

# Penetration of ultraviolet radiation into shallow water sediments: high exposure for photosynthetic communities

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**ABSTRACT:** Benthic photosynthetic microorganisms are widespread in shallow-water sediments, microenvironments that are commonly assumed to be virtually opaque to ultraviolet radiation (UVR). We used a newly developed optical microprobe to measure the submillimeter penetration of solar UVR into a variety of these microenvironments. UVR trapping due to strong scattering occurred at the surface of some sediments, resulting in a surface maximum of scalar irradiance ( $E_0$ ) that could be significantly larger than the incident radiation. In the subsurface,  $E_0$  was typically extinguished in a quasi-exponential manner, with attenuation coefficients (310 nm) ranging from 4 to 21  $\text{mm}^{-1}$ , depending on sediment type. Ultraviolet B (at 310 nm) was extinguished to 1% of the incident between 1.25 and 0.23 mm from the surface. Within the euphotic zones of these sediments, however, the space-averaged UVB scalar irradiance was very high, between 15 and 33% of the incident. In natural waters, for example, the same parameter varies between 3 and 9% of the incident. Thus, in fact, photosynthesis in these environments must develop under strong UV stress, and it must be regarded as potentially labile to the effects of ozone depletion.

**KEY WORDS:** Sediments · Ultraviolet radiation · Primary productivity

## INTRODUCTION

Shallow-water and intertidal sediments, both fresh water and marine, are usually colonized by phototrophic microorganisms. Since visible light is strongly attenuated in these microenvironments, photosynthesis can occur typically only within millimeters from the surface (Jørgensen & Des Marais 1988, Lassen et al. 1992, Kühl & Jørgensen 1994). In spite of this, the benthic phototrophic microbial communities serve a crucial ecological function, contributing significantly to local productivity (Van Raalte et al. 1976, Pinckney & Zingmark 1993), and decreasing sediment erosion (Grant & Gust 1987, Patterson 1994). Solar ultraviolet B radiation (UVB; 280 to 315 nm) can in principle damage the photosynthetic apparatus of microalgae and cyanobacteria, but the actual impact of UVB on

natural communities is clearly, among other variables, a function of *in situ* exposure. It is commonly assumed that ultraviolet radiation (UVR) does not penetrate significantly into sedimentary microenvironments, which have in fact been regarded as an effective refuge from UV exposure (Margulis et al. 1976). However, sediment microalgae may produce large quantities of UV-sunscreen compounds (Garcia-Pichel & Castenholz 1991, 1993) and, at least in some of these communities, several biological processes can be affected by natural UVB exposure (Bothwell et al. 1994, Garcia-Pichel & Castenholz 1994). These observations cannot be easily explained if exposure occurs only at the virtual surface, but no measurements of the penetration of UV into sedimentary environments were available. For this reason, we recently developed a fiber optic microprobe which made the measurement of UV scalar irradiance ( $E_0$ ) possible at a spatial resolution of ca 0.1 mm (Garcia-Pichel 1995), and used it to measure the UV radiative fields within shallow-water sediments on which we report here.

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## MATERIAL AND METHODS

**Samples.** All measurements were performed in the laboratory with freshly collected samples. North Sea samples were collected using PVC corers and transported within several hours to the laboratory. Overseas samples were air-lifted and reached the laboratory within 2 to 3 d of collection. Samples were incubated in water from the site, or in artificial seawater made up to the local salinity conditions, at room temperature in small acrylic flumes. Samples were illuminated from above at 200  $\mu\text{mol photon m}^{-2} \text{ s}^{-1}$  of PAR (photosynthetically active radiation, by convention 400 to 700 nm), until measurement. Measurements were made on core subsamples (5 to 7 cm in diameter, 3 to 10 cm deep). A brief description of each sample follows: Sample 1: fine beach quartz sand, washed and with a grain size of around 250  $\mu\text{m}$ ; Sample 2: natural intertidal carbonate sand crust, harboring a low-density community of diatoms and cyanobacteria (Bahamas); Sample 3: upper intertidal sandy cyanobacterial mat, *Sandwatt*, from Norderney Island (German Wadden Sea); Sample 4: shallow-water sediment from the Bay of Aarhus (Denmark), composed of benthic diatoms in a matrix of allochthonous opal and organic detritus; Sample 5: hypersaline, virtually organic, cyanobacterial mat from Solar Lake, Sinai, Egypt; Sample 6: silty intertidal mud, *Schlickwatt*, from the Wadden Sea near Dorum (Germany).

