (12 samples), and from June to mid-November at Stn D (7 samples).

Ice samples were collected in a homogeneous and solid layer of fast ice using 10 cm (internal diameter) ice-coring augers. Core-horizons were cut vertically into 20 cm segments with a sterile blade. To avoid contamination from the ice-auger, subsamples were taken from the centre of these segments, weighed, and stored in a cool sterile glass box prior to melting. Subsamples were gently melted in a known volume of sterile artificial seawater composed of 30 g l<sup>-1</sup> NaCl dissolved in filtered distilled water. This dilution technique reduces osmotic shock and cell loss. Following melting in the dark at 4°C for ca. 24 h, fixatives were added to specific subsamples. Bacterial samples for total bacterial counts were preserved with particle-free

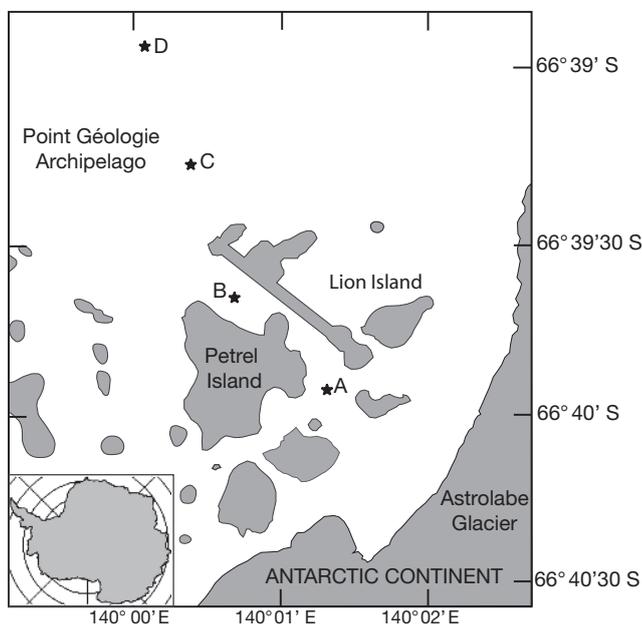


Fig. 1. Location of the 4 sampling stations (Stns A, B, C and D) at the Pointe Géologie Archipelago, Terre Adélie, Antarctica

formaldehyde (ca. 4 % final concentration). Protozoan samples were preserved with 1 % acid Lugol's iodine solution and algal samples with formalin (final concentration ca. 0.4 %). Samples were stored in the dark at room temperature until laboratory analysis.

Melt samples were filtered through Whatman GF/F glass fiber filter under low vacuum (<5 mm Hg). Nitrate, silicic acid and orthophosphate were determined using a standard automated method (Tréguer & Le Corre 1975). Due to the absence of salinity and temperature measurements, the nutrient concentrations were not normalized.

Chlorophyll *a* (chl *a*) was extracted with 90 % acetone and their concentrations measured using a Perkin Elmer MPF 66 spectrofluorometer (Neveux & Panouse 1987).

For taxonomic analysis and enumeration of algal communities, aliquots of 100 ml were collected once a month at Stn C from different parts of the ice cores. Algal cells were counted using an Olympus inverted microscope according to procedures described by Utermöhl (1958). Larger micro-sized (>20 µm) heterotrophic dinoflagellates (e.g. *Gyrodinium*, *Protoperdinium*) were identified on the basis of literature descriptions; however, only total dinoflagellate counts at Stn C are presented in this paper.

Protozoa were identified and enumerated with inverted microscopy (Utermöhl 1958). Subsamples (50 ml) were allowed to settle for 24 h and the whole cell counting was done with a Leitz Diavert microscope with ×25 and ×40 objectives and phase contrast illumination. This technique is suitable for the enumeration of ciliates, dinoflagellates and euglenophytes, but small (<10 µm) heterotrophic flagellates cannot be differentiated from autotrophs. All naked ciliates and total dinoflagellates were counted in size-classes with 10 µm intervals.

Total bacteria were enumerated by epifluorescence microscopy (Hobbie et al. 1977). Direct counts (AODC) were performed using an Olympus BHA microscope with acridine orange staining onto a 0.2 µm pore size black Nuclepore filter. A minimum of 500 fluorescing cells with a clear outline and definite cell shape were counted under oil immersion (×1000) in a minimum of 10 randomly chosen fields. The number of viable psychrotolerant aerobic heterotrophic microorganisms in each sea ice sample was estimated using the spread plate technique on Nutrient Agar 2216 (Oppenheimer & ZoBell 1952). Inoculated plates (6 replicates) were incubated for 20 d at 4°C. Heterotrophic counts (colony forming units, CFU) are only representative of culturable bacteria, however, they are an useful bacterial indicator corresponding to a small group of active bacteria that react immediately to changes in their nutrient supply (Delille & Bouvy 1989, Rheinheimer et al. 1989).

In addition, chl *a* and inorganic nutrients concentrations were measured on surface underlying sea water samples collected by opening a sterile glass bottle at each station.

## RESULTS

In Adélie Land, a land-fast ice cover develops each year from March-April to December. The ice thickness increases from a few centimeters to a maximum of ca. 2 m. In mid-January the ice cover begins to melt and breaks suddenly. Residual ice floes are then pushed away. Due to direct exposure to strong catabatic winds the snow cover is generally very thin (<2 cm). Thus the snow cover thickness could be considered as relatively negligible during the present study. An intense temporal variability in air temperature and solar irradiation occurs in Antarctic. During the study period, air temperature ranged from a maximum of -5°C during summer to a minimum of -30°C during winter and solar irradiation varied from 0.01 kJ cm<sup>-2</sup> in June to 2.8 kJ cm<sup>-2</sup> in November-December.

