

# Relative importance of phototrophic, heterotrophic, and mixotrophic nanoflagellates in the microbial food web of a river-influenced coastal upwelling area

Cristian A. Vargas<sup>1,\*</sup>, Paulina Y. Contreras<sup>1</sup>, José Luis Iriarte<sup>2</sup>

<sup>1</sup>Aquatic Systems Research Unit, Environmental Sciences Center EULA Chile, Universidad de Concepción, PO Box 160-C, Concepción, Chile

<sup>2</sup>Instituto de Acuicultura, Facultad de Pesquerías y Oceanografía, Universidad Austral de Chile, PO Box 1327, Puerto Montt, Chile

**ABSTRACT:** The contribution of phototrophic nanoflagellates (PNF) as primary producers and heterotrophic (HNF) and mixotrophic (MNF) nanoflagellates as major grazers of bacterioplankton was assessed during a 3 yr study in a highly productive, river-influenced coastal upwelling area under contrasting seasons (winter/non-upwelling vs. spring/upwelling). Sampling was conducted at 2 stations—around a river plume and at an intensive seasonal upwelling site—with contrasting environmental gradients. The MNF were evaluated functionally for the possession of photo-pigments and experimentally for the ability to take up prey, specifically fluorescently labeled bacteria (FLB). During the short-term experiments, we estimated grazing rates over FLB for 2 size categories of HNF and MNF. Both bacterial production (BP) and nanophytoplankton primary production (PP) were higher in the river plume area. PNF abundance ranged from  $6 \times 10^9$  to  $411 \times 10^9$  cells  $m^{-2}$ , whereas HNF abundance fluctuated between  $27 \times 10^9$  and  $267 \times 10^9$  cells  $m^{-2}$ . In contrast, the abundance of MNF was usually low, with a maximum of  $\sim 7 \times 10^9$  cells  $m^{-2}$ . For MNF, ingestion rates were between 7.3 and 30.7 bacteria per flagellate and hour (bact flag<sup>-1</sup> h<sup>-1</sup>), whereas HNF ingestion ranged from 2 to 7.5 bact flag<sup>-1</sup> h<sup>-1</sup>. However, since HNF dominated in terms of abundance, they were the dominant grazers on bacterioplankton populations. Estimates of grazing pressure for the microbial food web showed that MNF were capable of removing 1 to 51 % BP d<sup>-1</sup>, whereas HNF could control BP, eliminating from 24 % BP d<sup>-1</sup> up to more than 100 % BP d<sup>-1</sup>. Given the area's relatively high nutrient condition, the elevated MNF biomass in the river plume and the greater bacterivory impact from MNF in winter, it seems that light and, thus, the energy/carbon limitation could be the main trigger for mixotrophy in this river-influenced coastal upwelling area.

**KEY WORDS:** Mixotrophy · River discharge · Upwelling area · Plankton food web

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## INTRODUCTION

Protozoan nanoflagellates are ubiquitous members of planktonic food webs in both freshwater and marine aquatic ecosystems (Laybourn-Parry & Parry 2000). This group of small organisms encompasses a wide range of nutritional types, including purely photosynthetic cells, different degrees of mixotrophy, and heterotrophs, which consume mostly bacteria and small algal cells (González et al. 1990, Sherr &

Sherr 1994). The importance of bacterivorous nanoflagellates in planktonic food web dynamics and biogeochemical cycles has been recognized since the works by Pomeroy (1974) and Azam et al. (1983). By then, the traditional classical food web concept (phytoplankton-zooplankton-fishes) incorporated the major top-down role of prokaryotes and different protists into a 'microbial food web'. In this food web, about 40 to 60% of the carbon fixed by primary producers is thought to pass through heterotrophic

\*Email: crvargas@udec.cl

bacterioplankton (Cole et al. 1988) rather than directly into herbivorous zooplankton. This bacterial biomass is controlled by mortality through viral infection as well as grazing by both heterotrophic (HNF) and mixotrophic nanoflagellates (MNF), especially in oligotrophic environments (Christaki et al. 1999). Furthermore, grazing by HNF and MNF may commonly balance bacterial production (BP) (Anderesen & Fenchel 1985). Nanoflagellates may, in turn, transfer bacterial or small algal carbon indirectly to larger protozoans (ciliates and dinoflagellates) and/or directly to mesozooplankton grazers (Vargas et al. 2007, 2010), which would indeed strengthen the link between the microbial and classical food webs, increasing the trophic efficiency of the microbial loop (Sanders 1991).