**Optical measurements.** Core samples were illuminated from above with a 200 W Xe arc-lamp illuminator (Oriel Corp., Connecticut, USA). The lamp beam was collimated through a quartz condenser and reflected on a UV dichroic mirror so that the exiting beam (ca 5 cm in diameter) was collimated and highly enriched in UVB wavelengths. This set-up resulted in minimal effects of stray light signals in our measurement system (see 'Results') since, for any given spectrum, signals normally fell within 1 order of magnitude across the wavelength range 270 to 700 nm, even for spectra measured deep into the sediment.

Measurements of penetration of UVR and visible light were carried out with custom-built scalar irradiance microprobes with a sensing tip 80 to 105  $\mu\text{m}$  in diameter. The microprobes consisted of a tapered optical fiber with a UV-transparent vitroceraic tip diffuser, which was inserted into the samples at a 45° angle to the vertical using micromanipulators. The construction and use of the microprobes was described in detail elsewhere (Garcia-Pichel 1995). All operations were carried out under visual monitoring through a dissecting microscope focused on the sediment surface. Light collected by the microprobe was transmitted through the optical fiber, collimated at the opposite end and focused onto an IL 1785 (Internation-

Light 1785) spectroradiometer, which we modified to couple the fiber optic input. For the measurement of UVB profiles with depth, the spectroradiometer was set at 310 nm. For spectral measurements, we used a step size of 5 nm in the UV and of either 5 or 10 nm in the visible. The absence of stray light signals was checked using long-pass optical filters in the light source and spectrophotometer. The incident (beam) irradiance was determined by exchanging the subject with a black, nonreflecting well, maintaining the microprobe in place with respect to the beam. All data obtained are expressed here as a percentage of this parameter.

**Optical parameters.** The probes used measured scalar irradiance,  $E_0$ , or the radiation incident on a point in space from all directions in the solid angle, much like a typical  $4\pi$  sensor would. In highly scattering systems, such as sediments, it is important to measure this parameter because incident radiation becomes rapidly diffuse.  $E_0$ , also known as fluence rate or space irradiance, probably approximates the radiation to which the microorganisms are exposed better than other optical parameters, such as downwelling irradiance (see Kirk 1983, Kühl & Jørgensen 1994).

Subsurface attenuation coefficients of the scalar irradiance in the UVB,  $K_D(\text{UVB})$ , were calculated directly from the profiles (310 nm) using data below 100  $\mu\text{m}$  depth only, to avoid the surface light-trapping zones. For this we used linear regression of the ln-transformed mean values at each depth. The correlation coefficients were >0.98 in all cases. The depth at which incident UVB was attenuated to 1% of the incident,  $Z_{1\%}(\text{UVB})$ , was determined by interpolation in the regression line. The depth of the euphotic zone,  $Z_1(\text{vis})$ , was calculated from spectral measurements within the visible range (400 to 700 nm). For this, the spectral measurements were integrated to obtain attenuation profiles of visible light assigning equal weights to all wavelengths. From the averaged profiles, averaged  $K_D(\text{vis})$  and  $Z_{1\%}(\text{vis})$  were obtained in the same way as for the UVB profiles above. We note that the latter 2 parameters, because they average broadband radiation, are actually a function of the spectral composition of the incident light, and thus may vary somewhat under different spectral conditions.

