

Revised 21st century temperature projections

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ABSTRACT: Temperature projections for the 21st century made in the Third Assessment Report (TAR) of the United Nations Intergovernmental Panel on Climate Change (IPCC) indicate a rise of 1.4 to 5.8°C for 1990–2100. However, several independent lines of evidence suggest that the projections at the upper end of this range are not well supported. Since the publication of the TAR, several findings have appeared in the scientific literature that challenge many of the assumptions that generated the TAR temperature range. Incorporating new findings on the radiative forcing of black carbon (BC) aerosols, the magnitude of the climate sensitivity, and the strength of the climate/carbon cycle feedbacks into a simple upwelling diffusion/energy balance model similar to the one that was used in the TAR, we find that the range of projected warming for the 1990–2100 period is reduced to 1.1–2.8°C. When we adjust the TAR emissions scenarios to include an atmospheric CO₂ pathway that is based upon observed CO₂ increases during the past 25 yr, we find a warming range of 1.5–2.6°C prior to the adjustments for the new findings. Factoring in these findings along with the adjusted CO₂ pathway reduces the range to 1.0–1.6°C. And thirdly, a simple empirical adjustment to the average of a large family of models, based upon observed changes in temperature, yields a warming range of 1.3–3.0°C, with a central value of 1.9°C. The constancy of these somewhat independent results encourages us to conclude that 21st century warming will be modest and near the low end of the IPCC TAR projections.

KEY WORDS: Temperature projections · Climate change · Global warming · Climate models · Impact assessment

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1. INTRODUCTION

Much has been made of 21st century temperature projections from the Third Assessment Report (TAR) of the United Nations Intergovernmental Panel on Climate Change (IPCC) that indicate a rise of 1.4 to 5.8°C during the 1990–2100 period. The projections are based upon a number of 'storylines' describing what the IPCC believes are self-consistent social development possibilities for the next 100 yr. These are factored into what are known as 'SRES' scenarios, as they were published in an IPCC document known as the 'Special Report on Emissions' (IPCC 2000). These 'storylines' range from high population growth/high

emission assumptions to low population growth with substantial substitution of low greenhouse gas technologies such as wind, solar and nuclear.

This report re-examines the IPCC TAR range of 21st century temperature projections using 3 methods. The first adjusts the sensitivity of the thermal response to changes in atmospheric CO₂ based upon recent findings which imply that this sensitivity is less than previously assumed, and the second adjusts the SRES-projected changes in CO₂ using the actual trends established in the last quarter-century. The third adjusts the average slope of a large collection of General Circulation Models (GCMs) for observed increases in temperature during the recent warming of the past 25 yr.

2. ADJUSTMENT IN RADIATIVE FORCING, THERMAL SENSITIVITY, AND CLIMATE/CARBON CYCLE FEEDBACK

We used the 'Model for the Assessment of Greenhouse-gas Induced Climate Change' (MAGICC), a simple upwelling diffusion/energy balance (UD/EB) model (Wigley & Raper 1987, Raper et al. 1996), as the tool to investigate the effects of altering the climate parameters and emissions scenarios on average global temperatures for the next 100 yr. MAGICC takes as input trace gas emission scenarios (such as the ones from the TAR SRES) and outputs global average temperature for user-defined time steps. The use of a simple UD/EB model to emulate global-scale temperature results of fully integrated Atmosphere-Ocean General Circulation Models (A/OGCMs) has been a standard practice employed by the IPCC in order to simulate the effects of many different model parameter and emissions scenario combinations without the cost (both physical and temporal) of running full A/OGCMs.

The version of MAGICC that we used was extensively used by the IPCC in its Second Assessment Report (SAR) (IPCC 1996). We were not able to obtain the updated version employed by the IPCC in its TAR (IPCC 2001), despite repeated attempts to do so from its authors. The model used in the TAR produces higher temperatures for similar trace gas emissions than the SAR version because of the inclusion of a sea ice parameter, a shallower ocean mixed-layer, a climate/carbon cycle feedback, and several other minor modifications (Raper et al. 2001, IPCC 2001). Since the SAR version of MAGICC has only a limited set of user-configurable climate parameters, it is impossible to replicate the workings of the TAR version exactly; therefore, we had to work within the limitations of the software to produce output that best emulates the TAR results.