Recent research has demonstrated the widespread occurrence of mixotrophy as a significant mode of nutrition in both freshwater (e.g. Bennett et al. 1990) and marine nanoflagellate communities (e.g. Unrein et al. 2007). MNF forms are able to combine phototrophic nutrition with the uptake of particulate or dissolved organic material (Sanders 1991). The occurrence and high clearance rates of MNF (e.g. Bennett et al. 1990, Czipionka et al. 2011) suggest that they might contribute substantially to total bacterivory in aquatic ecosystems. Furthermore, photosynthesis by MNF may also contribute to primary production in the nanoplankton size fraction in coastal areas (Marshall & Laybourn-Parry 2002, Jost et al. 2004).

Mixotrophy has been suggested to be a nutritional strategy for algal cells during low nutrient conditions (Jost et al. 2004), which are not common at productive coastal upwelling sites (Vargas et al. 2007). Hence, the ecological role of mixotrophy has not been incorporated into carbon fluxes and food web models for these productive coastal environments and is probably underappreciated at the moment. Although bacterivory by both HNF and MNF is likely to play a major role in channeling bacterial production to higher trophic levels in situations of coastal environmental gradients, field studies contrasting the role and impact of MNF during different seasons are scarce for marine ecosystems (e.g. Tsai et al. 2007, Unrein et al. 2007). These studies can be relevant for estuarine and river-influenced areas where the flux of dissolved organic matter from freshwater runoff may be significant and bacterial biomass and production are relatively high (Dagg et al. 2004).

The incorporation of mixotrophy into food web models of highly productive, temperate coastal areas requires a robust data set with precise estimations of

the nanoflagellate grazing impact on bacterioplankton, as opposed to a data set of strictly heterotrophs over contrasting seasons (spring vs. winter) and environmental gradients (e.g. upwelling vs. river discharges). During this study, we experimentally investigated the nanoflagellate alga community in winter and spring, asking one simple question: Is mixotrophy a significant strategy (>20% of total biomass) in this highly-productive, temperate coastal upwelling area? We determined the relative importance of phototrophic nanoflagellates (PNF) as phytoflagellate primary producers and the relevance of HNF and MNF as major grazers of bacterioplankton in a productive, river-influenced coastal upwelling area. Although 'mixotrophy' describes many forms of nutritional strategies on the continuum from absolute phototrophy to absolute heterotrophy, we used this word to refer to algae combining phototrophic and phagotrophic heterotrophy. Mixotrophy was evaluated functionally by the possession of photo-pigments and experimentally by the ability to take up prey (i.e. fluorescently labeled bacteria, FLB).

## MATERIALS AND METHODS

### Study area

The river-influenced coastal upwelling area off Concepción, central Chile (36° S), is located on one of the widest parts of the continental shelf along the eastern boundary coastal upwelling margin. This area also receives freshwater discharges from 2 important rivers: the Itata River (mean flow: 296 m<sup>3</sup> s<sup>-1</sup>) and the Bío-Bío River (mean flow: 1699 m<sup>3</sup> s<sup>-1</sup>). The runoff from these rivers supplies the coastal ocean with substantial amounts of silicate, nitrate, and phosphate as well as trace metals (Sánchez et al. 2008). This area is characterized by seasonal upwelling events in austral spring–summer (Vargas et al. 2007), when upwelling-driven phytoplankton blooms, which are typically dominated by long-chain-forming diatoms, occur in the shallow embayment (Vargas et al. 2007) and result in one of the highest primary production (PP) rates (~4 to 20 g C m<sup>-2</sup> d<sup>-1</sup>) reported worldwide for coastal areas (Montero et al. 2007). Two nearshore stations with contrasting environmental gradients (maximum depth: 40 m) were established off the coast of central Chile. One station (Stn RV) was located around the plume of the Itata River and the other (Stn UW) approximately 20 km southeast of Coliumo Bay, in an area of intensive seasonal upwelling events (Fig. 1). The

selection of these stations was based in terms of the major driver for nutrient input in the euphotic zone (river discharge vs. coastal upwelling).

