

Seasonal variation of microalgal assemblages at a fixed station in King George Island, Antarctica, 1996

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ABSTRACT: Microalgal assemblages were measured daily from January to December 1996 at a fixed neritic station in Marian Cove, King George Island, Antarctica. The abundance of microalgae and carbon biomass exhibited clear seasonal variation. Annual mean of total microalgal abundance in surface water was 2.43×10^4 cells l^{-1} . Microalgal cell abundance in Marian Cove showed a multimodal distribution of standing crop during the study period. Microalgae started to bloom in October and increased abruptly during November. More than 45% (avg. 3.5 mg m^{-3}) of chl *a* was present in the 2 months November and December, dominated by microplanktonic diatoms ($>20 \mu\text{m}$) such as *Fragilaria striatula* Lyngbye, *Licmophora belgicae* Peragallo, and *Achnanthes groenlandica* Grunow. The increase of these diatoms were mainly due to resuspension of benthic microalgae by wind and tidal currents in spring and summer. In contrast, microalgal assemblages in winter were characterized by the dominance of pico- and nanoplanktonic microalgae ($<20 \mu\text{m}$) such as *Phaeocystis antarctica* Karsten, *Navicula glaciei* Van Heurck, and *Navicula perminuta* Grunow.

KEY WORDS: Microalgal abundance · Carbon biomass · Diatom · Picoplanktonic microalgae · Nanoplanktonic microalgae · Marian Cove · King George Island

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INTRODUCTION

Microalgae in Antarctic coastal environments have been studied mainly during austral summers (e.g. Dayton et al. 1986, Perrin et al. 1987, Brandini & Rebello 1994), because micro- and nanoplanktonic microalgal assemblages change considerably during austral summer. The changes are probably a cause of influences such as local wind and tide (Brandini & Rebello 1994). Although changes of microalgal assemblages in Antarctic coastal environments occur during all 4 seasons, seasonal studies at the species level, such as species abundance, species composition, and succession of phytoplankton in coastal regions, are practically nonexistent (Horne et al. 1969, Krebs 1983). Most studies

have analyzed chlorophyll *a* (chl *a*) concentration, because analysis at the species level requires a long time period, and taxonomy of neritic species is poorly known. To date, continuous year-round monitoring has been undertaken at only 1 Antarctic station (Clarke et al. 1988). Data for single complete years are available for a few other stations (e.g. Fukuchi et al. 1985, Tokarczyk 1986, Domanov & Lipski 1990, Kang & Kang 1997, Kang et al. 1997). Between 1969 and 1983, seawater temperature, macronutrients, and chlorophyll were measured regularly at a coastal site close to the British Antarctic Survey research station on Signy Island, South Orkney Islands (Clarke et al. 1988). In this early study, distinction between different size-fractions was not made. Clarke & Leakey (1996) measured seawater chl *a* concentration (size-fractionated at 20, 2, and $0.2 \mu\text{m}$) and temperature at Signy Island weekly from 1988 to 1994.

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There are 3 types of neritic microalgae in nearshore environments of Antarctica: phytoplankton, sea-ice microalgae, and benthic microalgae. The phytoplankton have been well examined in both offshore and inshore regions (Horne et al. 1969, Whitaker 1982, El-Sayed 1984, Fiala & Oriol 1990). Sea-ice microalgae have also received much attention. Productivity measurements of the sea-ice algae indicate that they are an important source of organic carbon at their peak activity (Palmisano & Sullivan 1983, McConville et al. 1985, Grossi et al. 1987, Palmisano et al. 1988, Kang et al. 1999). Benthic microalgae such as sediment-living

diatoms, epiphytic, and epilithic forms are also common in nearshore environments (Heywood & Whitaker 1984). There have been a few studies on microphyto-benthic algae, but investigations in the Antarctic and Arctic regions suggest that these microalgae may be important in primary production at certain times of the year (Horner & Schrader 1982, Dayton et al. 1986,). Suspended benthic diatoms may contribute significantly to total primary production (Picken 1985, Klöser et al. 1994, Ahn et al. 1997), and benthic diatoms are assumed to be an important food source for benthic animals (Berkman et al. 1986, Ligowski 1983). Temporal variations in abundance of benthic diatoms are thus expected to have a significant influence on ecosystems of Antarctic nearshore environments. However, distributional patterns of benthic diatoms have been rarely studied.

Maxwell Bay shows open-water characteristics, as the bay mouth is open to Bransfield Strait and also has the characteristic of nearshore ecosystems. Therefore, microalgal populations comprise both benthic and pelagic species. In the present study, abundance, biomass, and species composition of micro- and nanoplanktonic microalgae were monitored at a fixed station in Marian Cove, King George Island (Fig. 1). The objectives of this study were (1) to observe the dominant species in the area, including nanoplanktonic microalgae; (2) to identify the major carbon source at coastal site; and (3) to determine microalgal succession during the annual cycle.

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MATERIALS AND METHODS

Samples were collected at a fixed station in front of King Sejong Station from 11 January to 31 December 1996 ($n = 273$) (Fig. 1). At this station the water is about 10 m deep and the water column is well mixed by wind and tidal currents. Water samples were taken at a depth of 0.5 m with a PVC bottle, and water temperature and salinity were measured with Sea-Bird 911 plus CTD (Otronix).

Water samples were transferred to the research station immediately after collection and were used to measure chl *a* concentration with Turner Design field fluorometer (Model 10-005R). The *in vivo* fluorescence values were calibrated with

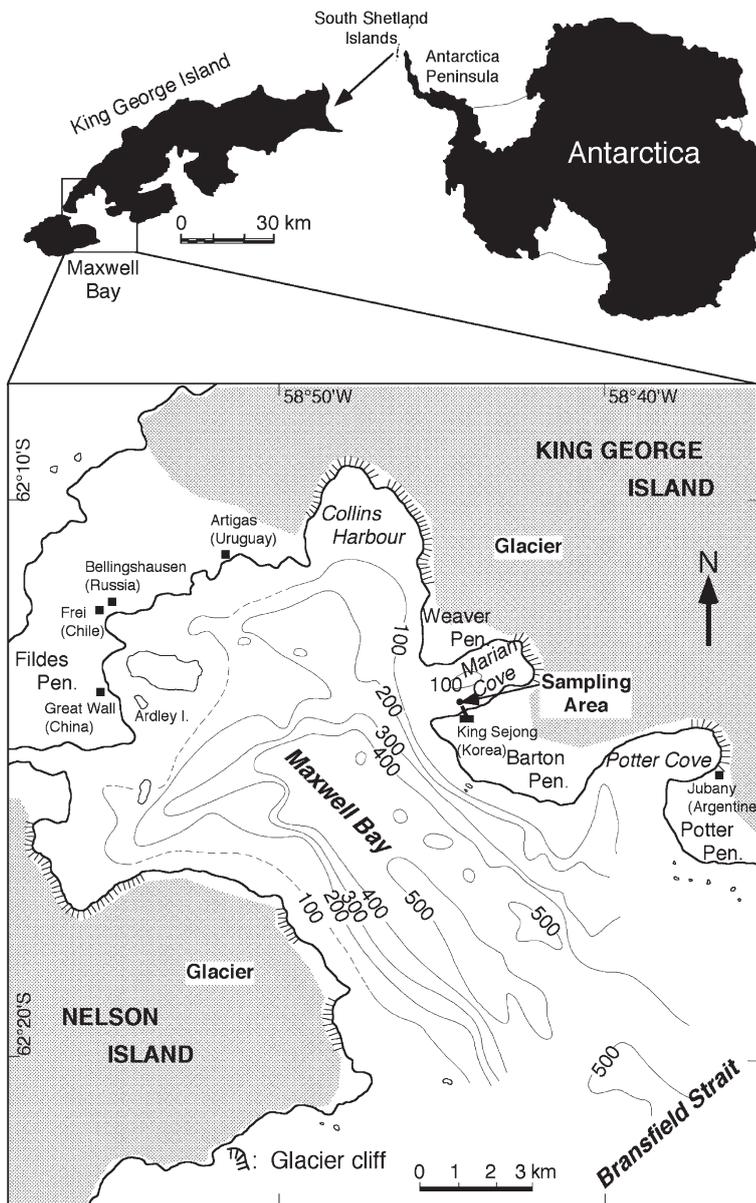


Fig. 1. Location of the sampling area (●) in Marian Cove, King George Island, Antarctica

absolute chl *a* values. The chl *a* concentration was estimated spectrophotometrically (Parsons 1984). Particulates were collected on 25 mm glass fiber filters (GF/F) and the filters were then left in the dark for 12 h in 90% acetone at 4°C for pigment extraction. The linear relation ($y = 0.8x - 0.11$, $n = 55$) between *in vivo* and absolute values was used to measure chl *a* concentrations (mg m^{-3}). Size-fractionated study of chl *a* was conducted with 20 μm nylon mesh. Cells which passed through the net were considered as pico- and nanoplankton chlorophyll (<20 μm) and consisted mostly of flagellates, small diatoms, cyanobacteria, and cell fragments. Microplankton chlorophyll (>20 μm) consisted essentially of colonial forms, chain-forming species and large single diatoms.