Concentrations of inorganic nutrients in the fast ice changed dramatically during the study period. Mean nitrate concentrations in the entire ice column varied from 4.0 at Stn A to 5.4 at Stn D. The highest mean values were observed from April to July at Stns A, B and C and between July and August at Stn D (Table 1). Maximum nitrate concentrations (10 to 12 µmol l<sup>-1</sup>) were observed in the bottom ice in May-June at Stns A, B and C and in July-August at Stn D. Phosphate mean concentrations ranged from <1 to 3.6 µmol l<sup>-1</sup> and their seasonal distribution did not show noticeable variability at the 4 stations (Table 1). At Stn A, the mean concentrations were high (1.3 to 3.6 µmol l<sup>-1</sup>) throughout the study period. At the 3 other stations, the mean values were <1.5 µmol l<sup>-1</sup>. Like phosphate distribution, silicic acid did not show a well defined seasonal pattern at the 4 stations. Mean concentrations reached very high levels over the study area (22.7 to 121.5 µmol l<sup>-1</sup>; Table 1). Maximum values (>100 µmol l<sup>-1</sup>) were generally observed in the ice bottom layer.

In the surface underlying water, mean nutrient concentrations were higher than in the sea ice (data not shown). Orthophosphate and nitrate values were 2 and 6-fold higher in underlying seawater than in sea ice cores, respectively, while silicic acid concentrations were equivalent.

In sea ice, chl *a* concentrations showed spatial and temporal distribution. The pigment biomass reached highest values at the near-shore Stn A and the lowest values at the offshore Stn D (Fig. 2). At Stn A, highest chl *a* concentrations (100 to 215 µg l<sup>-1</sup>) were encountered at the beginning of ice formation in the surface

Table 1. Monthly mean concentrations of nitrate, phosphate and silicate in the entire sea ice column at the 4 study stations. nd = no data

Month	Nitrate ( $\mu\text{M}$ )		Phosphate ( $\mu\text{M}$ )		Silicate ( $\mu\text{M}$ )	
	Mean	Range	Mean	Range	Mean	Range
<b>Stn A</b>						
Apr	4.71	1.9–7.5	3.62	1.0–6.3	52.49	21.5–102.2
May	5.21	2.0–8.0	2.68	0.9–5.5	73.08	49.2–105.1
Jun	4.78	2.0–10.2	3.32	1.3–7.3	54.89	21.9–150.9
Jul	4.07	1.9–7.1	2.73	0.7–6.4	47.15	33.5–64.1
Aug	nd		nd		nd	
Sep	3.94	1.8–7.0	3.22	0.9–6.5	58.58	25.5–103.1
Oct	3.97	1.7–8.2	3.52	0.9–10.2	43.03	29.9–93.1
Nov	2.99	2.6–4.1	2.33	0.7–5.4	48.38	12.4–87.1
Dec	2.20	1.4–3.5	1.28	0.5–2.9	57.18	43.8–73.9
<b>Stn B</b>						
Apr	7.11	5.8–8.2	1.53	0.9–2.3	22.71	16.4–29.4
May	5.32	2.0–11.1	0.96	0.4–1.7	30.14	13.4–80.4
Jun	5.18	2.2–10.2	0.91	0.4–1.8	62.39	27.2–156.0
Jul	4.66	2.2–6.9	0.74	0.5–1.0	92.24	52.1–130.1
Aug	4.26	2.1–7.8	1.23	0.6–2.4	59.64	40.2–92.8
Sep	3.85	2.2–5.6	0.67	0.5–0.9	32.26	21.9–38.4
Oct	3.24	1.7–6.8	0.53	0.3–1.1	53.18	38.4–112.4
Nov	3.94	2.3–9.9	1.15	0.5–2.2	51.22	20.6–155.2
Dec	2.81	1.9–3.7	0.87	0.4–2.9	68.5	57.9–79.9
<b>Stn C</b>						
Apr	4.99	3.9–7.0	1.37	1.2–1.5	24.60	20.1–27.2
May	5.65	1.9–10.7	0.89	0.6–1.4	43.67	29.4–68.1
Jun	6.07	3.0–10.8	0.91	0.4–2.2	82.04	36.2–107.7
Jul	4.08	1.7–5.7	0.78	0.6–0.9	41.70	32.3–59.6
Aug	3.89	1.9–5.1	2.82	0.8–3.2	59.24	39.0–85.4
Sep	4.41	2.2–6.5	0.90	0.7–1.2	74.72	32.4–134.7
Oct	3.33	1.9–6.5	0.91	0.7–1.3	99.00	88.8–114.2
Nov	3.26	2.9–4.2	0.97	0.5–1.3	46.04	39.7–54.7
Dec	2.44	1.9–3.7	1.31	0.6–3.9	60.67	50.0–70.4
<b>Stn D</b>						
Jun	7.24	6.5–8.5	1.09	1.0–1.2	35.29	23.1–47.4
Jul	7.16	3.4–12.4	1.48	1.0–2.3	80.80	26.9–139.6
Aug	5.68	4.2–10.7	0.91	0.7–1.2	55.24	28.2–94.4
Sep	4.42	2.9–7.3	0.86	0.6–1.2	33.95	27.3–39.3
Oct	3.96	2.3–6.5	0.57	0.4–0.9	121.51	86.0–261.9
Nov	3.86	2.1–7.5	0.66	0.4–1.2	58.40	40.6–81.8

layer. Thereafter a regular decrease occurred until ice breaking. A second chl *a* maximum (50 to 90  $\mu\text{g l}^{-1}$ ) was observed just before ice melting at the bottom ice. The same seasonal pattern was observed at Stns B and C with lower maximum chl *a* values of 50  $\mu\text{g l}^{-1}$  in April and 10 to 40  $\mu\text{g l}^{-1}$  in November. At Stn D, the concentrations were low ( $\leq 2 \mu\text{g l}^{-1}$ ) throughout the ice-covered period, except a maximum of 30  $\mu\text{g l}^{-1}$  at the bottom ice in November.

In the surface underlying seawater, chl *a* concentrations showed the same spatial and seasonal pattern as in ice at the 4 stations. During autumn ice formation, maximum chl *a* values were 8.8, 0.44, 0.49 and 0.03  $\mu\text{g l}^{-1}$ , at Stns A to D, respectively (Fig. 3). During winter, the values decreased and were  $< 0.2 \mu\text{g l}^{-1}$ . A second peak was observed in November. The highest chl *a*

concentration was recorded at Stn A (200.7  $\mu\text{g l}^{-1}$ ). The maximum values were lower at Stns B and C (156.3 and 166.5  $\mu\text{g l}^{-1}$ ), and Stn D (38.6  $\mu\text{g l}^{-1}$ ).