As a means to compare the overall exposure for the phototrophic organisms, we calculated the space-averaged mean UVB scalar irradiance within the euphotic zone,  $\bar{E}_0^{\text{eu}}(\text{UVB})$ , which is defined as

$$\bar{E}_0^{\text{eu}}(\text{UVB}) = \frac{1}{Z_{1\%}(\text{vis})} \int_0^{Z_1(\text{vis})} E_{0,310}(\text{UVB})(z) dz$$

where  $E_{0,310}(\text{UVB})(z)$  is the distribution of the UVB scalar irradiance with depth and  $z$  is depth, and was calculated as

$$\bar{E}_0^{\text{eu}}(\text{UVB}) = \frac{1}{n} \sum_1^n E_0(\text{UVB})(z)$$

where  $E_0^{\text{eu}}(\text{UVB})(z)$  is each of the measured values of  $E_0$  in the UVB profile that fell within the euphotic zone of the sediment, and  $n$  is the total number of such values.

Because the attenuation of UVR was in all cases stronger than that in the visible, the ratio of UV to visible decreased with depth. This may have implications for the *in situ* photobiology of the microorganisms, since a progressive 'refuge' from ultraviolet exposure may be gained by avoiding the surface zones if enough visible light for photosynthesis is available at depth. The dimensionless parameter,  $R$ , was designed as a means to compare the depth-refuge effect in different environments. It is defined, and calculated, as

$$R = [K_D(\text{UVB}) - K_D(\text{vis})] [K_D(\text{vis})]^{-1}$$

It takes positive values whenever the attenuation of UVB is stronger than that of visible, and it becomes larger as the difference in attenuation coefficients between UVB and visible grows. By normalizing to the attenuation coefficient of the visible, the parameter becomes dimensionless and environments of different (optical) depths can be readily compared. Environments with a high value of  $R$  will offer a significant depth-refuge from UV exposure, and vice versa. We note that, because the exponential extinction of  $E_0$  occurs only at the subsurface, this parameter cannot satisfactorily describe the zones where light trapping occurs.

## RESULTS

Depth profiles for  $E_0$  in the UVB, in samples ranging from organic-free mineral to virtually organic, are pre-

sented in Fig. 1. The optical parameters are presented in Table 1. Incident UVB was strongly attenuated in all samples (100-fold within 1.3 mm or less). The build-up of a surface  $E_0$  maximum, a light-trapping effect due to scattering, typical in strongly scattering media (Vogelmann & Björn 1986, Jørgensen & Des Marais 1988), and the apparently exponential extinction of  $E_0$  ( $R^2$  always  $>0.98$  for ln-transformed mean data) below 100 to 200  $\mu\text{m}$ , defined the shape of the profiles. Surface maxima [ $E_0(\text{UVB}), z = 0$ ] ranged from insignificant in silty and highly organic samples to values 157% higher than the incident irradiance in carbonate sand crusts. Subsurface exponential extinction occurred with apparent diffuse attenuation coefficients ranging from 4.1 in wet, clean quartz sand to 21.6  $\text{mm}^{-1}$  in silty mud (Table 1). The values of  $\bar{E}_0^{\text{eu}}(\text{UVB})$  calculated for these sedimentary microenvironments ranged between 15% of the incident irradiance in wet sand to 25% of the incident irradiance in carbonate sand and 33% of the incident in silty mud; these values were not well correlated with their respective UVB attenuation coefficients.  $\bar{E}_0^{\text{eu}}(\text{UVB})$  values attained in sediments were invariably much higher than those calculated for natural waters (also presented in Table 1), which varied between 3 and 9% of incident for a range of optically diverse localities.