For nutrients (nitrate, phosphate and silicate), water samples (500 ml) were filtered through a 47 mm GF/F and the filtrates were stored frozen in acid-cleaned polyethylene bottles at -45°C. The samples were kept frozen with dry ice during transport to the Korea Ocean Research & Development Institute (KORDI). Later in the lab, nutrient concentrations were determined with a Lachat Autoanalyzer. For suspended solids (SS), 1 l of seawater was filtered on a pre-dried and weighed GF/F, and the filter was dried in a 60°C drying oven for 24 h and weighed. Total weight of suspended solids was calculated as the value of the filtered GF/F after drying minus the pre-dried and weighed GF/F.

Aliquots of 125 ml were preserved with glutaraldehyde (final concentration 1%) for microscopic analysis of micro- and nano-sized algae in the laboratory. Samples were filtered and mounted on slide glass with 2-hydroxypropyl methacrylate (HPMA). The HPMA slide-counting method was used for quantitative analysis on cell concentration and biomass (Kang 1993). At least 30 fields or 300 cells were enumerated using a Zeiss Axiophot microscope with a combination of light and epifluorescence microscopy at 400 \times for microplankton (cells >20 μm), and at 1000 \times for autotrophic pico- and nanoplankton (Booth 1993). The micro- and nanoplanktonic microalgae were identified with a scanning electron microscope (SEM; Philips 515).

Conversions from cell count to cell carbon biomass were carried out as described previously by Kang et al. (1993). Cell dimensions of dominating microalgal species were measured to the nearest 1 μm for subsequent estimations of biovolume using appropriate geometrical shape (Kang et al. 2001). The carbon biomass was estimated from the cell biovolume with the modified Strathmann equations (Smayda 1978).

The computer program Systat was used for the statistical treatment of the microalgal cell abundance and environmental data. Statistical analysis yielded Pear-

son's correlation coefficient (r), which is the linear association between physicochemical variables and phytoplankton species that are normally distributed.

RESULTS

Environments

Maxwell Bay (62°10–19'S, 58°35–58'W) is surrounded by King George Island and Nelson Island (Fig. 1). The mouth of Maxwell Bay is open to the southeast and water mass in the bay is influenced by the Bransfield Strait (Chang et al. 1990). The bay is about 18 km long and 6 to 14 km wide. Water depth gently increases from the coastline to 200 m, but sharply increases from 200 to 400 m. The central part of the bay is relatively flat, ranging from 400 to 500 m in depth.

Marian Cove is a tributary inlet of Maxwell Bay and one of the big fjords at King George Island. The bottom of Marian Cove and the adjacent open coasts consist of hard substrata (boulders, stones, pebbles, gravel, and sand) with macroalgal vegetation of changing density (Chung et al. 1991, 1994). During winter, surface water is covered with sea-ice or icebergs. In summer, surface ice melts and occasionally drifting floes are carried into the bay by wind. Freshwater input also occurs from land and submerged glaciers in summer. The details of hydrographic features of the bay during summer have been described by Chang et al. (1990).

Seasonal characteristics of surface water at a fixed neritic station in Marian Cove

Table 1 shows the physicochemical data recorded during 1996 (Kang et al. 1997). The study period was a 'warm year' showing an annual mean air temperature of -0.7°C, which is about 1.3°C higher than the mean of the last 8 yr from 1988 to 1995 (avg. -2.0°C) measured at King Sejong Station (Kang et al. 1997). Annual mean temperature of surface water in Marian Cove was -0.28°C, about 0.4°C higher than air temperature (avg. -0.7°C). Annual mean salinity of seawater was 33.5 psu. Higher salinity (34.03 to 34.36 psu) was formed during austral winter (Table 1) and the lowest salinity was formed during February (avg. 31.97 psu), which is about 2.5 psu lower than the highest in July (34.36 psu). The lowest salinity was due to high runoff of snow and glacial meltwater, because highest air temperature was observed in February. Also, the high meltwater in February might be the cause of low concentrations of nitrate, phosphate, and silicate and increase of suspended solids (SS; Table 1). Dilution

Table 1. Monthly mean values of physicochemical variables (air temperature, water temperature, salinity, total chl *a* concentration, nanoplanktonic microalgal chl *a* concentration [$<20 \mu\text{m}$], microplanktonic microalgal chl *a* concentration [$\geq 20 \mu\text{m}$], suspended solids [SS], nitrate, phosphate and silicate) recorded during 1996 (STD = standard deviation, n = sample size, temp. = temperature)

	Air temp. (°C)	Water temp. (°C)	Salinity (psu)	Total chl <i>a</i> (mg m ⁻³)	Nano chl <i>a</i> (mg m ⁻³)	Micro chl <i>a</i> (mg m ⁻³)	SS (mg l ⁻¹)	Nitrate (μM)	Phosphate (μM)	Silicate (μM)
January STD (n = 21)	2.01 (1.07)	1.33 (0.72)	32.60 (0.79)	1.74 (0.96)	1.42 (0.73)	0.32 (0.49)	29.04 (1.69)	17.57 (4.04)	1.33 (0.19)	51.74 (9.11)
February STD (n = 29)	2.72 (1.37)	1.23 (0.39)	31.97 (0.60)	1.60 (0.86)	1.28 (0.76)	0.32 (0.21)	29.53 (2.47)	16.30 (3.99)	1.23 (0.17)	45.27 (14.60)
March STD (n = 29)	1.53 (2.44)	0.88 (0.45)	32.26 (0.48)	1.28 (1.29)	0.82 (0.39)	0.46 (0.94)	29.13 (3.33)	18.50 (2.93)	1.37 (0.17)	58.21 (9.08)
April STD (n = 30)	-1.02 (2.91)	0.08 (0.41)	33.39 (0.55)	0.79 (0.34)	0.61 (0.23)	0.19 (0.15)	27.04 (1.37)	16.84 (3.51)	1.28 (0.16)	51.85 (10.22)
May STD (n = 27)	1.00 (3.73)	-0.55 (0.55)	33.50 (0.53)	0.65 (0.38)	0.51 (0.29)	0.13 (0.11)	28.75 (3.28)	16.29 (3.25)	1.33 (0.21)	49.77 (11.02)
June STD (n = 19)	-4.78 (3.86)	-1.33 (0.38)	34.29 (0.49)	0.25 (0.11)	0.21 (0.08)	0.04 (0.04)	27.10 (1.35)	17.07 (3.31)	1.34 (0.17)	51.09 (9.94)
July STD (n = 6)	-3.40 (3.77)	-1.78 (0.07)	34.46 (0.26)	0.19 (0.07)	0.17 (0.06)	0.02 (0.03)	26.00 (1.51)	19.57 (5.29)	1.32 (0.22)	53.03 (12.03)
August STD (n = 6)	-3.59 (3.62)	-1.75 (0.09)	34.26 (0.59)	0.25 (0.08)	0.21 (0.08)	0.05 (0.04)	26.24 (1.26)	18.88 (2.13)	1.38 (0.06)	55.30 (7.45)
September STD (n = 17)	-1.38 (2.89)	-1.33 (0.48)	34.02 (0.89)	0.77 (0.48)	0.60 (0.29)	0.17 (0.21)	27.45 (2.00)	19.63 (2.58)	1.42 (0.14)	63.35 (8.22)
October STD (n = 29)	-1.28 (2.12)	-0.95 (0.44)	33.95 (0.40)	1.33 (1.53)	0.80 (0.78)	0.52 (0.81)	27.79 (2.24)	16.12 (3.80)	1.05 (0.32)	43.35 (17.40)
November STD (n = 30)	-0.08 (1.69)	-0.08 (0.41)	33.79 (0.57)	3.82 (6.87)	0.99 (1.02)	2.83 (5.88)	29.47 (4.10)	13.70 (3.20)	0.98 (0.21)	31.76 (9.80)
December STD (n = 30)	1.14 (1.30)	0.81 (0.70)	33.60 (0.53)	3.28 (4.94)	1.20 (0.97)	2.09 (4.06)	31.80 (6.57)	12.09 (3.01)	1.08 (0.22)	42.82 (12.36)

might have played a primary role in reducing surface nutrient concentrations during February (Brandini & Rebello 1994).