At Stn C, total cell concentrations ranged, throughout the entire period, from 1.1 to  $6.7 \times 10^6$  cells  $\text{l}^{-1}$  in the surface layer (0 to 20 cm) of the ice core to 1.2 to  $14.3 \times 10^6$  cells  $\text{l}^{-1}$  in the 60 to 140 cm stratum). The high cell densities ( $> 10^6$  cells  $\text{l}^{-1}$ ) in deeper ( $> 40$  cm) parts of the ice core were present throughout the study period, but a maximum ( $> 13 \times 10^6$  cells  $\text{l}^{-1}$ ) was found in May and June at 60 to 100 cm. In all months at different ice core levels, diatoms formed 86.4 to 99.1% of the total cells (Fig. 4a); however, there were a few exceptions at the surface in May, June, and October, when they contributed 29 to 64% of the total cells. Dinoflagellates, both heterotrophic and autotrophic, were either not found (as in May), or formed only 0.3 to 3.3% of the total cells (Fig. 4b). Exceptionally, between 100 and 140 cm in November they made up 7.8% of the cells and displayed a maximum of 5.7 to  $6.2 \times 10^4$  cells  $\text{l}^{-1}$ . Total nanoflagellates, which included Prymnesiophytes, Prasinophytes, and 4 to 10  $\mu\text{m}$  monads, usually formed  $< 5\%$  (Fig. 4c); only in October and November they contributed 28 to 50% between 20 and 60 cm. The highest numbers ( $5.6$  to  $6.6 \times 10^5$  cells  $\text{l}^{-1}$ ) were observed at 40 to 80 cm in November. Diatoms, the prevalent algae in the ice flora assemblage, were represented mainly by a few pennate species of the genera *Fragilariopsis*, *Nitzschia*, *Navicula* and *Pseudo-nitzschia* (Fig. 5). It is

interesting to note that in July and in October–November samples obtained especially in the lower ice core parts (60 to 100 cm), but also at the surface (July, November), contained about 60 to 98% of empty frustules of *Fragilariopsis curta* and *F. cylindrus*. On the other hand, in May, June and September, only full, young cells were found of chiefly *F. curta*. *F. cylindrus* and *F. curta* were the overall dominant diatoms along the ice core; the former species was more numerous in the 40 to 60 cm layer, while the latter at the bottom of the ice, mainly in May and June (Fig. 5). In the upper parts of the ice column where the lowest cell densities ( $< 0.4 \times 10^4$  cells  $\text{l}^{-1}$ ) were observed, 2 other species, *Nitzschia longissima* and *N. closterium*, were relatively abundant in most months (Fig. 5). In July, when none, or very few, live cells of *F. curta* were found in the sur-

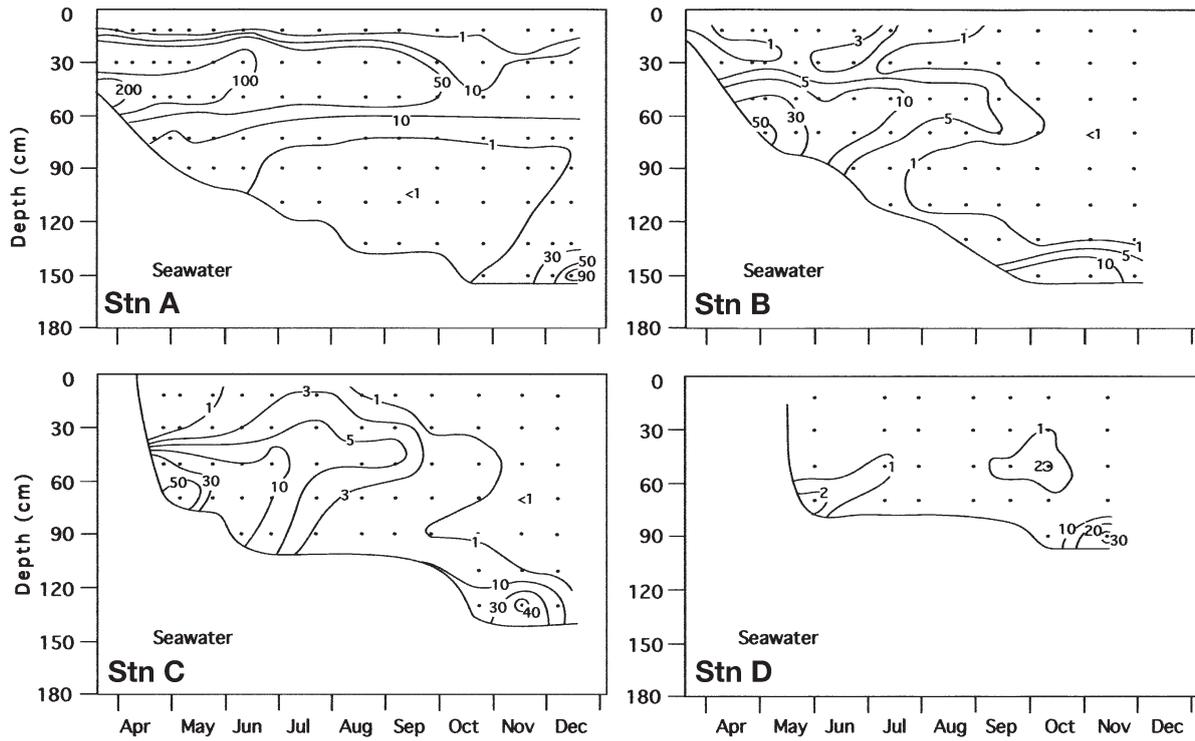


Fig. 2. Vertical distributions of chlorophyll a concentration ( $\mu\text{g l}^{-1}$ ) in sea ice of Stns A to D