### Seawater sampling and hydrography

Samples were taken during seven 2 d cruises aboard the LC 'Kay Kay' (Universidad de Concepción) in spring/summer and winter as part of the OMMIX Project (influence of OMnivory and MIXotrophy in food-web dynamic). Field sampling was conducted twice in December 2006 (spring), June 2007 (winter), and December 2007 (spring), and once in August 2008 (winter). At each station, temperature, salinity, oxygen, and fluorescence profiles were recorded from near the bottom to the surface using a SeaBird SBE-19 plus CTD equipped with a YSI-calibrated Beckman oxygen sensor and a Wetstar fluorometer. Water samples (1 l) for PNF and HNF abundance and biomass were collected at discrete depths (1, 5, 10, 25, and 35 m) with a rosette system equipped with twelve 8 l Niskin bottles. Nanoplankton samples were preserved in glutaraldehyde (2.0% w/v in 0.2  $\mu\text{m}$  prefiltered seawater). Additional samples for nutrient concentrations and short-term grazing experiments (5 l) were also collected from the surface, maximum fluorescence depth, and the base of the thermocline (see 'Short-

term grazing experiments' below). Samples for nutrient analysis, including nitrate ( $\text{NO}_3^-$ ), nitrite ( $\text{NO}_2^-$ ), and phosphate ( $\text{PO}_4^{3-}$ ), were filtered (GF/F filters) on board and frozen ( $-20^\circ\text{C}$ ) until analysis in the laboratory.  $\text{NO}_3^-$  and  $\text{PO}_4^{3-}$  were determined spectrophotometrically following Strickland & Parsons (1968).  $\text{NO}_2^-$  was analyzed using an automated nutrient analyzer (ALPKEM, Flow Solution IV) following the protocol of the US Environmental Protection Agency; Method 353.2. River flow data were obtained from the National Water Directorate ([www.dga.cl](http://www.dga.cl)), whereas PAR (photosynthetically active radiation) time-series were obtained from a HOBO weather station (Onset Computer Corp., USA) installed by the COPAS Center (Universidad de Concepcion) at  $36^\circ 31.687' \text{S}$  and  $72^\circ 57.955' \text{W}$ .

### Bacterioplankton and nanoplankton abundance and biomass estimates

The abundances of bacteria and cyanobacteria were estimated by flow cytometry. Samples of 1350  $\mu\text{l}$  were collected in 1500  $\mu\text{l}$  vials, fixed with paraformaldehyde (1% final concentration) and quick-frozen in liquid nitrogen. Subsamples of 150  $\mu\text{l}$  were processed on a FACSCalibur flow cytometer equipped with an Ion-Argon Laser of 488 nm and 15 MW (Becton Dickinson). Coccioid cyanobacteria (*Synechococcus*) and photosynthetic eukaryotes were distinguished by differences in side light scatter and by fluorescence in the orange (cyanobacteria) and red (eukaryotes) wavelengths. Heterotrophic bacteria abundance was obtained from samples previously stained with SYBR Green I (Molecular Probes) (Marie et al. 1997). The bacterial volume was calculated from length and width measurements of at least 20 cells per sample using standard geometric forms for a sphere and an ellipsoid (i.e. rods and bacilli), respectively. Finally, the volume-to-carbon conversion was carried out using a factor of  $0.35 \text{ pg C } \mu\text{m}^{-3}$  (Bjørnsen 1986).

Nanoflagellates were quantified by the proflavine technique (Haas 1982). For the enumeration of nanoflagellates, subsamples were filtered on a  $0.8 \mu\text{m}$  polycarbonate membrane filter, stained with proflavine