The annual mean values of nitrate, phosphate, and silicate are 17, 1.3, and 49.9 μM, respectively. Annual mean value of SS including organic and inorganic matters (Ahn et al. 1997) was 29.05 mg l⁻¹. The concentrations of nitrate, phosphate, and silicate generally increased throughout the year until September and decreased abruptly with an increase of microalgal biomass in early October (Table 1). The concentration of SS is also high from October to December. An increase of microalgal biomass from October onwards is indicative of a primary role of microalgal bloom on the low concentrations of nutrients and high SS concentration. The influence of glacial meltwater might have been minimal, since air temperature was still low and no significant increase in meltwater run-off was observed. The relatively high salinity during this period is also suggestive of a minimal role of glacial meltwater.

Annual mean of total chl *a* concentration in surface water was 1.38 mg m⁻³. A feature of microalgal biomass in Marian Cove was characterized by multimodal distribution of standing crop during the study period. chl *a* concentration showed a marked seasonality (Table 1). Monthly mean of total chl *a* concentrations in January and February were 1.74 and 1.61 mg m⁻³, respectively. After March (1.27 mg m⁻³), the chl *a* concentration declined rapidly, showing low values $<1 \text{ mg m}^{-3}$ during the period of April to September. A bloom started in October (1.33 mg m⁻³), resulting in high chl *a* concentrations during November and December (3.82 and 3.90 mg m⁻³, respectively). The pattern of monthly mean of nanoplanktonic microalgal ($<20 \mu\text{m}$, including pico-planktonic microalgae) chl *a* concentrations showed a similar pattern to that of total chl *a* concentration. Mean value of nanoplanktonic microalgal chl *a* concentration during austral spring (September, October and November) and summer (December, January and February) accounted for 0.81 and 1.34 mg m⁻³,

respectively. During autumn (March, April and May) and winter (June, July and August) the values were 0.65 mg m^{-3} and 0.19 mg m^{-3} , respectively. Microplanktonic microalgal (>20 μm) chl *a*, however, showed low concentration from January to October (Table 1). The microplanktonic microalgal chl *a* increased abruptly during November (2.83 mg m^{-3}) with a peak concentration of 31.17 mg m^{-3} on 28 November.

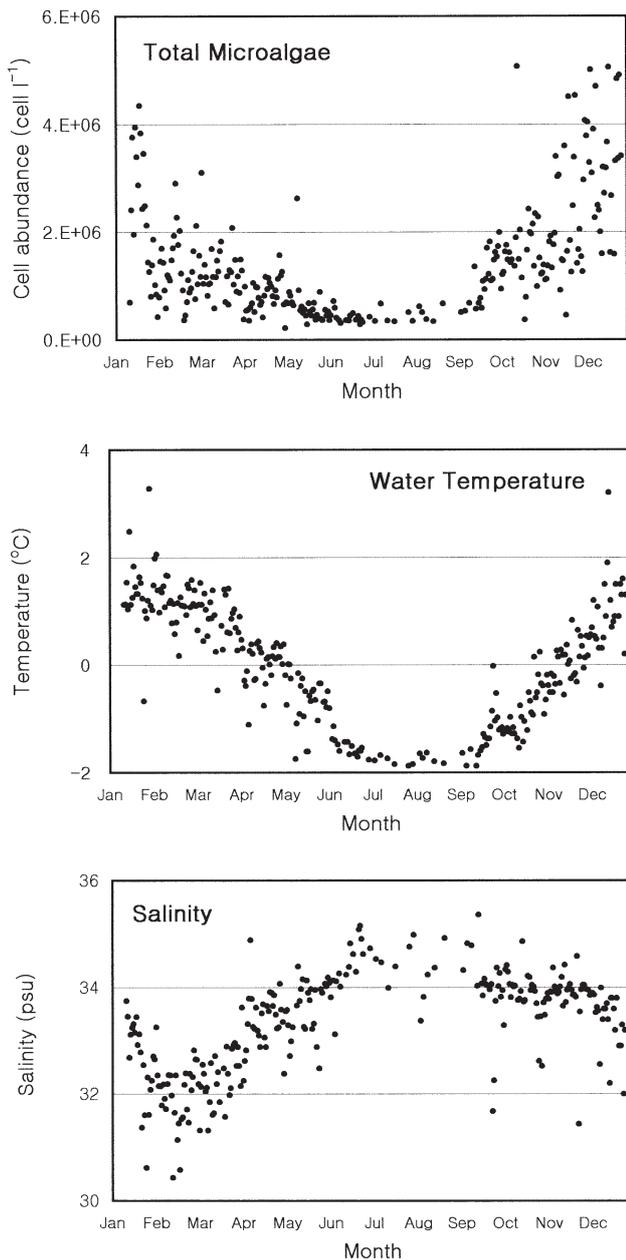


Fig. 2. Seasonal variation of total microalgal cell abundance (MCA), water temperature, and salinity

Seasonal variation of microalgal abundance

Annual mean of total microalgal cell abundance (MCA) in surface water was $2.43 \times 10^4 \text{ cells l}^{-1}$ (Fig. 2). Monthly mean of total MCA in January and February were 2.21×10^6 and $1.31 \times 10^6 \text{ cells l}^{-1}$, respectively. After March ($1.28 \times 10^6 \text{ cells l}^{-1}$) the abundance declined, reaching levels below $1 \times 10^6 \text{ cells l}^{-1}$ during April and August. A bloom started in September ($1.02 \times 10^6 \text{ cells l}^{-1}$) and the abundance of microalgae increased abruptly during October and November (1.68×10^6 and $2 \times 10^6 \text{ cells l}^{-1}$, respectively) peaking in December ($3.17 \times 10^6 \text{ cells l}^{-1}$). The highest MCA was $5.07 \times 10^6 \text{ cells l}^{-1}$ on 17 October.

The nanoplanktonic microalgae (including nano-sized flagellates and diatoms) such as *Cryptomonas* spp., *Phaeocystis antarctica* and *Navicula glaciei* were observed in high numbers in January ($2.20 \times 10^6 \text{ cells l}^{-1}$) and decreased in number to $0.21 \times 10^6 \text{ cells l}^{-1}$ on 2 May (Fig. 3). Monthly mean of nanoplanktonic MCA in January and February were 2.20×10^6 and $1.27 \times 10^6 \text{ cells l}^{-1}$, respectively (Fig. 4). After March ($1.24 \times 10^6 \text{ cells l}^{-1}$), the abundance declined, reaching levels below $1 \times 10^6 \text{ cells l}^{-1}$ during April and August. Nanoplanktonic microalgae started to bloom in October ($1.57 \times 10^6 \text{ cells l}^{-1}$) and abruptly increased in number during October and November (avg. $1.81 \times 10^6 \text{ cells l}^{-1}$) with a peak in December ($3.04 \times 10^6 \text{ cells l}^{-1}$).

The abundance of microplanktonic microalgae (including most diatoms), however, was low from January to September and then increased to $1.47 \times 10^6 \text{ cells l}^{-1}$ on 28 November (Fig. 3). *Fragilaria striatula* was the most abundant species among microplanktonic microalgae on that day. Monthly mean abundance of microplanktonic microalgae from January to September was $0.02 \times 10^6 \text{ cells l}^{-1}$. Monthly mean of microplanktonic MCA increased dramatically during October and November (0.11×10^6 and $0.19 \times 10^6 \text{ cells l}^{-1}$, respectively) and decreased in December ($0.13 \times 10^6 \text{ cells l}^{-1}$) (Fig. 4).