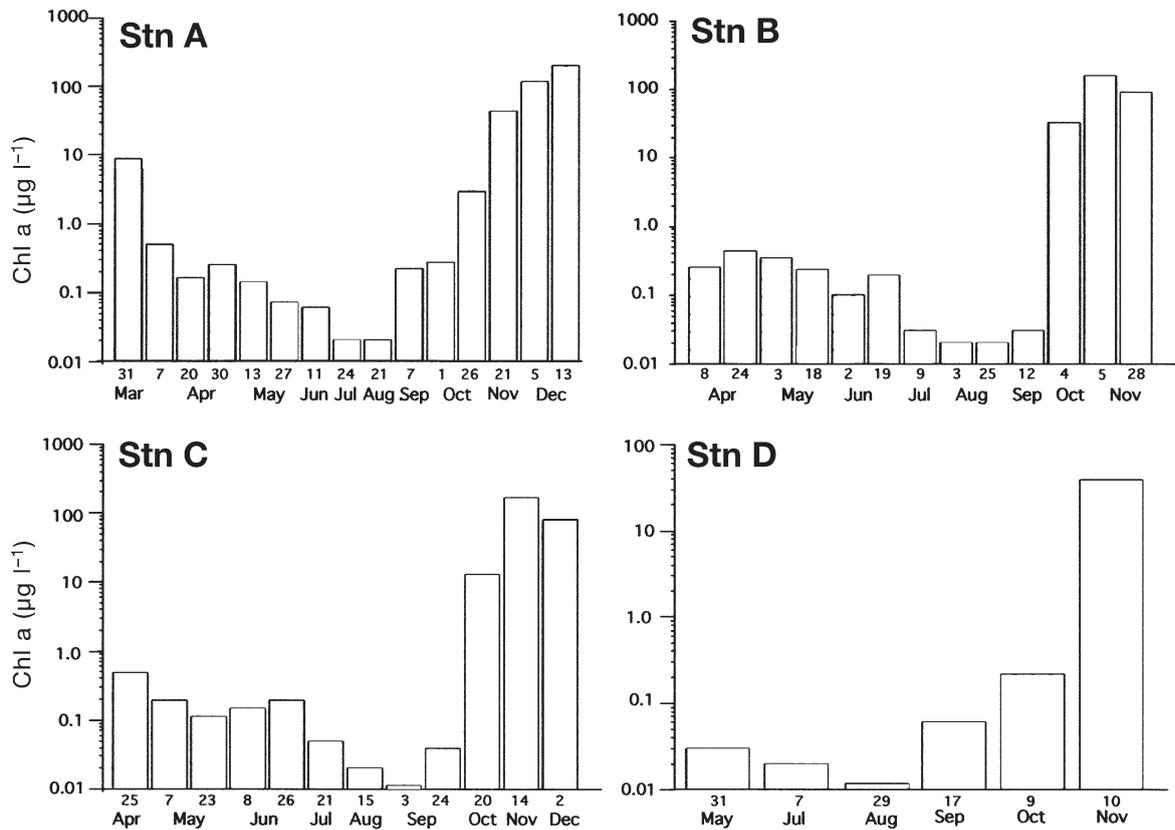


Fig. 3. Chlorophyll a concentrations ( $\mu\text{g l}^{-1}$ ) in the surface underlying water at Stns A to D

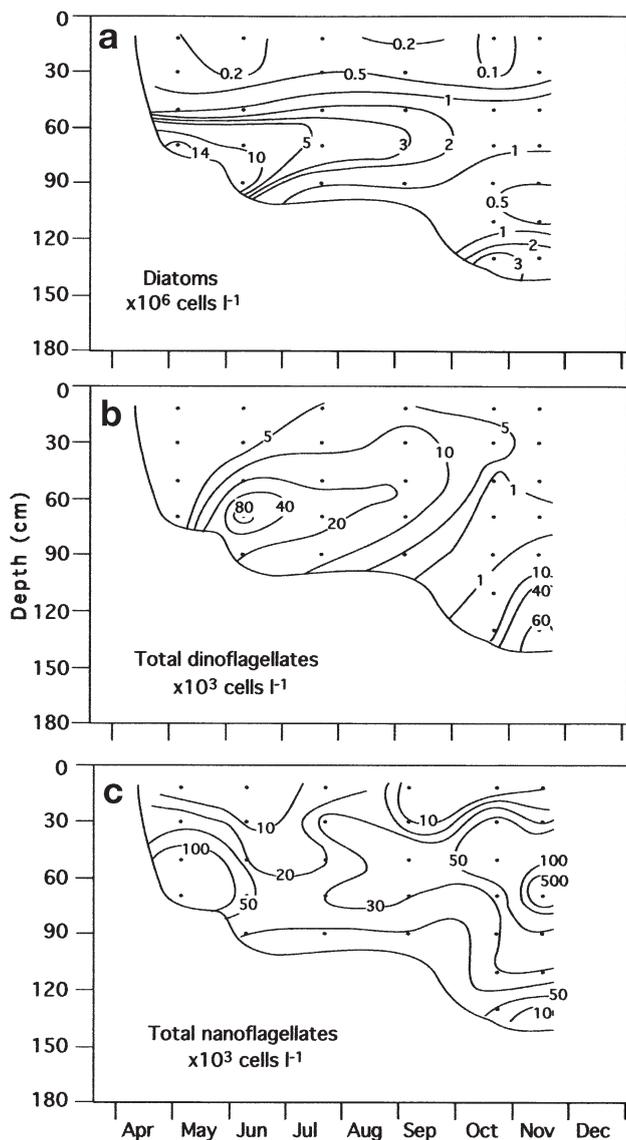


Fig. 4. Vertical distributions of (a) diatoms ( $\times 10^6$  cells  $l^{-1}$ ), (b) total dinoflagellates ( $\times 10^3$  cells  $l^{-1}$ ), (c) total nanoflagellates ( $\times 10^3$  cells  $l^{-1}$ ) in sea ice of Stn C from May to late November

face stratum, these 2 species formed together 70 to 80% of diatoms. They were absent from lower ice levels, or contributed less than 1%. *Navicula glaciei* formed 7 to 33% of the cells in the 0 to 40 cm stratum. Small young cells and spores of *Corethron* made up about 12% of diatoms in October at 20 to 80 cm; *Nitzschia subcurvata* was present from July to November in the 20 to 140 cm ice core stratum, forming 1 to 8.6% of the cells. *Chaetoceros dichchaeta* and *Pseudonitzschia prolongatoides* were encountered (4 to 8 and 12%, respectively) between 40 and 80 cm. Some species, e.g. *Nitzschia stellata*, appeared in abundance in June (60 to 80 cm:  $2.0 \times 10^5$  cells  $l^{-1}$ ) and in July (80 to