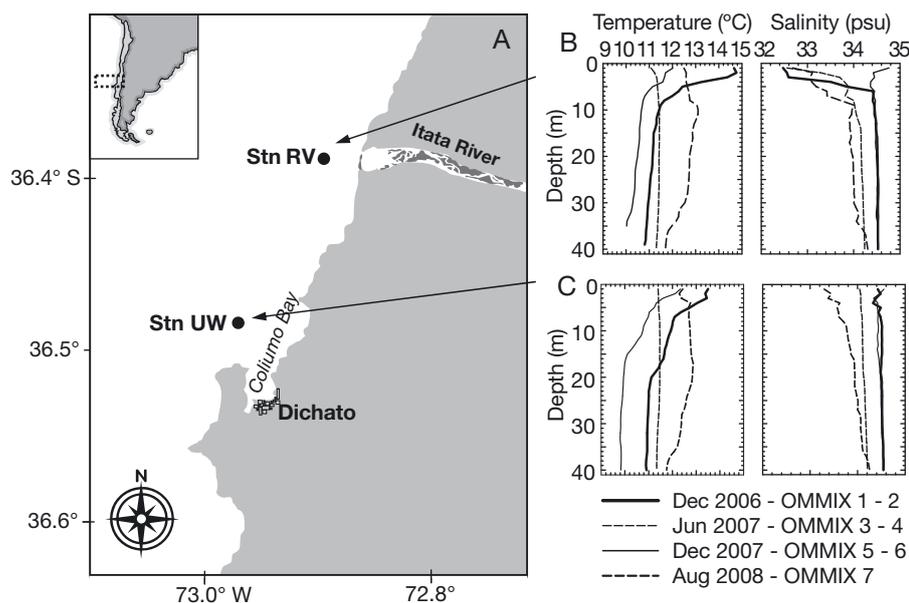


Fig. 1. (A) Study area and locations of the sampling stations in the Itata River plume area (Stn RV) and upwelling site off Coliumo Bay (Stn UW), including seasonal averages of temperature ( $^\circ\text{C}$ ) and salinity (psu) at (B) Stn RV and (C) Stn UW during OMMIX cruises 1 to 7

(0.033% w/v in distilled water) following Haas (1982), and fixed with glutaraldehyde (as above) for subsequent analysis. Nanoflagellates were counted with an inverted microscope OLYMPUS IX-51 equipped with UV model UMWU2 (width band pass 330 to 385 nm) and FITC model U-MWB2 (width band pass 450 to 480 nm) filter sets. Nanoflagellates were measured and their biovolume was estimated from a minimum of 80 cells per group (PNF, MNF, and HNF). Nanoflagellate counts considered 2 major size classes: <5  $\mu\text{m}$  and 5–20  $\mu\text{m}$ , with a mean volume of 11.7 ( $\pm 0.5$ ) and 47.6 ( $\pm 2.8$ )  $\mu\text{m}^3$  for each size class. Biomass was estimated using a size-dependent carbon:volume ratio as suggested by Verity et al. (1992). The trophic mode of nanoflagellates (phototrophic, heterotrophic, and/or mixotrophic) was assessed by estimating the FLB uptake during a short-term incubation (see 'Short-term grazing experiments' below).

#### Nanoplankton primary production and bacterial production

Water samples for PP estimates were collected in a 5.0 l PVC Go-Flo bottle (General Oceanics) at 4 depths: 0 m, maximum fluorescence depth (ca. 28% of surface irradiance), the thermocline base (i.e. halocline base in winter, ca. 51% of surface irradiance), and 20 m (ca. 2% surface irradiance). Samples were incubated in 125 ml polycarbonate bottles (2 clear + 1 dark bottle) and placed in a natural-light incubator for ca. 4 h (roughly 10:00 to 14:00 h). Running surface seawater over the incubation bottles regulated ambient temperature and light intensity was attenuated using a screen (i.e. nylon fibers matrix, gray colour) to approximate light at the depth where the water was collected (Iriarte & González 2004). Radio-labeled sodium bicarbonate (30 to 40  $\mu\text{Ci NaH}^{14}\text{CO}_3$ ) was added to each bottle. Primary production was measured using the method described by Steemann-Nielsen (1952). Samples were manipulated under subdued light conditions during pre- and post-incubation periods. Filters (0.7  $\mu\text{m}$ ) were placed in 20 ml plastic scintillation vials and kept at  $-15^\circ\text{C}$  until reading (15 d later). To remove excess inorganic carbon, filters were treated with HCl fumes for 24 h. A cocktail (8 ml, Ecolite) was added to the vials and radioactivity was determined in a Beckmann scintillation counter. Differential size fractionation of phytoplankton to determine nanoplankton PP and chlorophyll *a* (chl *a*) was carried out by pre-filtering the seawater samples using a 20  $\mu\text{m}$  Nitex mesh. The