The most abundant phytoflagellate, *Phaeocystis antarctica*, occurred only in the motile flagellate form. Abundance of *P. antarctica* and other unidentified nanoplanktonic microalgae ($5 \pm 2 \mu\text{m}$) was relatively high during the study period (Fig. 5). Monthly mean of abundance in January was $2.09 \times 10^6 \text{ cells l}^{-1}$ (Fig. 5). After February ($1.15 \times 10^6 \text{ cells l}^{-1}$), abundance declined, reaching levels below $1 \times 10^6 \text{ cells l}^{-1}$ from April to September. A bloom started in September ($0.85 \times 10^6 \text{ cells l}^{-1}$) and abundance increased abruptly during October and November (1.13×10^6 and $1.36 \times 10^6 \text{ cells l}^{-1}$, respectively) and peaked in December ($2.53 \times 10^6 \text{ cells l}^{-1}$). Peak abundance of the prasinophyte *Pyramimonas* sp. ($0.07 \times 10^6 \text{ cells l}^{-1}$) was on 30 December, coinciding with the bloom of *P. antarctica*

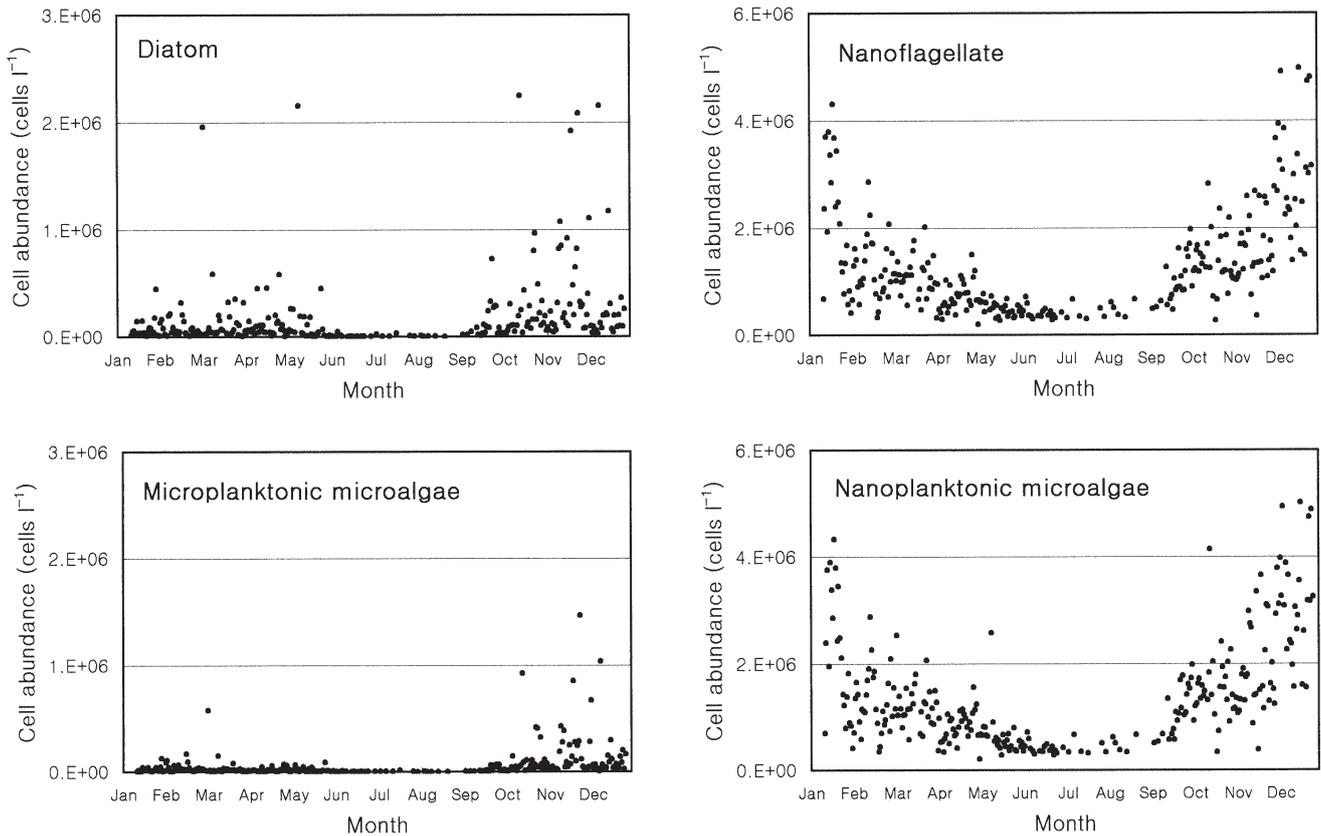


Fig. 3. Seasonal variation in cell abundance of diatoms, nanoflagellate, microplanktonic microalgae and nanoplanktonic microalgae

(Fig. 5). *Cryptomonas* spp., the second-most abundant phytoflagellate, was observed in high numbers during November and December (avg. 0.25×10^6 cells l^{-1} ; Fig. 5), accounting for 65.4% of total abundance of *Cryptomonas* spp.

Total diatom abundance showed a bimodal distribution (Fig. 4) with high values in March to May and November. In contrast, abundance of microplanktonic microalgae (most diatoms except *Minidiscus* spp., *Navicula glaciei/perminuta*, and *Rhoikosphenia* sp.) did not show a peak during March and May (Fig. 4), which is the time of highest *N. glaciei/perminuta* abundance (Fig. 5). The concentration of most diatoms increased rapidly during September and October. Especially benthic diatoms, such as *Achnanthes groenlandica*, *Fragilaria striatula*, *Licmophora belgicae*, and *L. gracilis* bloomed during October and November.

The most abundant microplanktonic diatom was *Fragilaria striatula*. This species showed large bimodal distribution (Fig. 6). The number of *F. striatula* was low from January to early October. Abundance of *F. striatula* increased dramatically to 8.42×10^5 cells l^{-1} on 28 November and then decreased during December (avg.

6.70×10^4 cells l^{-1}). Abundance between October and December (avg. 8.37×10^4 cells l^{-1}) accounted for 79.3% of annual mean abundance of *F. striatula* (Fig. 6). In the case of *Licmophora belgicae*, abundance was low from January to October and increased abruptly in November. Abundance during November (avg. 3.45×10^4 cells l^{-1}) accounted for 48.8% of total abundance of *L. belgicae* (7.06×10^4 cells l^{-1}).

On the other hand, temporal variations of planktonic diatoms, such as *Thalassiosira* spp. (mainly *T. antarctica*, *T. tumida*, *T. dichotomica* and *T. gracilis*), *Actinocyclus actinochilus*, and *Nitzschia lecointei* were dissimilar to those of benthic diatoms. Both size classes of *Thalassiosira* spp. (<20 μm and >20 μm) did not show a decrease in austral summer (Fig. 6). Highest mean abundance of both groups of *T. spp.* appeared in July (9.61×10^4 and 7.94×10^4 cells l^{-1} , respectively). The large planktonic diatom *Corethron pennatum* (= *C. criophilum*) was observed in low numbers from January to early October (Fig. 6). Abundance increased dramatically to 4.51×10^4 cells l^{-1} on 20 November and then decreased rapidly during December (avg. 1.80×10^2 cells l^{-1}). Abundance during November (avg. $6.20 \times$