100 cm:  $7.4 \times 10^4$  cells  $l^{-1}$ ). *Synedropsis* sp. and *Tropidoneis* sp. contributed 9 to 23% of the cells at the bottom of the ice in October and November ( $3.7$  to  $5.5 \times 10^5$  cells  $l^{-1}$ ). Other species found sporadically included *Thalassiosira* spp., *Fragilariopsis kerguelensis*, *Pseudonitzschia turgidula*, *P. lineola/barkleyi*, *P. heimii*, *Planktoniella* sp., *Nitzschia taeniiformis*, *Eucampia* sp., *Dactyliosolen antarcticum*, *Cocconeis* sp. Spores of *Nitzschia* and *Corethron* spp. (6 to 9% of diatoms) were found in October, and of the latter species in July in the surface layer. Dinoflagellates observed in June, July, September and October, included the cells and spores of the heterotrophic *Gyrodinium* spp., *Gymnodinium* spp., and autotrophic *Prorocentrum* spp.; in November there were the cells and spores of *Prorocentrum* spp. and heterotrophic *Protoperidinium* spp..

Throughout the ice period, total protozoan abundance, including total ciliates, dinoflagellates, euglenoids and heliozoa, ranged from  $10^3$  to  $80 \times 10^3$  cells  $l^{-1}$  (Fig. 6). Two maxima were observed, one at Stns A and C in April-May and at Stn D in July and, the other at Stns A and C in August-September. Maximum cell numbers decreased from  $80 \times 10^3$  cells  $l^{-1}$  at the coastal Stn A to  $15$  to  $20 \times 10^3$  cells  $l^{-1}$  at the offshore Stns C and D. At Stn D, except a maximum of  $15 \times 10^3$  cells  $l^{-1}$  in the ice bottom in July, total protozoan cell numbers were  $< 2 \times 10^3$  cells  $l^{-1}$  throughout the ice period. The protozoan assemblage was mainly dominated by ciliates during the first maximum (Fig. 6). They contributed more than 75% of the total protozoan numbers during April-May at Stns A and C. At the 3 stations, ciliate assemblages were mainly dominated by 10 to 90  $\mu$ m sized unidentified species. Small quantities of Tintinnidae (*Codonellopsis* spp., *Cymatocylix* spp., *Laackmanniella* spp.), Strombididae (*Strombidium* spp.) and Strobiliidiidae (*Strobilidium* spp.) were also encountered. Heterotrophic euglenoids were numerically the second component of the protozoan communities in April-May at Stns A and C (data not shown). On the other hand, they were the major component of the cell maxima observed in July at Stn D between 60 and 100 cm and in September at 20 to 40 mm at Stns A and C. Heterotrophic dinoflagellates were either not found, or present in only small numbers at the 3 stations. Heliozoa were absent, except at Stn A in December at 20 to 40 cm ( $78 \times 10^3$  cells  $l^{-1}$ ). Adult copepods were absent and copepod nauplii were scarce. Large quantities of unidentified cysts were present in the ice cores throughout the study period.

There were no significant differences between the abundance of total culturable bacteria and mean cell volumes distribution observed at the 4 stations. The general seasonal patterns found for these 2 parameters were similar to those previously reported (Delille et al.