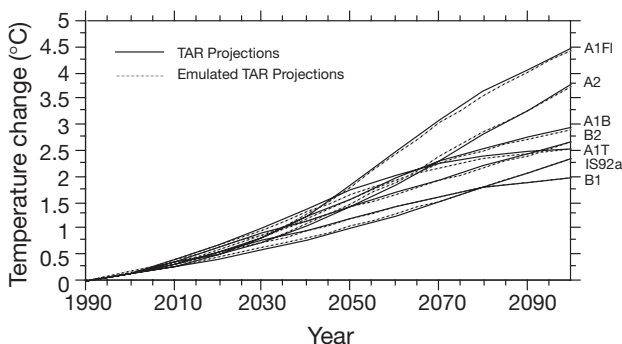


Fig. 1. Temperature projections for the 6 illustrative SRES scenarios and the IS92a scenario from the IPCC Third Assessment Report (TAR) (solid lines) and emulation of these scenarios using the SAR version of MAGICC (dotted lines)

The inclusion in TAR of a climate/carbon cycle feedback results in higher atmospheric CO₂ concentrations for a particular emissions scenario than when the climate feedback is not included. An important cause of the elevated CO₂ concentrations is enhanced soil respiration from elevated temperatures (IPCC 2001). To reproduce the enhanced CO₂ concentrations in the version of MAGICC we employed, we increased the CO₂ emissions at each time step (every 10 yr) within each of the 6 SRES illustrative scenarios (IPCC 2001) until we were able to closely match the CO₂ concentration curves given by Wigley (2000), which are the concentration profiles used in the TAR. However, this change alone was not enough to replicate the higher temperatures resulting from the MAGICC model differences. Therefore, we altered the climate sensitivity so that for a given input emission scenario (adjusted to account for the enhanced CO₂ concentrations produced by the CO₂ /climate feedback), the output temperature change for the year 2100 best matched the TAR values. In every case, we had to increase the climate sensitivity. Fig. 1 shows the comparison of the average temperature projections for 6 SRES illustrative scenarios as well as the IS92a scenario (the 'central' emissions scenario from SAR) as depicted in the TAR with the projected temperatures calculated by the SAR version of MAGICC with adjusted climate sensitivities. The close correspondence (the temperature projections never differ by more than 0.1°C, and typically the differences are less than 0.05°C) indicates that our adjustment procedures allow us to very closely emulate the TAR results. Although we cannot be sure that this close emulation will continue when model parameters are subsequently altered, our experience with the model indicates that it is well behaved under a wide range of parameter settings and the alterations that we make fall within the range of values that MAGICC has been developed to incorporate.

The range of warming estimated in the TAR from the full set of 35 SRES scenarios and the MAGICC emulation of 7 GCMs is 1.4 to 5.8°C by the year 2100. The range of the average projected warming from the 7 GCMs using the 6 illustrative SRES scenarios is 2.0 to 4.5°C (Fig. 1). To assess how these projections may change based upon alternative determinations of climate sensitivity and future emission scenarios, we adjusted the inputs and parameters of the MAGICC model to account for the following: (1) the finding of Jacobson (2001) that the direct warming effect of black carbon (BC) aerosols (soot) is on the order of 0.55 W m⁻², a value greater than that used in the TAR projections; (2) the reduction of the estimate of climate sensitivity to doubled CO₂ as calculated by Lindzen et al. (2001); (3) adjustment of the positive feedback

between global temperatures and emissions of CO₂ based upon the findings of Luo et al. (2001).

By focusing on only these 3 aspects of climate change, we are not implying that other climate and climate model parameters are better constrained. In fact, many are not. For instance, the level of uncertainty of the effects of other forcing agents, such as sulfate aerosols, far exceeds that of BC, at least within the framework of current determinations of these uncertainty levels (IPCC 2001). Our work is not intended as an exhaustive sensitivity analysis of the uncertainties of climate and climate modeling, but instead we examine the effect of refinements to the understanding of several key parameters—refinements that have occurred since the release of the TAR, and that have significant impacts on the TAR findings.