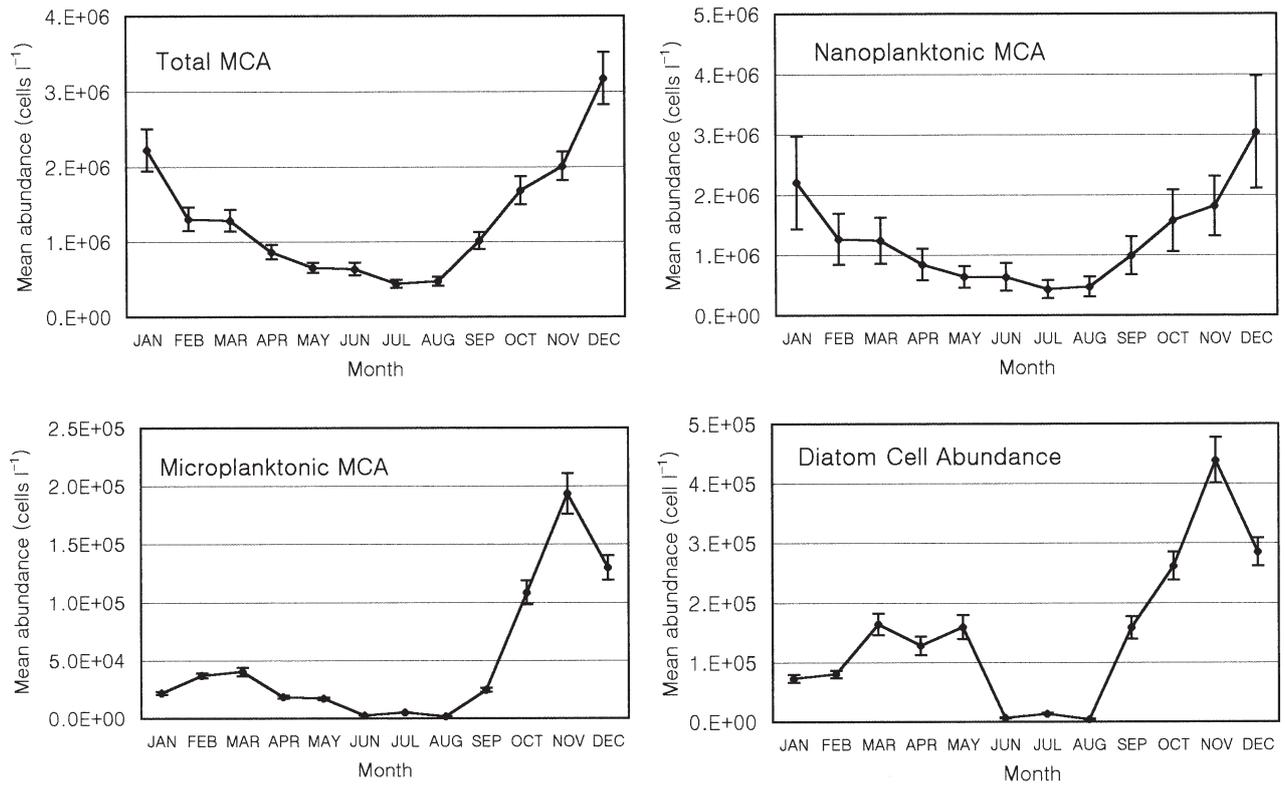


Fig. 4. Seasonal changes in monthly mean values of total microalgal cell abundance (MCA), microplanktonic MCA, nanoplanktonic MCA, and diatom cell abundance

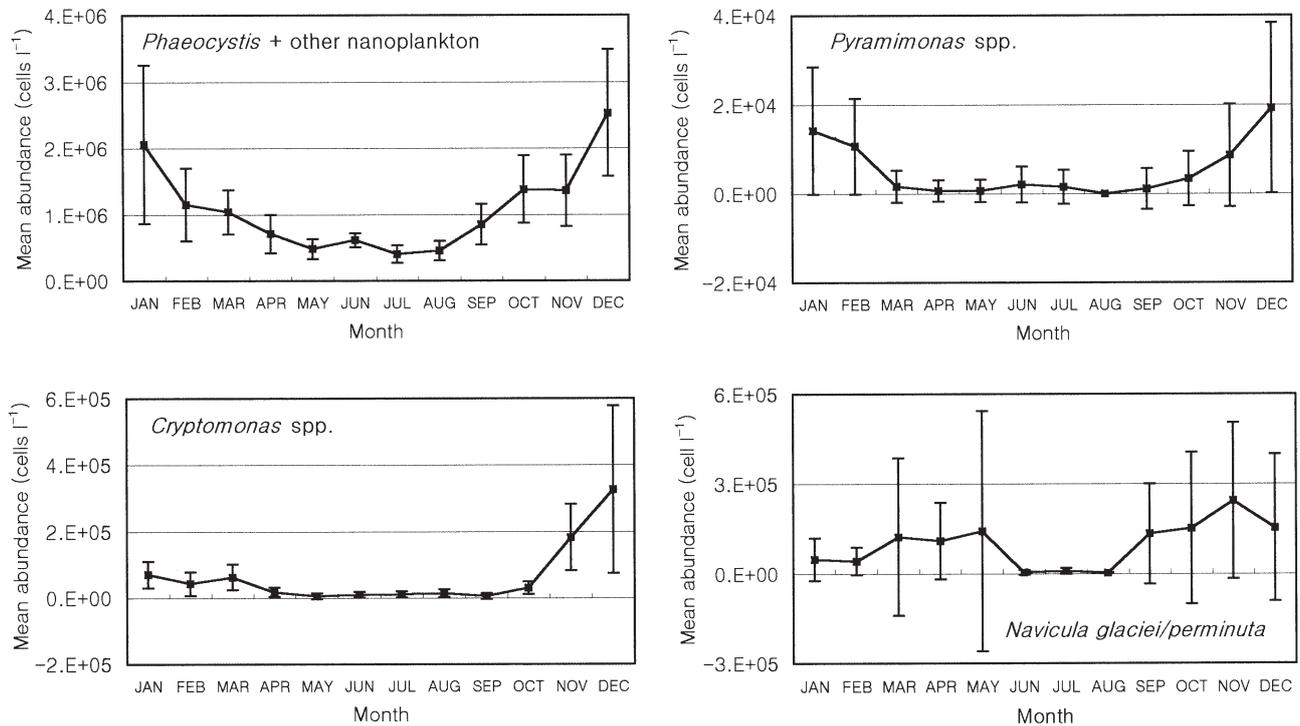


Fig. 5. Seasonal changes in monthly mean values of nanoplanktonic microalgal cell abundance (*Phaeocystis antarctica* and other unidentified nanoplankton, *Pyramimonas* spp., *Cryptomonas* spp., and *Navicula glaciei* and *N. perminuta*)

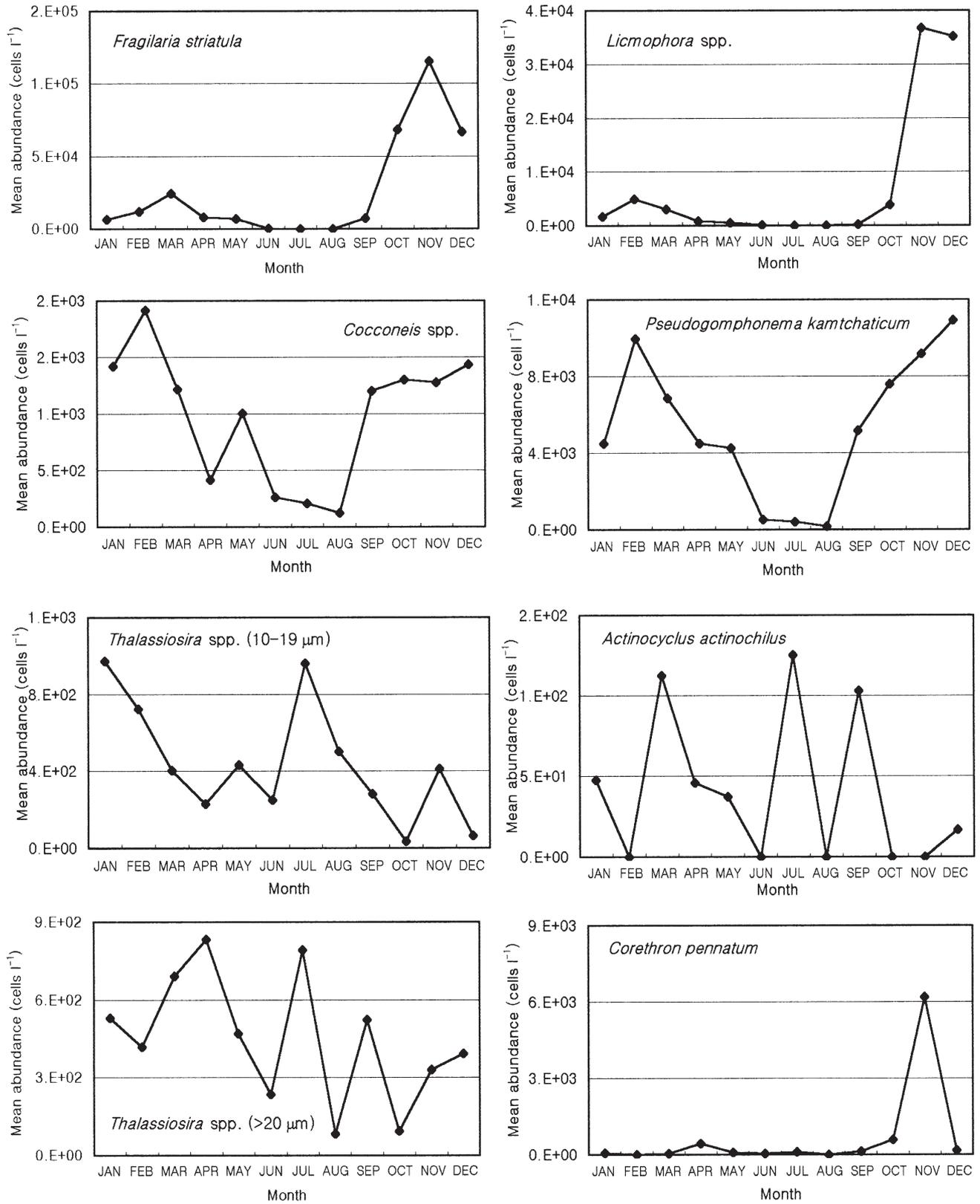


Fig. 6. Seasonal changes in monthly mean values of microplanktonic microalgal cell abundance (*Fragilaria striatula*, *Licmophora* spp., *Cocconeis* spp., *Pseudogomphonema kamtchaticum*, *Thalassiosira* spp. (10 to 19 μm), *Actinocyclus actinochilus*, *Thalassiosira* spp. (>20 μm) and *Corethron pennatum*)