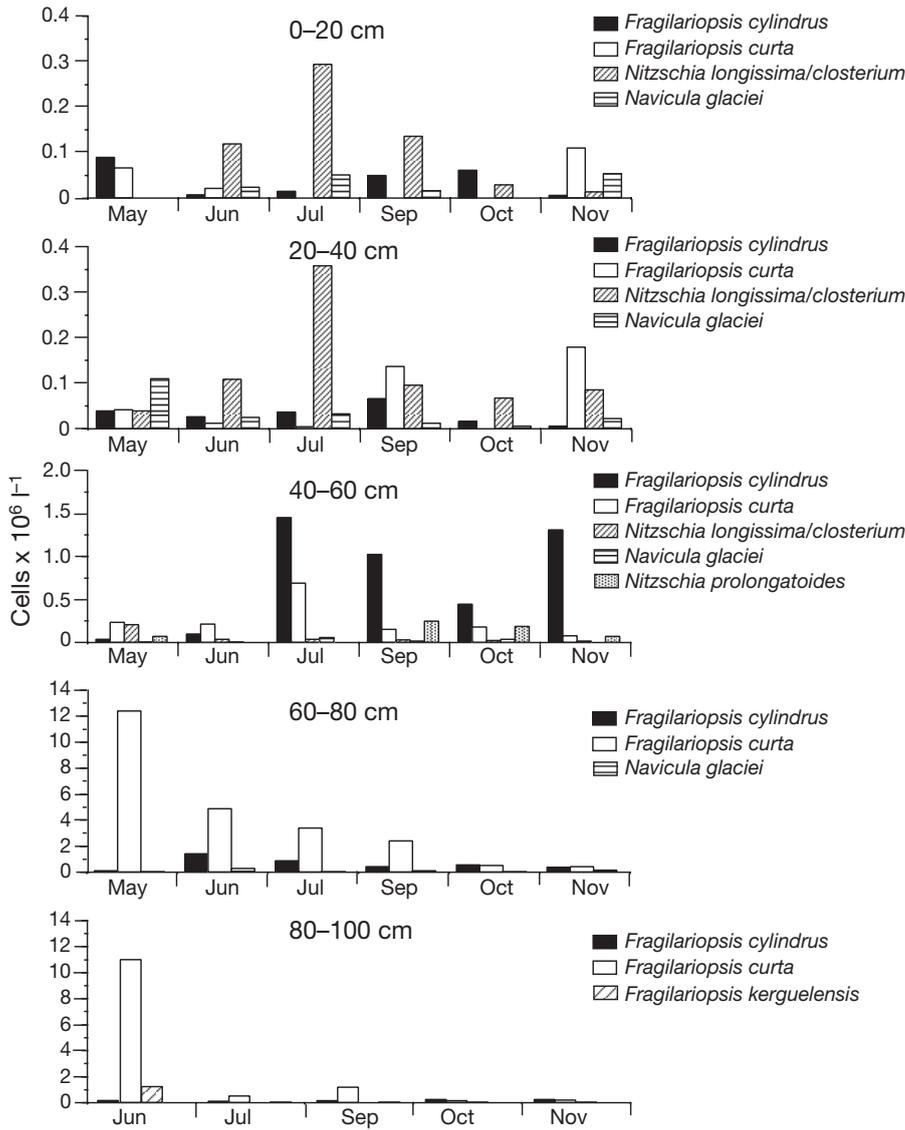


Fig. 5. Cell distribution of dominant diatoms ( $\times 10^6$  cells  $l^{-1}$ ) in the different layers between surface and bottom sea ice of Stn C

2002, data not shown). In contrast, there was a clear decreasing gradient of culturable bacteria from the near-shore Stn A to the offshore Stn D (Fig. 7). Maximal levels of bacteria abundance were detected during ice formation and just before the summer thaw. However, both autumn and spring maximal values observed at Stn A were severely reduced at the offshore stations. Autumn values higher than  $10^7$  CFU  $l^{-1}$ , which were recorded until July at Stn A, fade in May at Stn B and were never recorded at Stn D. In the same way, spring abundance of bacteria, which could be higher than  $2 \times 10^7$  CFU  $l^{-1}$  in surface layer of Stn A in November/December, was never higher than  $2 \times 10^6$  CFU  $l^{-1}$  in similar layers of Stns C and D.

## DISCUSSION

We observed seasonal and spatial changes in the autotrophic biomass distribution. At the 4 stations, 2 peaks of chl *a* concentrations were found during the study period. The higher one occurred in the surface layer during the ice formation in April–May. The second one was observed in the bottom layer, just before ice break-up. Maximum chl *a* values ( $>200$  mg  $m^{-3}$ ) recorded at the coastal Stn A in April were similar to those observed in the first year fast ice in other Antarctic areas (Gunther & Dickman 1999), but they were less than those found by Palmisano & Sullivan (1983) and Watanabe et al. (1990). The seasonal distribution of autotrophic biomass was in accordance with that recorded previously in the same area. However, the maximum values of chl *a* recorded in this study at Stn A were about 4 to 5 times as high as those observed at the same station in 1993 and 1997 (Fiala & Delille 1999, Delille et al. 2002). A decrease in chl *a* concentration was observed from south to north. Mean chl *a* values decreased dramatically from the near-shore Stn A to the offshore Stn D. The highest chl *a* levels observed in surface layer during autumn could be attributed to mechanical inclusion and concentration of microalgae into the newly formed sea ice and also to *in situ* growth (Stewart & Fritsen 2004). During ice formation, a high level of biomass in the ice surface layer corresponds to a

high level of phytoplankton biomass in sea water (Watanabe et al. 1990). In our study, sea-ice algal biomass can be  $>20$  times higher than in the interface water. As observed at Stn A the highest chl *a* concentration ( $8.8$   $\mu g$   $l^{-1}$ ) in the surface underlying water corresponded to a maximum chl *a* concentration ( $200$   $\mu g$   $l^{-1}$ ) in sea ice. The second chl *a* biomass increase observed in the ice bottom layer in October–November just before ice-breakup was lower than that observed during the autumn. Such spring biomass increases have been previously reported (Palmisano & Sullivan 1983, Garrison et al. 1987, Watanabe et al. 1990, Syvertsen & Kristiansen 1993, Stoecker et al. 1998). A concomitant increase of chl *a* biomass in the