2.1. Adjustment to the radiative forcing due to the effects of black carbon aerosol

Jacobson (2001) determined that the global mean direct radiative forcing from BC aerosols was greater than any previous estimates, including the mean value used in the TAR models (0.55 vs 0.40 W m⁻²). There are 2 primary anthropogenic sources of BC aerosols of roughly equivalent importance: incomplete combustion of fossil fuels and biomass burning. While BC emissions pathways are not explicitly defined in the SRES scenarios, the assumption used in TAR MAGICC is that the fossil fuels contribution to the total BC emissions is scaled linearly with sulfur dioxide emissions while the biomass burning component is scaled with gross deforestation (IPCC 2001). For the case of fossil fuel burning, the direct effect of BC emissions can be combined with the direct effect of the emissions of organic carbon (OC) and sulfate aerosols to create an 'effective' direct sulfate forcing—a user-configurable value in MAGICC. The 1990 reference value used in TAR for this sum is -0.3 W m⁻², (-0.4 from sulfate, -0.1 from OC, and +0.2 from BC). If we apply the TAR assumption that the forcing from BC is equally split between biomass and fossil fuel sources, then the TAR value of +0.2 W m⁻² should be increased to +0.275 in accordance with the results of Jacobson (2001), bringing the 1990 reference effective sulfate forcing value to -0.225 W m⁻². This adjustment was explicitly made in the version of MAGICC used here.

There are no user-adjustable parameters for biomass burning in the SAR version of MAGICC. In half of the 6 SRES illustrative scenarios, there is little change over time in the amount of global deforestation, while in the other half, total deforestation declines about 5 to 30 % by the year 2100. The TAR 1990 reference value for the combined direct effect of BC and OC aerosols as a

result of biomass burning is -0.2 W m⁻² (-0.4 from OC aerosols, +0.2 from BC aerosols). Using the results of Jacobson (2001) would increase the total forcing to -0.125 W m⁻² (-0.4 + 0.275). This number declines to -0.092 W m⁻² (a change of only 0.033) for a 30 % decline in deforestation. We will consider this change to be of negligible climate consequence and ignore the impact of Jacobson (2001) on this half of the BC aerosol equation.

2.2. Adjustment to the climate sensitivity resulting from the iris effect

Lindzen et al. (2001) demonstrated evidence for an 'adaptive infrared iris' that acts as a negative feedback to control global temperature fluctuations. Through examinations of cloud observations made by the Japanese Geostationary Meteorological Satellite-5 over parts of the western Pacific, Lindzen et al. (2001) determined that the relative proportion of cirrus-to-cumulus cloud coverage varied with the sea surface temperature (SST). Higher SSTs resulted in fewer cirrus clouds while lower SSTs were associated with a greater proportion of cirrus clouds. Since cirrus clouds have a net positive radiative forcing, this results in a negative feedback in that fewer cirrus clouds allow more infrared radiation to escape from the warmer surface, which subsequently cools, leading to an increase in the cirrus coverage and a closing of the atmospheric infrared window. Lindzen et al. (2001) find that this negative feedback, if applicable throughout the tropics, would have a magnitude that would be nearly equivalent, but of opposite sign, to the magnitude of the sum of all other positive feedbacks present in the current generation of climate models. Lindzen et al. (2001) go on to calculate that the iris effect would lower the climate sensitivity in current GCMs by about 60 %. Several subsequent studies (Fu et al. 2001, 2002, Harrison 2002, Hartmann & Michelsen 2002, Lin et al. 2002) have challenged the Lindzen et al. (2001) results; however, in careful comments to these challenges (Chou et al. 2002a,b,c, Chou & Lindzen 2002, Bell et al. 2002, Lindzen et al. 2002) the results of Lindzen et al. (2001) have been confirmed. In the series of responses to the critiques, it has been suggested that the effect of the iris on the global temperature sensitivity was initially overestimated by perhaps 15 to 20 %. Thus, in accordance with these new results, we will take the reduction of the values of temperature sensitivity used by current generation GCMs to be 50 %—a value about 17 % less than that originally proposed by Lindzen et al. (2001).

While this may seem as a rather large reduction of the current manifestations of climate sensitivity as gen-