10^3 cells l^{-1}) accounted for 78.7 % of total abundance of *C. pennatum* (7.88×10^3 cells l^{-1}).

Cryophilic diatoms *Navicula glaciei* and *N. perminuta* were the most abundant species of nanoplanktonic diatoms, showing higher numbers than other diatoms during all seasons. Because both species were not easy to identify with the use of light microscopy (LM), we did not split up the 2 species. Monthly mean abundance of *Navicula glaciei/perminuta* is shown in Fig. 5. Annual mean abundance of *N. glaciei/perminuta* in surface water was 9.69×10^4 cells l^{-1} . Abundance was low from December to February and June to August, whereas it was high from March to May and September to November (Fig. 5). Monthly mean abundance was highest in November (2.44×10^5 cells l^{-1}) and then decreased during December (1.54×10^3 cells l^{-1}).

Microalgal abundance in relation to environmental factors

In order to see the relation between the physico-chemical factors and microalgal abundance, we used Pearson's correlation matrix based on daily values (Table 2).

Total MCA was positively correlated with air temperature (r-value = 0.29, $n = 273$, $p < 0.01$, Table 2a). No significant correlations were observed between microplanktonic MCA and either air temperature or water temperature. However, significant correlations were seen between nanoplanktonic MCA and both temperatures (r-values = 0.30 and 0.41, respectively, $n = 273$). Also, MCA of nanoflagellate showed a positive correlation with air and water temperatures (r-values = 0.31 and 0.43, respectively, $n = 273$), whereas

Table 2. Pearson's correlation matrix of biological or environmental parameters ($n = 273$). The r-values shown in this table indicate statistical significance when p-values are < 0.01 . Asterisks indicate that r-values are not significant. (a) Correlation coefficients between salinity (SAL), air temperature (AIR), water temperature (WAT), wind speed (WIN), suspended solids (SS), nitrate (NIT), phosphate (PHO), silicate (SIL), total microalgae (TOT), nanoplanktonic microalgae (NAN), microplanktonic microalgae (MIC), nanoflagellate microalgae (NA-F) and diatom (DIA). (b) Correlation coefficients between salinity (SAL), air temperature (AIR), water temperature (WAT), wind speed (WIN), suspended solids (SS), nitrate (NIT), phosphate (PHO), silicate (SIL), *Cryptomonas* sp. (CRY), *Phaeocystis antarctica* (PHA), *Cocconeis* spp. (COC), *Corethron pennatum* (COR), *Fragilaria striatula* (FRA), *Licmophora* spp. (LIC), *Navicula* spp. (NAV), *Pseudogomphonema kamtchaticum* (PSE) and *Thalassiosira* spp. (THA)

(a)	SAL	AIR	WAT	WIN	SS	NIT	PHO	SIL	TOT	NAN	MIC	NA-F	DIA
Salinity	1												
Air temperature	-0.58	1											
Water temperature	-0.60	0.60	1										
Wind speed	*	*	-0.22	1									
Suspended solids	*	0.25	0.17	*	1								
Nitrate	*	*	-0.20	*	*	1							
Phosphate	*	*	-0.07	*	*	0.80	1						
Silicate	*	*	*	*	*	0.81	0.94	1					
Total microalgae	*	0.29	*	*	0.31	-0.31	-0.28	-0.20	1				
Nano-microalgae	*	0.30	0.41	*	0.27	-0.29	-0.26	-0.18	0.99	1			
Micro-microalgae	*	*	*	*	0.42	-0.26	-0.24	-0.22	0.53	0.42	1		
Nanoflagellate microalgae	*	0.31	0.43	*	0.22	-0.26	-0.22	*	0.95	0.97	0.29	1	
Diatom	*	*	*	*	0.37	-0.26	-0.24	-0.20	0.50	0.41	0.87	0.20	1

(b)	SAL	AIR	WAT	WIN	SS	NIT	PHO	SIL	CRY	PHA	COC	COR	FRA	LIC	NAV	PSE	THA
Salinity	1																
Air temperature	-0.58	1															
Water temperature	-0.60	0.60	1														
Wind speed	*	*	-0.22	1													
Suspended solids	*	0.25	0.17	*	1												
Nitrate	*	*	-0.20	*	*	1											
Phosphate	*	*	-0.07	*	*	0.80	1										
Silicate	*	*	*	*	*	0.81	0.94	1									
<i>Cryptomonas</i> sp.	*	0.20	0.30	*	0.24	-0.33	-0.27	-0.22	1								
<i>Phaeocystis</i> + other nanoplanktonic microalgae	*	0.31	0.42	*	0.20	-0.23	-0.21	-0.14	0.59	1							
<i>Cocconeis</i> spp.	-0.21	0.32	0.17	0.27	0.31	*	*	*	*	0.20	1						
<i>Corethron pennatum</i>	*	*	*	*	*	-0.27	-0.32	-0.34	0.27	*	0.17	1					
<i>Fragilaria striatula</i>	*	*	*	*	*	0.35	-0.25	-0.25	-0.22	0.23	0.51	0.30	0.26	1			
<i>Licmophora</i> spp.	*	*	*	*	0.52	-0.23	*	*	0.26	0.25	0.22	0.16	0.80	1			
<i>Navicula glaciei/perminuta</i>	0.23	*	*	*	0.27	-0.22	-0.20	*	*	*	0.31	0.29	0.64	0.43	1		
<i>Pseudogomphonema kamtchaticum</i>	-0.20	0.19	*	0.18	0.44	*	*	*	*	0.22	0.59	*	0.68	0.61	0.65	1	
<i>Thalassiosira</i> spp.	*	*	0.17	*	*	*	0.17	*	*	0.17	-0.07	*	*	*	*	*	1

diatom was not significantly correlated with either temperature. This suggests that the bloom of nanoplanktonic flagellate was induced by temperature increase. Nano-flagellates, *Cryptomonas* spp. and *Phaeocystis antarctica* were the only species showing significant positive correlation with both temperatures (Table 2b).

Salinity was negatively correlated with air and water temperatures (r -values = -0.58 and -0.60 , $n = 273$, Table 2a). Little correlation was observed between MCA and salinity, but *Cocconeis* spp. and *Pseudogomphonema kamtschaticum* were negatively correlated with salinity (Table 2b). These species, which were positively correlated with air temperature and wind speed, were found only on subtidal macroalgae such as the red alga *Plocamium cartilagineum* (<5 to 30 m) and the brown algae *Desmarestia* spp., the most dominant macroalgae in the depth range of 5 to 15 m during austral summer (Ahn et al. 1997).

Significant correlations were seen between MCA and concentration of SS (Table 2a). Abundant microalgal species, except the planktonic diatoms, *Corethron penatum* and *Thalassiosira* spp., were positively correlated with SS. Ahn et al. (1997) also reported a significant correlation between chl *a* concentration and SS in Maxwell Bay during austral summer. These correlations indicate that increase of SS was influenced by benthic diatoms during summer (see 'Discussion').