Spectral determinations showed (Fig. 2) that there was a general marked trend of increased attenuation with decreasing wavelength. Whereas this occurred in all cases (not all data shown), spectral signatures for photosynthetic pigments (particularly chlorophyll *a*) were sometimes clear wherever microalgal populations were present at sufficient densities. Spectral signatures for mycosporine-like amino acid derivatives, thought to act as UV-sunscreen compounds, may have been responsible for the spectral signatures in the UVB

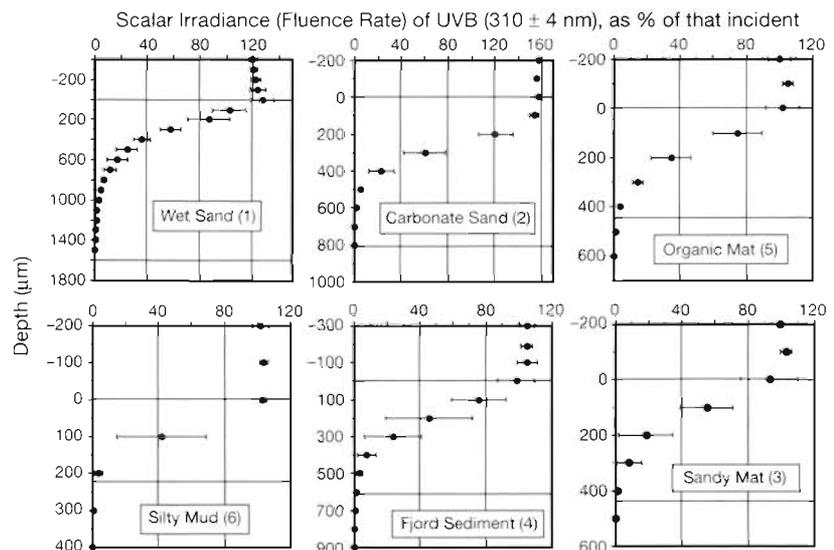


Fig. 1 Profiles of UVB (310 nm) scalar irradiance within sedimentary environments. Means and standard deviations of 4 to 5 profiles at different locations in each sample are shown. Inserted boxes identify each sample with a short name and the sample number (see 'Materials and methods' and Table 1 for further information). In each panel, the upper horizontal line represents the (approximate) surface and the lower line the depth at which the visible light (400 to 700 nm) was attenuated to 10% of the incident ( $\frac{1}{2}$  of the euphotic zone), which was determined separately from data such as those in Fig. 2

Table 1. Optical characteristics of sediments and waters. See 'Materials and methods' for sample details and parameter definitions. For waters, the optical parameter measured was downwelling irradiance, not scalar irradiance; in natural waters scalar irradiance was only slightly larger than downwelling irradiance

Characteristics	$Z_{1\%}(\text{UVB})$ (mm)	$Z_{1\%}(\text{vis})$ (mm)	$\bar{E}_0^{\text{eu}}(\text{UVB})$ (% incident)	$E_0(\text{UVB}; z = 0)$ (% incident)	$K_D(\text{UVB})$ ( $\text{mm}^{-1}$ )	$R$
<b>Sediments</b>						
1. Fine beach sand (wet)	1.25	3.10	15	127	4.1	1.56
(dry)	0.98	2.40	23	131	6.5	1.62
2. Carbonate sand crust	0.64	2.41	22	157	12.1	3.97
3. Sandy intertidal mat	0.36	0.72	24	103	17.2	1.28
4. Diatom/opal fjord sediment	0.65	1.23	18	98	7.9	0.88
5. Cyanobacterial mat	0.50	0.95	25	105	10.5	0.75
6. Silty intertidal mud	0.23	0.45	33	103	21.6	0.94
<b>Waters</b>						
Sargasso Sea <sup>a</sup>	$25 \times 10^3$	$150 \times 10^3$	6	–	$0.12 \times 10^{-3}$	2.8
Southern Ocean <sup>b</sup>	$17 \times 10^3$	$120 \times 10^3$	3	–	$0.26 \times 10^{-3}$	5.9
	$16 \times 10^3$	$63 \times 10^3$	6	–	$0.28 \times 10^{-3}$	2.8
Admiralty Bay (Antarctica) <sup>c</sup>	$22 \times 10^3$	$44 \times 10^3$	6	–	$0.26 \times 10^{-3}$	1.0
Dutch Wadden Sea <sup>c</sup>	$1 \times 10^3$	$3 \times 10^3$	9	–	$4.60 \times 10^{-3}$	2.5