filtrate was then collected on a 2.0  $\mu\text{m}$  Nuclepore polycarbonate filter. Water samples for chl *a* measurements were taken at the same depths as the samples collected for PP analysis. Seawater samples (200 ml) were filtered following the same procedure used for PP, extracted in 90% v/v acetone, and analyzed using a digital PS-700 Turner fluorometer (Parsons et al. 1984). Depth-integrated nanoplankton PP and chl *a* values through the photic layer were estimated by the trapezoidal integration method.

The ambient BP was estimated from the L-[ $^{14}\text{C}$ (U)]-leucine incorporation rates (Simon & Azam 1989). From each depth, 10 ml of seawater from each depth were transferred to sterile tubes (3 replicates and 1 blank). The blank was poisoned with 0.2  $\mu\text{m}$  filtered formaldehyde. We added radio-labeled leucine (300 to 330 mCi mmol $^{-1}$  s.a., Sigma T-6527) to obtain final saturation concentrations of 50 nM, which has been already reported as the mean saturation level for the same study area by Hernández et al. (2006). All tubes were incubated in the dark for 1 to 2 h at the surface *in situ* water temperature. We also selected the incubation period for both sampling stations considering previous studies in the same coastal region (i.e. Troncoso et al. 2003, Hernández et al. 2006). After incubation,  $^{14}\text{C}$ -leucine incorporation was stopped by adding 2% formalin. The samples were extracted with ice-cold trichloroacetic acid (TCA, final conc. of 5%) for 15 min. The extracted samples were kept cool, filtered onto 0.45  $\mu\text{m}$  cellulose nitrate filters (Whatman), and rinsed 3 times with 5 ml of 5% ice-cold TCA. The filters were placed into scintillation vials and dissolved in 1 ml of ethylacetate. A volume of 10 ml of liquid scintillation cocktail (Ecolitel) were added and the samples were radio assayed. Leucine uptake was estimated from dpm using a Packard (Model 1600TR) liquid scintillation counter; counting efficiency was calculated from the non-quenched standard of  $^3\text{H}$ -toluene. BP from leucine incorporation was calculated using a ratio of cellular carbon to protein of 0.86 and a fraction of leucine in protein of 0.073 (Simon & Azam 1989). The cell production rates, obtained from moles of leucine incorporated (see Fuhman & Azam 1982), were transformed to BP assuming a widely used conversion factor of  $2 \times 10^{18}$  (Lee & Fuhman 1987).

#### Short-term grazing experiments

Seawater samples for short-term grazing experiments were taken from 3 different layers: (a) the maximum fluorescence depth (FMD), (b) thermocline

base (i.e. halocline base in winter; BT), and (c) a deeper layer (25 m depth). Samples were filtered through a 20  $\mu\text{m}$  mesh-sized Nitex-net to obtain the nanoplankton fraction, whereon they were left undisturbed for 1 h at the *in situ* temperature until setting up the experiment. The functional analysis of the nanoplankton community was done using FLB following the methods of Sherr et al. (1987). FLB uptake experiments were carried out in triplicate in 250 ml acid-cleaned polycarbonate bottles (Nalgene) and incubated in an on-deck rack equipped with a shading device to mimic the *in situ* temperature and the *in situ* light regime when necessary. In the assay, FLB were added to a final concentration of  $10^5$  FLB  $\text{ml}^{-1}$ , following Sherr et al. (1987), which in fact represented between 10 to 20% of the ambient bacterial abundance. Subsamples of 30 ml were collected at 0, 10, 30, and 50 min, fixed with glutaraldehyde (final concentration of 2%; Christaki et al. 1999), and stained with a combination of 1200  $\mu\text{l}$  of DAPI (diamidino phenylindole; final concentration: 4  $\mu\text{g ml}^{-1}$ ) and 600  $\mu\text{l}$  of proflavine (final concentration: 0.033%; Haas 1982). Stained samples were filtered on 0.8  $\mu\text{m}$  Nuclepore filters for counting nanoflagellates and stored at  $-20^\circ\text{C}$  until analysis. Analysis was performed at a magnification of  $\times 1000$  with an Olympus IX-51 using one filter (excitation UV-light 340 to 380 nm, emission 43 to 485 nm) to observe DTAF (dichlorotriazinyl aminofluorescein) and DAPI fluorescence and a second filter (excitation 450 to 490 nm, emission 515 nm) to observe proflavine and chlorophyll auto-fluorescence. At least 250 flagellates per filter were classified according to their size (2–5 and 5–20  $\mu\text{m}$ ) and nutritional regime: phototrophic (i.e. chloroplast autofluorescence), heterotrophic (i.e. proflavine stain without autofluorescence), and/or mixotrophic (i.e. chloroplast autofluorescence and FLB-ingestion). Food vacuole content was estimated for both mixotrophic and heterotrophic cells based on their FLB ingestion. Since mixotrophic forms were classified based on their FLB uptake, without consideration of inactive grazers as for HNF forms, ingestion rate was overestimated relative to heterotrophic cells by using this methodology. For both mixotrophs and heterotrophs, the mean food vacuole content was used for clearance estimations. We calculated the number of ingested FLB  $\text{ind.}^{-1} \text{h}^{-1}$  from the slope of food vacuole content versus incubation time for each bottle multiplied by the FLB concentration. Average values of ingested FLB  $\text{ind.}^{-1} \text{h}^{-1}$