erated by climate models, it arises from the possibility that current climate models do not handle tropical cloud and moisture processes adequately. Lindzen et al. (2001) found that in a small sampling of climate model simulations forced by observed SSTs the observed iris effect was not present. Additionally, results consistent with the adaptive infrared iris mechanism have been recently reported by Chen et al. (2002), Wielicki et al. (2002) and Wang et al. (2002) when analyzing trends in longwave and shortwave radiation in the tropics from data collected by the satellite-borne instrumentation of a series of experiments designed to study the earth's radiation balance including the second Stratospheric Aerosol and Gas Experiment (SAGE II), the Earth Radiation Budget Experiment (ERBE), and the Clouds and the Earth's Radiant Energy System (CERES). These researchers found positive trends in outgoing longwave radiation and negative trends in reflected shortwave radiation from the mid-1980s to the late 1990s. A redistribution of the tropical cloudiness and moisture, including decreasing high level clouds and humidity and increasing low level clouds, has been identified as contributing to the observed trends. Furthermore, Wielicki et al. (2002) found that a suite of current climate models were unable to reproduce the observed data, even when forced with the observed SSTs, indicating a potential problem with the cloud response to SST forcing in the models, as implied by Lindzen et al. (2001). Harvey (2000), using a simple energy balance climate model, showed that a downward vertical moisture redistribution in the tropics could lead to a decrease in the temperature sensitivity of about 15%—an effect not replicated by more complex climate models and one that is independent of actual cloud properties. Gaffen et al. (2000) report that temperature trends derived from data collected by microwave sounding units show that the temperatures in the lower troposphere in the tropics have declined since 1979 while temperature data collected from surface-based instruments show a warming trend during the same period. Independent data from radiosonde observations confirm these trend differences. These disparate temperature trends have led to an increase in the vertical lapse rate in the tropics, a behavior not captured by the 3 climate models that were examined (Gaffen et al. 2000). Hegerl & Wallace (2002) further confirm that the temperature trends from the lower troposphere are less positive than those from the surface since the late 1970s, and for the period 1964–1981 this trend differential was reversed. They were unable to replicate the observed behavior using the most recent version of the Max Planck Institute for Meteorology climate model in either a control or a transient run and suggest that this deficiency is not limited to this model alone.

These results confirm that current climate models are not accurately capturing the cloud and moisture processes active in the tropics—processes which play a critical role in the earth's temperature sensitivity. Furthermore, these model inaccuracies result in an overestimation of the true climate sensitivity.

2.3. Adjustment to the climate/carbon cycle feedback

The results of Luo et al. (2001) challenged the generally accepted notion that there exists a strong positive feedback between average global temperature and average biospheric respiration, primarily the amount of CO₂ released from the soil. They found that in tall grass prairie plots in the central United States, there was no elevation in soil respiration with an elevation in temperature. These results are supported by other long-term field studies (Verville et al. 1998, Liski et al. 1999). Since IPCC TAR incorporates a climate change/CO₂ feedback that includes a positive feedback between temperature and soil respiration (which IPCC SAR did not consider), the results of Luo et al. (2001) imply that the CO₂ emissions pathways in TAR are overestimated.

Soil respiration is not the only factor involved in the temperature/biospheric feedback loop, although it is a major factor (Cox et al. 2000, Luo et al. 2001). In MAGICC, the effect of incorporating a climate/CO₂ feedback is an increase of atmospheric CO₂ of about 4% by 2100 over the non-climate/CO₂ feedback version (the exact value depends on the particular SRES scenario). Therefore, we simulate the impact of the findings of Luo et al. (2001) by reducing the total climate/CO₂ positive feedback by 75%, an amount that allows for the continued possibility of a small positive feedback from soil respiration or other terrestrial and oceanic processes. We implement this reduction by developing a CO₂ emissions pathway that results in only a 1% increase in atmospheric CO₂ levels rather than the 4% that would be required to achieve a match between the SAR and the TAR emissions pathways in MAGICC (as described earlier in this section).

2.4. Results from adjustments to the radiative forcing, thermal sensitivity and climate/carbon cycle feedback

The results from running the 6 SRES illustrative scenarios and the IS92a scenario with the adjustments for the new determinations of the radiative forcing of BC aerosol, the temperature sensitivity, and the climate/CO₂ feedback are shown in Fig. 2. The range of projected warming for these scenarios drops from 2.0–4.5°C in the TAR to 1.1–2.8°C. As an example of

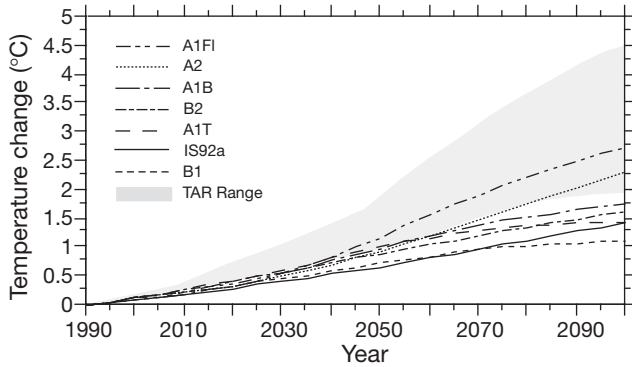


Fig. 2. 6 illustrative SRES scenarios and the IS92a scenario adjusted for an enhanced black carbon forcing (Jacobson 2001), a reduced temperature sensitivity (Lindzen et al. 2001), and a reduced climate/carbon cycle feedback (Luo et al. 2001). Gray: range of temperature projections from the IPCC Third Assessment Report

'central' scenarios, warming from 1990 to the year 2100 in IS92a drops from 2.6 to 1.5°C and TAR scenario A1B declines from 2.9 to 1.7°C.