<sup>a</sup>Calculated from spectral attenuation coefficients (Smith & Baker 1981)  
<sup>b</sup>Data from Helbling et al. (1994) gathered at 2 different stations near Elephant Island.  $Z_{1\%}(\text{vis})$  is from measured values  
<sup>c</sup>Data from Vosjan & Pauptit (1992). Measurements of 'visible' included infrared wavelengths, which are strongly absorbed in water, so  $Z_{1\%}(\text{vis})$  is probably underestimated and, consequently,  $\bar{E}_0^{\text{eu}}(\text{UVB})$  and  $R$  are probably overestimated.  $Z_{1\%}(\text{vis})$  is from measured values

region that we found in densely populated cyanobacterial mats (Sample 3), and thus such sunscreens may have contributed significantly to the extinction of UVB in these systems.

The parameter  $R$ , which is high in those environments in which UVB is strongly attenuated relative to

visible and which gauged the depth-refuge from UVB exposure, varied around 1 in most sedimentary samples, indicating that UVB was attenuated approximately twice as strongly as the averaged visible. The carbonate sand crust had a much higher value, and UVB was attenuated in the subsurface close to 5 times as strongly as the visible. The depth-refuge effect tended to be larger in waters than in sediments; it was, however, very much environment-specific. Interestingly, the lowest  $R$  values were found in highly organic samples (i.e. cyanobacterial mats), which may have had to do with the increased attenuation of visible light due to their high concentrations of photosynthetic pigments.

## DISCUSSION

The range of sediments studied were diverse with respect to the attenuation of UVR, from silty mud sediments being the most absorbant (Sample 6) and quartz sand the least (Sample 1), paralleling their optical characteristics in the visible. Carbonate sand sediments seemed to behave rather differently than sandy, silty and organic sediments, in that both light trapping and differential extinction of UV with respect to visible were very strong. This may have

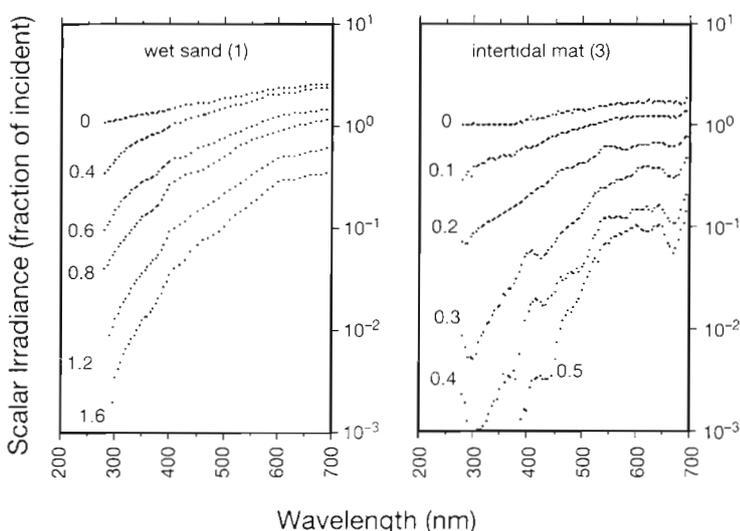


Fig. 2.  $E_0(\lambda)$  at selected depths in sedimentary benthic environments. The depth corresponding to each spectrum is noted at the left end. In sand (Sample 1), a rather featureless increase of attenuation with decreasing wavelength occurred. In the sandy mat (Sample 3), the spectral signatures for photosynthetic pigments are superimposed on this general increase. A signature for mycosporine-like amino acids may be responsible for the maximum in attenuation at ca 300 to 310 nm in the intertidal mat

been due to the fact that each carbonate particle scatters light through diffuse surface reflectance (each particle has a lot of microstructure and many refraction events may occur within each one), rather than through multiple refraction, as it might occur with living cells or quartz particles. Thus, carbonate sand sediments may offer a stronger refuge from UVB exposure in the deeper layers of the euphotic zone than do sandy or organic sediments.