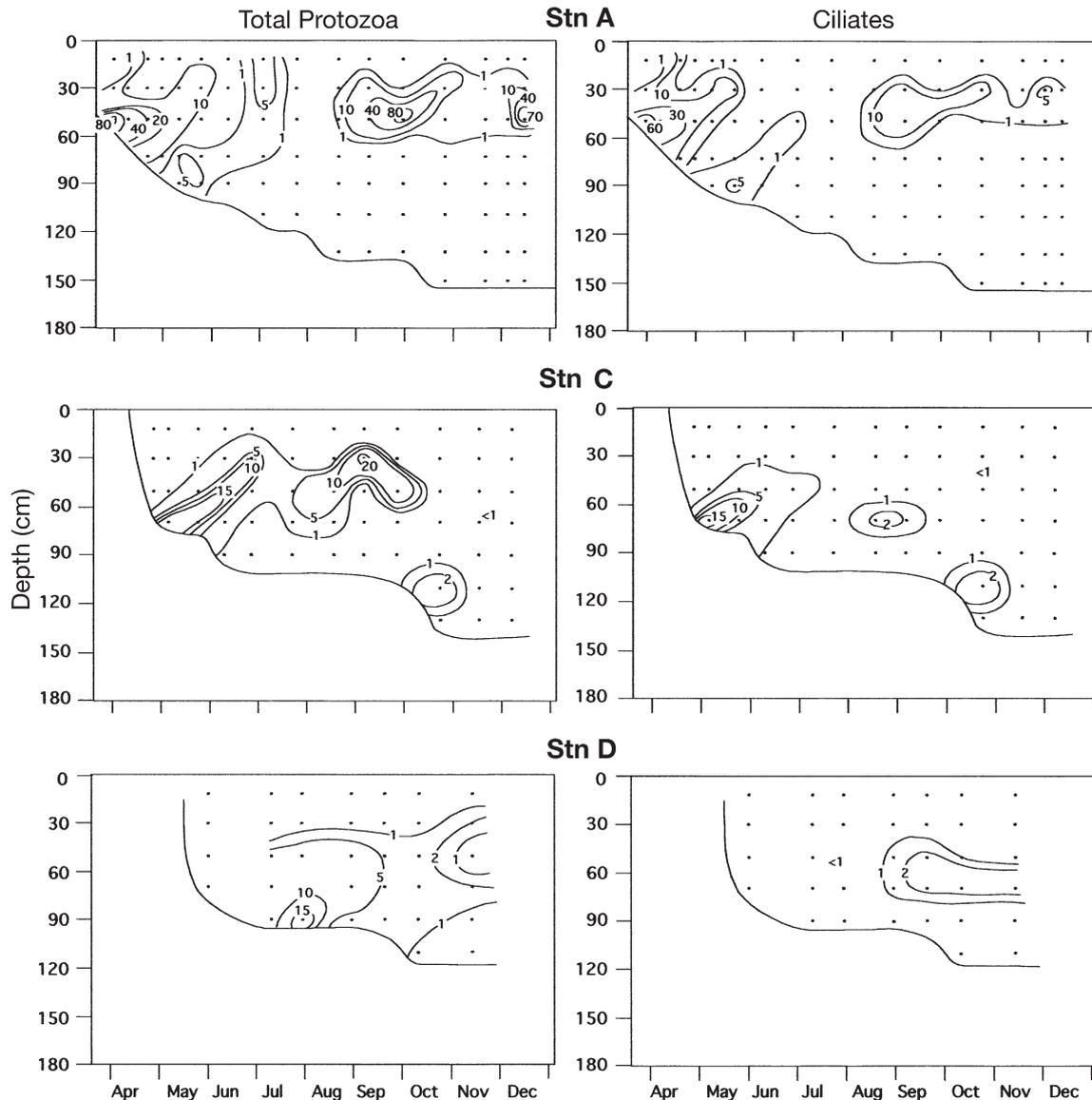


Fig. 6. Vertical distributions of total Protozoa ( $\times 10^3$  cells  $l^{-1}$ ) and heterotrophic ciliates ( $\times 10^3$  cells  $l^{-1}$ ) in sea ice of Stns A, C and D

sea ice and in the underlying surface water during this period was observed. Unfortunately, underlying water species were not determined and it is not possible to know if the 2 communities were similar or different. Generally, the communities in fast ice are quite different from those of the underlying water (Palmisano & Sullivan 1983, Riaux-Gobin et al. 2003), and would not provide an algal inoculation for a spring phytoplankton bloom (Palmisano & Sullivan 1983, McMinn 1996). The south-north positive gradient in nitrate and silicic acid concentration clearly reflected a link with ice algal biomass. During the autumn bloom silicic concentration was low in the ice surface layer ( $<25 \mu M$  at Stn A) suggesting its uptake by diatoms. At all 4 study stations, algal ice communities were dominated by diatoms.

Dominance of diatoms and especially of the pennates is a common feature which has been observed elsewhere in Antarctic coastal sea ice (Grossi & Sullivan 1985, Medlin & Hasle 1990, Watanabe et al. 1990, Archer et al. 1996, McMinn 1996). Dominant diatoms showed a vertical zonation along the ice column, a pattern observed also at other Antarctic localities, such as in the sea ice at Lutzholm Bay (Hoshiai 1977), or in bottom congelation ice community at McMurdo Sound (Grossi & Sullivan 1985). The presence of young cells of *Fragilariopsis curta* and *F. cylindrus* in May-June and also in September suggests 2 periods of these species reproduction. They were followed by an abundance of empty frustules in July and in October-November, providing further evidence for earlier

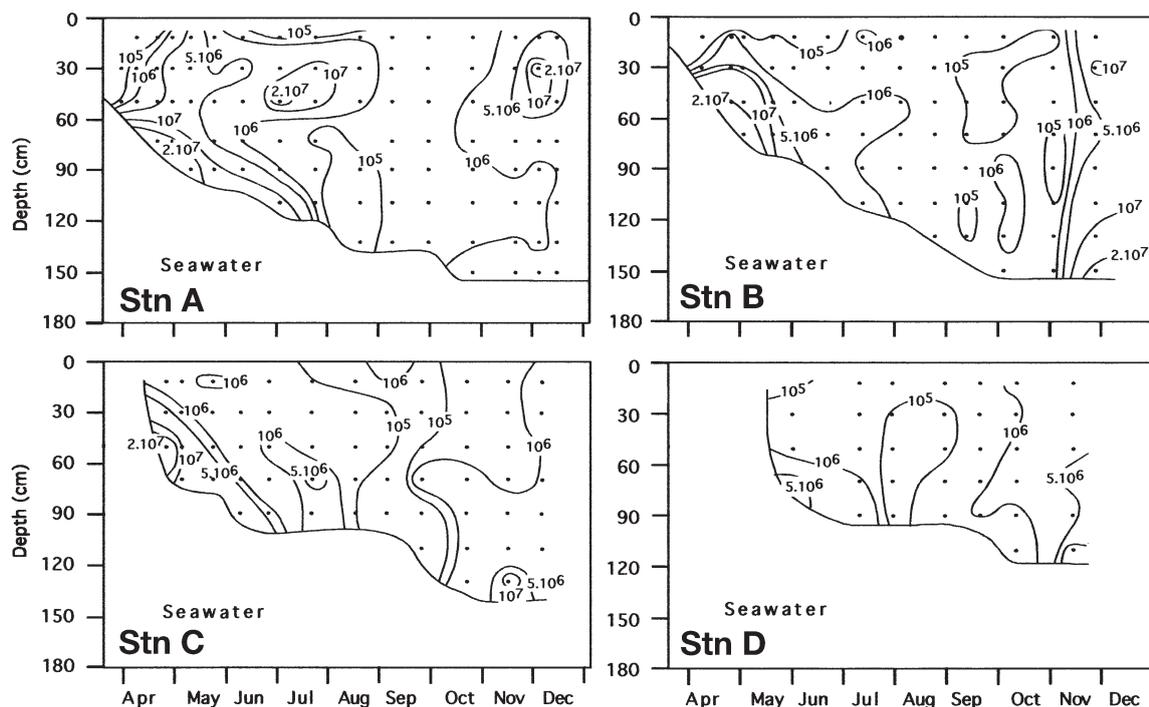


Fig. 7. Vertical distribution of culturable bacteria ( $\text{CFU l}^{-1}$ ) in sea ice of Stns A to D