3. ADJUSTMENT FOR OBSERVED TRENDS IN CO₂ EMISSIONS

As an alternative to the CO₂ emission pathways as prescribed by the IPCC TAR storylines and the time series of atmospheric CO₂ loadings that result (derived from the particular carbon cycle model employed), one could take the observed behavior of the atmospheric CO₂ concentrations as a total integrator of technological, biological, and geochemical processes that impact the atmospheric CO₂ burden. For most of the history of climate modeling, including a compendium of recent results detailed by Meehl et al. (2000), GCM projections assume a 1% yr⁻¹ increase in effective CO₂ concentration from background conditions. This is clearly not true, as it would produce a doubling in 70 yr, while the effective change in CO₂ concentration (counting the assumed forcing from all greenhouse emissions, including methane, chlorofluorocarbons, etc.) is usually assumed to be about 50 to 60% above the pre-industrial background (IPCC 2001).

The use of this wrong assumption results in overestimation of future warming. In addition, the non-CO₂ greenhouse enhancers are becoming increasingly inconsequential. Radia-

tively active chlorofluorocarbons have been effectively banned by the Montreal Protocol and atmospheric concentrations are beginning to decrease (Hansen & Sato 2001). Methane concentrations have leveled (Dlugokencky et al. 1998). CO₂ is the major greenhouse gas that continues to increase. These changes are reflected somewhat in the TAR SRES scenarios, but nevertheless, the majority of the SRES scenarios result in a rate of atmospheric CO₂ buildup that greatly exceeds observed trends.

The temporal behavior of the historical series of annual atmospheric CO₂ concentration as measured at Mauna Loa (Keeling & Whorf 2002) from 1958 to 2000 is best fit with a quadratic model including both a linear and a quadratic term. However, when successive years from the start of the record are withheld from the analysis, we find that the quadratic term becomes non-significant beginning in 1975 (as indicated by its p-value in the regression equation). This implies that the non-linear term contains no statistically significant information during the post-1974 time period and that the observed changes in CO₂ concentration in the last quarter-century are best fit by a simple linear model with a trend of 1.54 ppm yr⁻¹. Table 1 summarizes the results of the 2-parameter quadratic and 1-parameter linear models for the period 1975–2000. Fig. 3a graphically demonstrates the stabilization in growth rate (ppm yr⁻¹) of atmospheric CO₂.

This stabilization has occurred despite a continued increase in the annual rate of global CO₂ emissions (Fig. 3b; Marland et al. 2002). From 1975 to 1998, global CO₂ emissions increased by 42% yet the growth rate of atmospheric CO₂ burden remained steady. The reasons for a constant CO₂ growth rate in the face of steadily increasing emissions are not fully understood

Table 1. Regression results from the 2-parameter quadratic and 1-parameter linear models fit to the observed atmospheric CO₂ concentration for the period 1975–2000. SC: standardized coefficient

	Count	R	R ²	p	Regression SS	Residual SS
Quadratic	25	0.999	0.998	<0.0001	3090.78	6.185
Linear	25	0.999	0.998	<0.0001	3090.38	6.584
Regression		Coeff.	SE	SC	t	p
Quadratic						
Intercept		-2574.5	121.2	-2574.5	-21.24	<0.0001
Linear		1.47	0.06	0.95	24.03	<0.0001
Quadratic		0.0027	0.002	0.047	1.19	0.2463
Linear						
Intercept		-2714.6	29.5	-2714.6	-92.02	<0.0001
Linear		1.54	0.015	0.999	103.9	<0.0001