In clear waters, absorptive losses dominate the attenuation of radiation, and water absorption is significantly larger in the UV than in the visible. In sediments, however, both scattering and absorption processes contributed significantly to attenuation. The magnitude of scattering losses is manifest in the differences found between wet and dry sand. The difference in refractive index between particles and interstitial medium, largely responsible for the scattering of light, is larger in dry than in wet sand. Since scattering is strongest in dry sand, UVB penetrated less (Table 1). But because the radiation lost through scattering to the vertical path is not immediately lost to the system, but contributes to the (diffuse) fluxes at other nearby points, the average scalar irradiance within the system was actually higher in dry sand (Table 1). This phenomenon is responsible for the near-surface maxima, or light-trapping effect, which occurred in some sediments. These effects were considerably weaker in the UVB than in the UVA, and even weaker than in the visible in the same samples (Fig 2), as well as in other sediments and mats (Lassen et al. 1992, Kühl & Jørgensen 1994), so that the ratios of UVB and UVA to PAR in the light-trapping surface zone, and, thus, at all depths, were smaller than those in the incident radiation. Additionally, the high concentration of photosynthetic pigments in some of these communities resulted in strong attenuation of visible light, in effect narrowing the euphotic zone. Both the near-surface UV light-trapping effects and the strong attenuation of visible light are probably responsible for the comparatively high levels of exposure measured in the euphotic zones of sedimentary microenvironments.

Studies have identified significant decreases of primary productivity attributable to present-day solar UVB on planktonic photosynthetic communities (Smith et al. 1992, Holm-Hansen et al. 1993). Provided that UV sufficiently penetrates the overlaying waters, one should expect a much more marked impact of UVB on the metabolically analogous communities living in sediments, based only on the optical characteristics of their microenvironments. Indeed, in irradiation experiments, we could detect UVB-induced photodamage to photosynthesis in several of these communities, both at the surface and in the subsurface (Bebout & Garcia-Pichel 1995). In some cases, however, we found signif-

icant damage deeper than it would be expected from the measured UVB penetration (see also Garcia-Pichel & Castenholz 1994). This may be the result of a combination of factors unique to sedimentary ecosystems: a near-surface zone containing UV-absorbing organic components, high oxygen tensions (due to localized, high photosynthetic rates) and exposure to UV, which constitutes the right mixture for photosensitized and radical-mediated photooxidative reactions. The detrimental effects of UV may thus be locally enhanced and certain long-lived reactive oxygen species (singlet oxygen, hydrogen peroxide) may diffuse down where little UV exposure exists, producing some additional indirect damage.

In summary, the exposure of microphytobenthic communities to ultraviolet radiation does not only depend on the optical transparency of the overlaying waters but also on the optical properties of the sediments themselves. The optics of the microbial photosynthetic communities we examined allowed UV to play a significant role in their functioning, not only in the case of bare substrates (hence in the colonization process), but even more so in well-formed, highly organic communities, such as microbial mats, where the impact of UV may be enhanced by synergistic effects of high concentration of photosensitizers and elevated oxygen tensions. Our results underscore the need, recently expressed by other investigators (Vincent & Quesada 1994), to give serious consideration to benthic communities with regard to the ecological implications of increases of UVB due to stratospheric ozone depletion. Field experimentation is needed to establish the sensitivity of these communities to UVR. The extent to which biological adaptation to the physicochemical characteristics described above may prevent or counteract the detrimental role of UV in these communities remains largely to be explored. The use of sunscreen compounds and the regulation of organismal exposure through vertical migrations seem to be important, but not universal, adaptations to avoid excessive exposure.

*Acknowledgements.* We thank S. Neuer and L. Prufert for critically reading the manuscript and M. Kuhl, L. Prufert, L. Stal, and B. Thamdrup for samples. The authors were supported by stipends from the Max-Planck Society.

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*This article was submitted to the editor*

*Manuscript first received: May 24, 1995*

*Revised version accepted: July 18, 1995*