Nutrients such as nitrate, phosphate, and silicate were negatively correlated with MCA, which probably reflects microalgal uptake (Table 2). Not only microplanktonic MCA and diatoms but also nanoplanktonic MCA were significantly correlated with nutrients. Abundance of nanoflagellates, which have no silicate theca, was not negatively correlated with silicate. This means that the dramatic increase of microalgae significantly affected the decrease of nutrients. Benthic diatoms such as *Cocconeis* spp. and *Pseudogomphonema kamtschaticum*, however, were not significantly correlated with nutrients.

DISCUSSION

Seasonal variation of microalgal abundance in relation to environmental factors

Total chl *a* concentration showed low values <1 mg m^{-3} during the period of April to September (Fig. 5). In Factory Cove, the lowest chl *a* concentration was observed in July (0.23 mg m^{-3}) (Clarke et al. 1988). In Marian Cove, MCA was low during winter, with a minimum monthly mean of 4.4×10^5 cells l^{-1} in July. The abundance of *Phaeocystis antarctica* and other unidentified nanoplanktonic microalgae (5 ± 2 μm) was relatively high during winter.

Total chl *a* increased during summer. The nanoplanktonic diatoms, *Navicula glaciei/perminuta*, started to increase in abundance at the beginning of spring, when other microplanktonic microalgae (especially diatoms) also began to bloom (Fig. 4). Microalgal bloom started in October and abruptly increased in abundance during November 1996 (Table 1). More than 45% (avg. 3.5 mg m^{-3}) of chl *a* was concentrated during November and December. In Factory Cove, chl *a* started to increase in November, climbed very rapidly during December, and peaked in January (Clarke et al. 1988). The microalgal bloom started on average about 2 mo earlier in Marian Cove than in Factory Cove.

Nanoplanktonic microalgae also increased during summer (Fig. 4). In October, nanoflagellate *Cryptomonas* spp. started to increase in abundance and then continuously increased until December (Fig. 5). According to Moline et al. (2000), cryptophytes were dominant in low salinity water characteristic of the meltwater input near Anvers Island. Also, a cryptophyte bloom was observed in the upper 20 m of the water column within the coastal polynya between 76.0 and $76.5^\circ S$ (Arrigo et al. 1999). In the Weddell-Scotia-Bellingshausen Confluence areas cryptophytes have been found to make a major contribution to the total biomass along the retreating ice-edge zones (Buma et al. 1992).

In our data, *Cryptomonas* spp. was significantly correlated with air and water temperatures during 1996, showing positive correlations with both temperatures from October to December (r values = 0.38 and 0.56 , respectively; $n = 89$, $p < 0.01$). It was not significantly correlated with salinity, wind speed and SS. Salinity showed neither normal correlation nor lag correlation with *Cryptomonas* spp. Such a positive correlation between *Cryptomonas* spp. and temperature has been observed in the Bransfield Strait (Kang & Lee 1995, Mura & Agustí 1998, Kang et al. 2001). There, the growth of cryptophytes was highly responsive to small changes in water temperature and relatively warm water (ca. $2.4^\circ C$) was the optimal temperature for the growth of the cryptomonad (Mura & Agustí 1998). In Anvers Island, cryptophytes also showed a positive correlation with the daily mean air temperature during the 5-yr study period (Moline et al. 2000). The increase in air temperature decreased salinity because of ice melting, and subsequently increased water temperature. The latter stimulated the growth of cryptophytes (Mura & Agustí 1998). In our study, the insignificant correlation between salinity and *Cryptomonas* spp. is also suggestive of the minimal role of salinity in the growth of cryptophytes. Sea ice had already retreated far from the study site during the period of *Cryptomonas* spp. bloom (November to December).

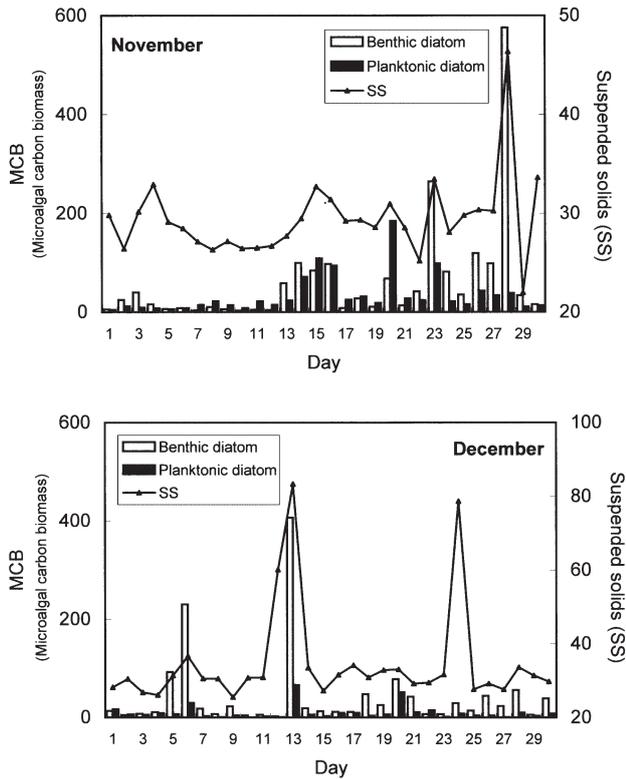


Fig. 7. Comparison of suspended solids (SS) and microalgal carbon biomass (MCB) of benthic and planktonic diatoms in November and December

Microplanktonic microalgae were also increased during summer (Fig. 4). High diatom biomass during November and December was caused by tycho planktonic benthic diatoms from the sea bottom, not by *in situ* production and growth of phytoplankton. Fig. 7 shows the microalgal carbon biomass of benthic and planktonic diatoms in November and December. As planktonic biomass was low in the water column for a large part of the summer growing season, such an increase in benthic diatoms would have an important role on the increase of chlorophyll. The proportion of total benthic diatoms to the total microalgal species was also consistently high in November and December (83 and 61%, respectively) near Davis Station (Everitt & Thomas 1986).

Total diatom and benthic diatom carbon biomass were positively correlated with SS (*r*-values 0.43 and 0.47, respectively; *n* = 60, *p* < 0.01) in November and December. SS comprise inorganic and organic matters that were usually supplied by meltwater streams, resuspension

of benthic material, and subglacial discharge during summer. The inorganic matter (mostly sediment) accounted for >80% of total SS in Marian Cove (Ahn et al. 1997). It was most likely supplied by wind-driven resuspension of sea-bottom material and by terrigenous input from meltwater run-off. Although SS was not significantly correlated with wind speed, resuspension of benthic diatoms from the sea bottom due to wind-induced turbulence might be affecting the secondary production of adjacent waters (Demers et al. 1987, Brandini & Rebello 1994). Brandini & Rebello (1994) reported that the strong turbulence generated by winds resuspended benthic diatoms from the sediment of the inner shallow sections of the Admiralty Bay. In other areas of Antarctic coastal environments, resuspension of sublittoral deposit has been suggested as a primary process for increasing suspended solids (Krebs 1983, Berkman et al. 1986, Everitt & Thomas 1986, Dunbar et al. 1989, Gilbert 1991, Brandini & Rebello 1994, Klöser et al. 1994, Ahn et al. 1997). The resuspension of benthic microalgae would significantly affect secondary production and the diversity of nearshore phytoplankton assemblages (Ahn 1996).