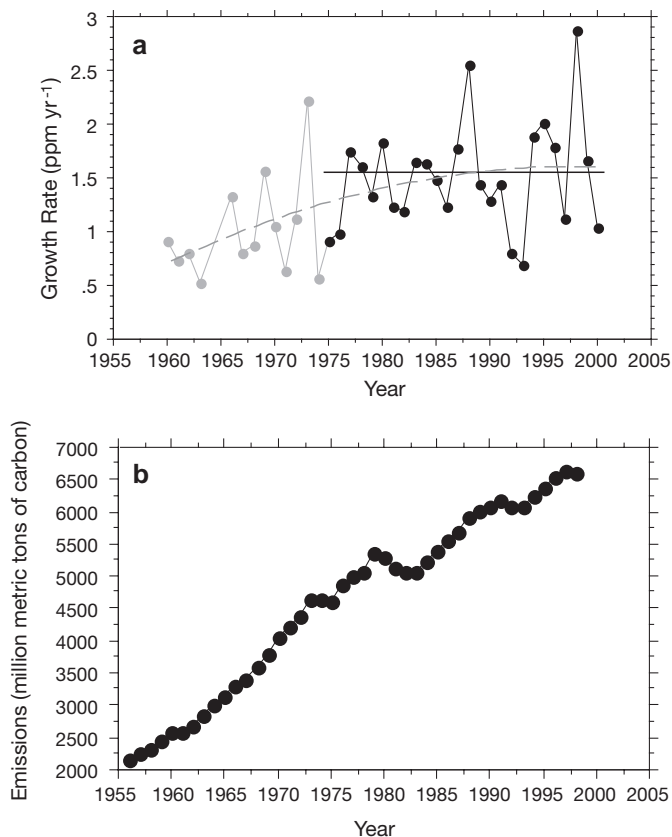


Fig. 3. (a) Annual atmospheric CO₂ growth rates (1958–2000); a quadratic fit to the data (gray dotted curve) and a linear fit since 1975 (solid black line) both indicate that the current atmospheric growth rate has levelled off at 1.5 to 1.6 ppm yr⁻¹. (b) Annual global CO₂ emissions (1956–1998)

but require enhanced terrestrial and/or oceanic CO₂ uptake. While there exist many carbon cycle models that attempt to explain this behavior, all include varying degrees of uncertainties surrounding every aspect of the models including, but not limited to, the effects of climate interactions, CO₂ fertilization, land use alterations, and marine productivity and chemistry changes. The observations, however, include no such uncertainty. They are the perfect integrators of all processes that are currently active. While there have been numerous suggestions that the flat CO₂ growth rate of the past 25 yr cannot be maintained (e.g. IPCC 2001, Hansen 2001), there is no indication that the current trend is breaking down. Therefore, an extrapolation of the observed trend should represent a reasonable, if not preferable, scenario of the future CO₂ concentrations.

To investigate the effect of a continued linear rise in atmospheric CO₂ concentration on projected temperatures, we examined this alternative to the SRES CO₂ buildup in the modified MAGICC SAR model. To do

this, we developed a CO₂ emission scenario that results in an atmospheric growth curve that matches the form of the observations; that is, an extension of the observed linear trend during the past 25 yr. By the year 2100, this produces a CO₂ concentration of 522 ppm. The range of atmospheric CO₂ concentrations that result from the 6 IPCC SRES illustrative scenarios coupled with the variations of the IPCC carbon cycle models is 541 to 970 ppm (Wigley 2000). Each of the 6 SRES illustrative scenarios and 33 of the full set of 35 SRES scenarios creates a greater concentration of atmospheric CO₂ than a simple extrapolation of reality does.

We replace the SRES CO₂ emissions pathways with the one that is based upon the extrapolation of observations. All other aspects of the SRES scenarios are left unchanged (e.g. SO₂ emissions, CH₄ emissions, etc.). We then perform 2 sets of MAGICC runs: the first with the adjusted SRES scenarios and the temperature sensitivity fitted to match the TAR results, and the second with the adjustments for BC and the iris effect. We no longer adjust for the climate/carbon cycle feedback, as we assume that it is implicit in the observed behavior.

Results are presented in Fig. 4, comparing the temperature projections based upon the original SRES scenarios with the ones with a linear CO₂ extrapolation from the last quarter-century. In the runs with the fitted SRES temperature sensitivities the range of warming by 2100 is 1.5 to 2.6°C, a reduction of warming from the IPCC's range of 25 to 42%. The runs with the adjusted temperature sensitivities produced a warming range for 2100 of 1.0 to 1.6°C, a reduction from the IPCC range of 50 to 64%.