reproduction and/or grazing by micro-heterotrophs. Seasonality of ice diatom distribution, both in terms of cell abundance and vertical occurrence, has also been reported from other Antarctic areas (Grossi et al. 1984). Apparently, algal ice assemblages may differ substantially depending on the type of ice (Scott et al. 1994), and the mode of ice formation (Clarke & Ackley 1984, Spindler et al. 1990). The dominant species, *F. curta* and *F. cylindrus*, found in the present study in the fast ice were also reported prevalent in both, fast ice and pack ice around Prydz Bay (Scott et al. 1994). These are planktonic species found in highest abundance in offshore and coastal Antarctic waters (Hasle 1969, Kopczyńska et al. 1986). Their highest numbers in the bottom part of the ice column in May, June and July suggest their entrapment from a bloom in water beneath the ice, although they apparently multiply successfully within the ice. Another species, *Nitzschia stellata*, very common in the bottom ice in July was first described from Adélie Land area (Manguin 1957), and its occurrence there was confirmed by Medlin & Hasle (1990). Vertical distribution of these species, which attain the highest concentrations at the bottom of the ice column, seem to be mainly controlled by the availability of nutrients in the underlying water (Grossi & Sullivan 1985). In contrast, *Nitzschia closterium*/*N. longissima* and *Navicula glaciei*, which dominate the less dense upper ice (10 to 20 cm) diatom assemblage, probably require higher light conditions than

other species for growth in the ice, as was suggested for the latter species by Whitaker (1977).

During the study period, distribution of culturable bacteria paralleled chl *a* distribution with maxima during autumnal sea ice formation and just before spring ice break up ( $100$  to  $200 \mu\text{g chl } a \text{ l}^{-1}$  and  $>2 \times 10^4 \text{ CFU}$ ). A clear decreasing gradient from the near-shore Stn A to the offshore Stn D was also observed. The good correlation between microalgal and bacterial abundance and biomass has been previously observed during spring and autumn in different Antarctic regions (Grossi et al. 1984, Lochte et al. 1997, Delille et al. 2002, Stewart & Fritsen 2004). However, the strength of the coupling between microalgae and bacteria varies seasonally. Heterotrophic bacteria are reliant on microalgae for their energy supply through dissolved organic matter that is provided from algae via processes including degradation of cells (Riemann & Sondergaard 1986), extracellular release (Suttle et al. 1991) and production of exopolymeric substances (Meiners et al. 2004). It has been suggested that, during the spring and autumn algal blooms, high levels of dissolved organic matter are available for bacterial consumption (Grossi et al. 1984, Fritsen & Sullivan 1999, Stewart & Fritsen 2004). In contrast, during winter, algal biomass is low and consequently, the low dissolved organic matter available limits the bacterial growth. During winter, the very low temperature of sea ice could also exert a strong control on bacterial

growth by the way of a decrease of substrate assimilation (Nedwell 1999, Stewart & Fritsen 2004). It has been suggested that during ice formation a differential trapping of bacteria and microalgae from sea water may also influence the bacteria-algae association in the pack ice (Stewart & Fritsen 2004).

Large amounts of algal and bacterial biomass observed during the blooms provide an important food source for other microbial consumers, predominantly protozoans. Among heterotrophic protozoan communities, nanoflagellates and dinoflagellates can be the major contributors to the biomass in fast ice (Archer et al. 1996). In this study, heterotrophic dinoflagellates accounted for less than 25% of the total protozoan numbers, and protozoan communities were usually dominated by ciliates. Although few studies address the abundance, biomass and distribution of Antarctic ciliates, their contribution in the sea ice biota has been recognized (Gradinger et al. 1999, Scott et al. 2001). Their concentrations are generally higher than in the underlying waters and may reach more than 70% of the total protozoan biomass (Burkill et al. 1995). In the present study, the ciliate numbers were lower and the autotrophic biomass was higher than that recorded in the previous study in the same area (Delille et al. 2002). Although it is difficult to relate abundance estimates to grazing impact, low ciliate numbers coupled with high autotroph biomass during this study are probably indicative of a lower grazing pressure than in the previous study. The highest ciliate numbers associated with maximum abundance of bacteria and microalgae observed during the fall ice formation suggests a mechanical trapping and concentration of these 3 components into the newly formed ice. Heterotrophic euglenoids were dominant in the ice surface layer of Stns A and B in September and in ice bottom layer of Stn D in July. Although they are not generally regarded as a major component of the heterotrophic biomass (Garrison & Buck 1989), they could contribute more than 50% of the heterotrophic biomass at certain times in the bottom and interior of the coastal sea ice (Archer et al. 1996). In ice samples low abundance of metazoa, dominated by copepods nauplii, suggests that their role in grazing effect was negligible.

Beside differences in the vertical distribution occurring over a very short distance (<1 m), a clear spatial heterogeneity from inshore to offshore in the different sympagic biota distribution was observed. A dramatic decrease occurred from south to north in the algal biomass. A similar gradient in chl *a* concentrations was reported in the same area by Riaux-Gobin et al. (2000). During ice formation, planktonic algae are physically concentrated by frazil ice, which affects results in chl *a* values in the newly formed ice. Evidence of algal biomass enrichment is provided in ice cores of Stns A, B

and C, during autumnal ice formation. Among the different factors controlling algal concentration in new ice, the date of ice formation is one of the most important (Fritsen & Sullivan 1999, Stewart & Fritsen 2004). In our study, the date of ice formation varied from the beginning of April at the coastal station to May at the offshore station. This could explain the south-north decrease of the underlying water phytoplankton biomass available for incorporation in the newly formed ice (from 8.8  $\mu\text{g chl } a \text{ l}^{-1}$  at Stn A, to 0.03  $\mu\text{g l}^{-1}$  at Stn D). This spatial decrease of the ice-incorporated algal biomass may affect the bacteria development by way of a decrease of the algal-derived inorganic substrate.

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