Abundant benthic diatoms such as *Fragilaria striatula*, *Licmophora belgicae*, and *L. gracilis* were densely attached on Rhodophyta *Bangia* sp. during austral summer 1996 (Fig. 8A). *F. striatula* and *Licmophora* spp. were significantly correlated with SS (Table 2B). The resuspension process would also make benthic diatoms available for benthic suspension feeders (Klöser et al. 1994, Ahn et al. 1993, 1997). Ahn et al. (1993) reported that benthic diatoms were abundant in the gut content of suspension-feeding bivalve *Laternula elliptica*. The tycho planktonic species in intertidal and subtidal diatom communities seem to contribute

Table 3. Seasonal changes in monthly mean values of total microalgal carbon biomass (MCB), nanoflagellate carbon biomass (CB), diatom CB, nanoplanktonic MCB (<20 µm), and microplanktonic MCB (>20 µm)

	Total MCB	Nano-flagellates CB	Diatoms CB	Nano-planktonic MCB	Micro-planktonic MCB
January (n = 21)	20.07	9.50	10.58	12.31	7.76
February (n = 29)	23.18	7.29	15.89	9.62	13.56
March (n = 29)	27.80	6.76	21.05	13.56	14.24
April (n = 30)	18.21	3.36	14.85	9.39	8.82
May (n = 27)	15.23	2.32	12.91	10.16	5.07
June (n = 19)	4.17	2.68	1.49	2.91	1.26
July (n = 6)	5.90	2.40	3.50	2.87	3.03
August (n = 6)	2.73	2.14	0.59	2.29	0.44
September (n = 17)	16.50	2.62	13.88	9.98	6.52
October (n = 29)	40.87	5.35	35.52	13.77	27.10
November (n = 30)	109.28	12.43	96.61	25.94	83.34
December (n = 30)	75.33	21.50	53.82	30.04	45.29
Integral	29.94	6.53	23.39	11.90	18.03

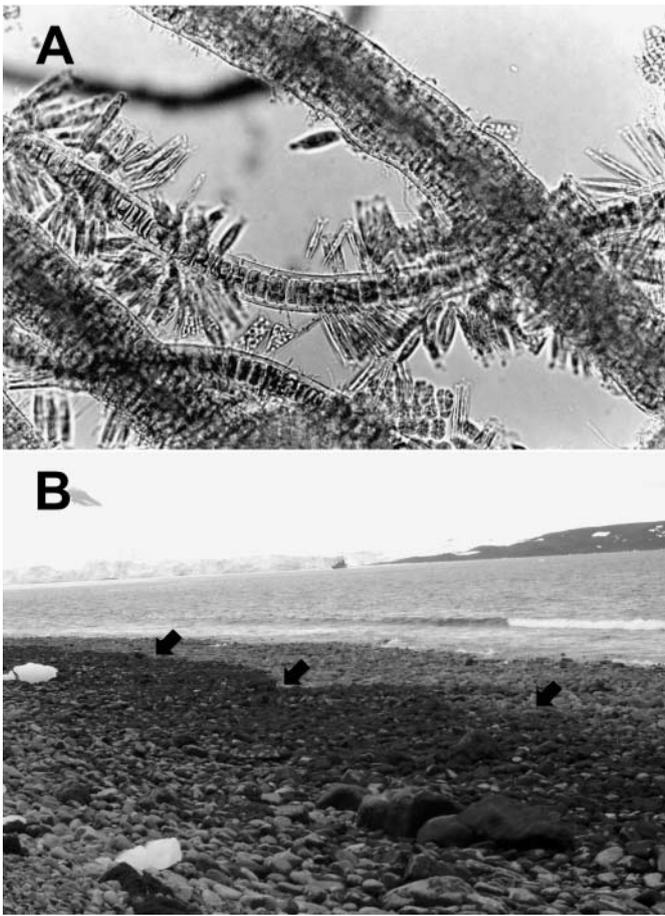


Fig. 8. (A) Microphotographs of benthic diatoms *Fragilaria striatula* and *Licmophora belgicae* attached on macrophyte *Bangia* sp. (B) Photographs of *Bangia* sp. and benthic diatom mat formed in intertidal zone. Note the boundary (see arrows) between the mat-forming area and grazed area by *Nacella concinna*

substantially to the carbon source for the nearshore ecosystems of Marian Cove.

The benthic diatoms *Fragilaria striatula* and *Licmophora belgicae* increased abruptly during November and started to decrease in December (Fig. 6). This is probably due to grazing by the Antarctic limpet *Nacella concinna*, which is the most important grazer in the Antarctic shallow water (Walker 1972). *N. concinna* usually grazed filamentous algae *Bangia* sp. and benthic diatoms, and increased in abundance during summer near King Sejong Station (Dr. Dohong Kim pers. obs.; Fig. 8B), which might have caused the decrease of benthic-diatom abundance during summer.

The benthic diatoms *Cocconeis* spp. and *Pseudogomphonema kamtschaticum* were negatively correlated with salinity and positively correlated with air temperature and wind speed (Table 2b). These species are usually attached on subtidal macroalgae, *Des-*

marestia mensiesii (Ahn et al. 1997, Klöser 1998). The increase of *Cocconeis* spp. and *P. kamtschaticum* during summer (low salinity and high air temperature) was probably due to favorable physicochemical conditions of summer, which are suitable for the growth of these diatoms. The positive correlation with wind speed suggests that *Cocconeis* spp. and *P. kamtschaticum* were detached from the macroalgae and suspended into the water column during periods of strong wind.

The seasonal variation of *Cocconeis* spp. and *Pseudogomphonema kamtschaticum* also seems to have been influenced by substrate conditions (Fig. 6). The specific growth rate and photosynthetic capacities of the brown algae *Desmarestia mensiesii* increased during the late winter to spring and declined in summer (Gómez & Wiencke 1997). The rapid growth rate of macroalgae and changes in pH at the thallus surface tend to prevent epiphytism (den Hartog 1972) and might have contributed to the paucity of the species during winter and spring. When the macroalgae grew slowly during summer, epiphytes would cover them rapidly, resulting in high abundance of *Cocconeis* spp. and *P. kamtschaticum*.

In conclusion, monitoring the seasonal variations of environmental factors and microalgal abundance in Marian Cove during 1996 has provided insights into the interactions of biological and environmental events. Air and water temperatures mainly affected the seasonal variation in nearshore microalgal abundance. The increase of microalgal abundance resulted in the depletion of nutrients. As Clarke & Leakey (1996) pointed out, we need continuous and interdisciplinary year-round studies to understand the highly seasonal and complex nature of nearshore marine ecosystems.

Succession of microalgae in the nearshore of King Sejong Station during 1996

Diatoms accounted for the major carbon source in nearshore waters of King Sejong Station during 1996. Annual mean of total microalgal carbon biomass (MCB) was $29.9 \mu\text{gC l}^{-1}$ (Table 3). Annual mean of diatom carbon biomass was $23.3 \mu\text{gC l}^{-1}$, which accounted for 77.9% of total MCB. Annual means of nanoplanktonic MCB and microplanktonic MCB were 11.9 and $18.0 \mu\text{gC l}^{-1}$, respectively. This suggests that microplanktonic microalgae were more important as a carbon source than nanoplanktonic microalgae, although nanoplanktonic MCA ($1.26 \times 10^6 \text{ cells l}^{-1}$) was

relatively higher than microplanktonic MCA (5×10^4 cells l^{-1}).

During summer (December through February), major carbon sources in January were the nanoflagellate *Phaeocystis antarctica* and other unidentified nanoplanktonic microalgae ($5.15 \mu\text{g C l}^{-1}$), whereas in December and February it was *Licmophora belgicae* (14.9 and $3.14 \mu\text{g C l}^{-1}$, respectively). During fall (March through May), *Navicula glaciei/perminuta* was a major carbon source (6.77 , 6.02 , and $7.80 \mu\text{g C l}^{-1}$, respectively). During winter, carbon mainly originated from *P. antarctica* and other unidentified nanoplanktonic microalgae (June: $0.99 \mu\text{g C l}^{-1}$), *Thalassiosira* spp. (July: $1.33 \mu\text{g C l}^{-1}$), and *N. glaciei/perminuta* (August: $0.14 \mu\text{g C l}^{-1}$). Mean of total chl *a* was $<0.3 \text{ mg m}^{-3}$ during winter. In spring, *N. glaciei/perminuta* (September: $7.32 \mu\text{g C l}^{-1}$), *Fragilaria striatula* (October: $14.9 \mu\text{g C l}^{-1}$), and *Licmophora belgicae* (November: $29.5 \mu\text{g C l}^{-1}$) were major carbon sources.

Consequently, *Navicula glaciei/perminuta* was the major carbon source during late winter and early spring, when other microalgae were not yet in bloom. Subsequently tychoplanktonic benthic diatoms *Fragilaria striatula* and *Licmophora belgicae* started to increase in number and were abundant during summer. *N. glaciei/perminuta* was a major carbon source again during fall. *Phaeocystis antarctica* and other unidentified nanoplanktonic microalgae and *Thalassiosira* spp., which were abundant in winter, were probably introduced from outside of Maxwell Bay.

In conclusion, microalgae as the carbon source for benthic animals underwent changes seasonally with respect to abundance and species composition. Benthic diatoms were the major carbon source during summer and the cryophilic diatom *Navicula glaciei/perminuta* was the major source during spring, fall, and early winter. The planktonic species of *Thalassiosira* spp., *Phaeocystis antarctica*, and other unidentified nanoplanktonic phytoflagellates were major food sources for benthic animals during winter.

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