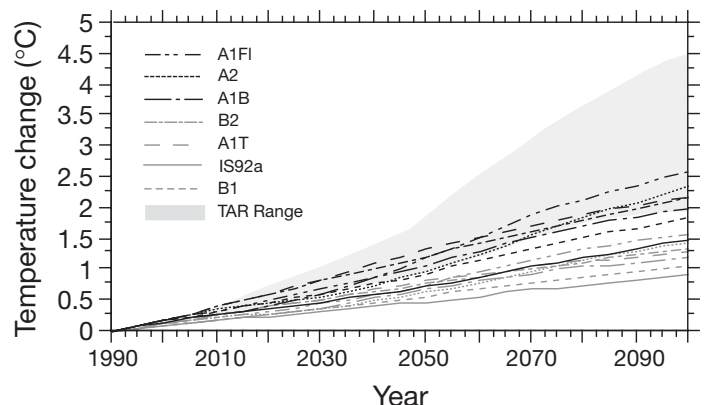


Fig. 4. Temperature projections for the 6 illustrative SRES scenarios and the IS92a scenario adjusted to include a linear extrapolation of observed atmospheric CO₂ growth rates. The black curves result from the use of the fit TAR values of temperature sensitivity, while the gray curves result from the use of the adjusted temperature sensitivities. Gray: range of temperature projections from the IPCC Third Assessment Report

4. ADJUSTMENT OF MODEL OUTPUT WITH OBSERVED TEMPERATURES

As shown by the IPCC TAR, the average of a large sample of climate models run with a $1\% \text{ yr}^{-1}$ increase in effective atmospheric CO_2 produces an essentially linear (rather than exponential) increase in temperature for the foreseeable future (Meehl et al. 2000). This largely results from the combination of a logarithmic response in temperature to CO_2 increases coupled with an assumed exponential growth, all adjusted implicitly for some thermal lag largely found in the ocean. It is noteworthy that the early parts of the Mauna Loa record (1957–1975) are much more strongly exponential than the last quarter-century. As a result, if there is a lag of a quarter-century or more, we should in fact be in a period of linear warming.

The IPCC SAR noted in 1995 that ‘the balance of evidence suggests a discernible human influence on global climate,’ and the TAR strengthened this statement by saying that ‘there is new and stronger evidence that most of the warming observed over the last 50 yr is attributable to human activities.’ Balling et al. (1998) and Michaels et al. (2000) demonstrated that the preponderance of warming is in cold, continental winter anticyclones. Calculations by Staley & Jurica (1970, 1972) on the overlap between CO_2 and water vapor imply that greenhouse warming largely would be concentrated in such air masses, and this is also implicit in many GCMs. Our hypothesis does not imply that the findings of Thompson & Wallace (2001) relating warming to temporal circulation variations are in question. Rather, circulation changes and direct greenhouse warming of surface anticyclones are not at all mutually inconsistent.

Consequently, the notion that greenhouse warming is the largest signal in recent surface temperature histories, coupled with the notion that the average of all GCMs is linear or very nearly so, leads to an obvious adjustment of the various slopes of model warming with the observed warming. This was first published by Michaels & Balling (2000) using a limited number of models, and subsequently a very similar calculation was published by Allen et al. (2000) based upon the Hadley Centre Model. Both concluded that warming in the next 50 yr is likely to be around 0.75°C . More recently, in an analysis which relied exclusively on observations, Hansen & Sato (2001) carefully examined trends in climate forcing agents, noting that the growth rate of greenhouse gas forcing agents has been declining for the past 20 yr. After reviewing the possible reasons for this decline, they state that ‘future global warming can be predicted much more accurately than is generally realized’ and that ‘practical constraints on changes in emission levels suggest that

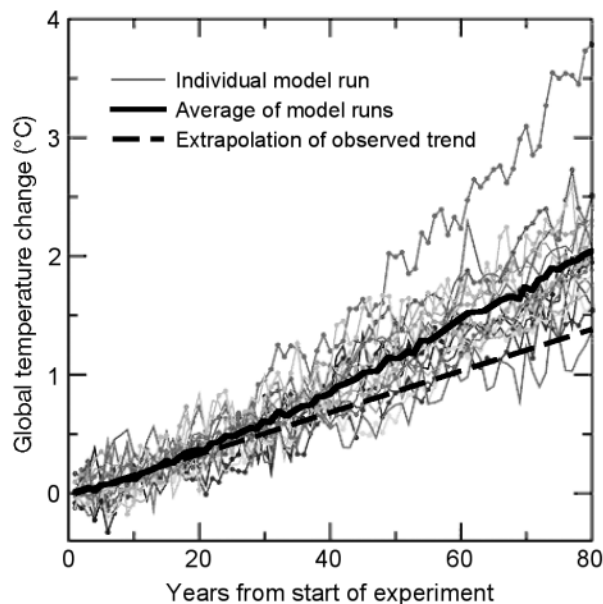


Fig. 5. Extrapolation of the observed temperature trend during the past 25 yr (thick dashed line) plotted with a collection of climate model output run with a $1\% \text{ yr}^{-1}$ effective CO_2 increase (thin gray curves). Thick solid line: mean of all models (adapted from IPCC 2001)

global warming at a rate of $+0.15 \pm 0.05^\circ\text{C}$ per decade will occur over the next several decades.’

Fig. 5 superimposes the observed warming of the last quarter-century upon IPCC TAR Fig. 9.3a, which includes projections from a large suite of models each run under the simple assumption of a $1\% \text{ yr}^{-1}$ increase of atmospheric CO_2 . If we project the warming rate of each model in Fig. 5 to encompass 11 decades of exponential CO_2 increase (so that it is comparable to the change from 1990–2100), the range of temperature increase is 1.9 to 4.5°C with a mean value of 2.75°C . The extrapolation of the warming trend of 0.17°C per decade that is characteristic of the past 25 yr produces a warming from 1990–2100 of 1.9°C or a value that is 68% of the mean of the individual model projections. Adjusting the high and low extreme model projections by this amount results in a range of 1.3 to 3.0°C .

5. DISCUSSION AND CONCLUSIONS

Our adjustments of the projected temperature trends for the 21st century all produce warming trends that cluster in the lower portion of the IPCC TAR range. Together, they result in a range of warming from 1990 to 2100 of 1.0 to 3.0°C , with a central value that averages 1.8°C across our analyses. These results are from somewhat independent analyses inasmuch as our first set of adjustments is largely a modification of climate

model input, while the second is an adjustment to climate model output. The first incorporates the observed behavior of the sensitivity (of both temperature and the biosphere), an adjustment for infrared radiation flux, and a simple integration of the planet's response to what the IPCC claims is responsible for 'most of the warming' in the last several decades. The second assumes that climate models have the form of the warming correct but generally overestimate its magnitude.

Our analysis differs from attempts to assign probabilities to future temperature projections in that we explore the effect of recent refinements to the understanding of key climate parameters while other studies (Reilly et al. 2001, Wigley & Raper 2001) examine the implication of the uncertainty range of climate parameters, a range that is guided by IPCC science rather than the recent refinements thereof, or the actual distribution statistics of the full set of IPCC projections (Schneider 2001).

The parameters that we examine in this study are only a subset of all those which have an impact on the earth's climate. The level of understanding of these parameters and the processes in which they are involved will continue to become more refined with time. In that light, our findings themselves will surely have to be modified as the results of new and ongoing research becomes available. Future results may serve to further reduce the range of expected 21st century temperature changes, or, in fact they may serve to broaden the range.

Despite the uncertainty that future results may hold, the set of climate observations continues to expand and the trends in fundamental quantities such as atmospheric CO₂ concentrations and global temperatures are becoming better established. Since these quantities serve as integrators of all processes acting on and within these systems, the only uncertainty they contain is measurement uncertainty, which is arguably small. Therefore, these trends should serve as the de facto standard for future expectations, at least in the near term and should be better incorporated into longer-term projections of future temperature changes.

These trends suggest that even the temperature range and central values determined in our study may be too great. Observations of atmospheric CO₂ buildup and global CO₂ emissions demonstrate that the family of IPCC 'storylines' that lead to exponential warming, characterized by A1FI in Fig. 1, is not based upon the reality of recent decades. The 'worst case' warming now appears to be merely linear, subject to the modifications described in this paper. Furthermore, both Table 1 and Fig. 3 indicate that any exponential rise in atmospheric CO₂ concentrations is weak at best. Con-

sequently, the current linear warming may in fact be the adjustment to the exponential growth in CO₂ that took place *prior* to 1975. Levitus et al. (2000) documented a warming of 0.06°C in the top 3 km of a large-area ocean sample over the course of 40 yr. A lag correlation between that deep-water record and the sea-surface temperature record from Quayle et al. (1999) is very suggestive that oceanic thermal lag maximizes around 35 yr (Michaels et al. 2001). Thus, the truly exponential phase of concentration growth in the atmosphere, which ended about 25 yr ago, should induce a linear warming for the next decade or two before it could actually begin to damp. Therefore, if the impact of the mid-century exponential CO₂ rise has nearly worked its way through the oceanic thermal lag, our warming projections derived from the assumption of a continued linear buildup of the atmospheric CO₂ burden should appear as the upper end of the likely range of warming during this century.

LITERATURE CITED

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