were calculated for triplicate bottles and clearance was measured over the linear portion of the uptake curve. Finally, ingestion rates were calculated by multiplying the clearance rate by the ambient bacterial abundance at the sampling station. Carbon balances were made using data for ingestion rates assuming a gross growth efficiency of 32% (Straile 1997). Mixotrophs were assumed to be only 50% as efficient at fixing inorganic carbon as strict autotrophs (Hansen et al. 2000, Adolf et al. 2006). Consequently, cell specific PP was determined by calculating their share of the total nanoplankton PP.

Potential differences in biomasses and production between stations RV and UW were tested for using a randomized paired *t*-test that compared depth-integrated biomass estimates production for each taxonomic group, and bacterial and size-fractioned primary within each season (see Manly 1997). A Spearman's rank correlation between MNF and HNF ingestion rates and environmental variables were also conducted on the  $\log(x + 1)$  transformed data.

## RESULTS

### Hydrography

Values of surface temperature, salinity, and inorganic nutrient concentrations clearly reflected the different physical/chemical environments of the sampling periods (Figs. 1, 2 & 3A). As typical at this latitude, spring and winter periods showed contrasting scenarios in light conditions (PAR), with values from 1500 to 2000  $\mu\text{E m}^{-2} \text{s}^{-1}$  in spring and  $< 1000$   $\mu\text{E}$

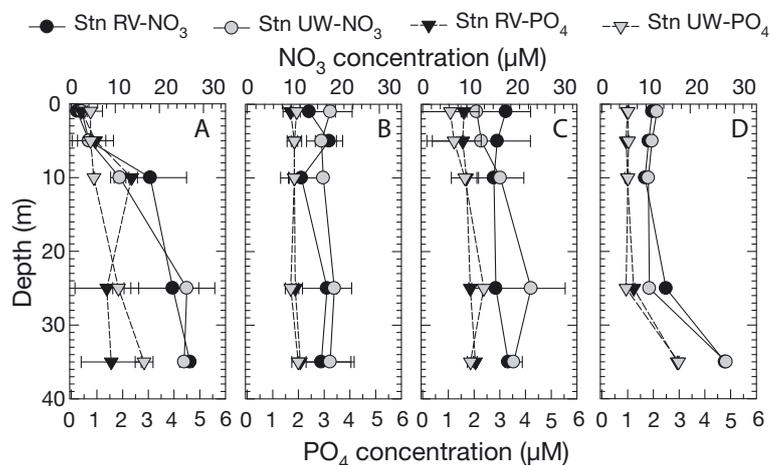


Fig. 2. Concentration (mean  $\pm$  SD;  $\mu\text{mol}$ ) of nitrate ( $\text{NO}_3$ ) and orthophosphate ( $\text{PO}_4$ ) at both sampling stations and averaged for each sampling season (A: December 2006; B: June 2007; C: December 2007; D: August 2008)

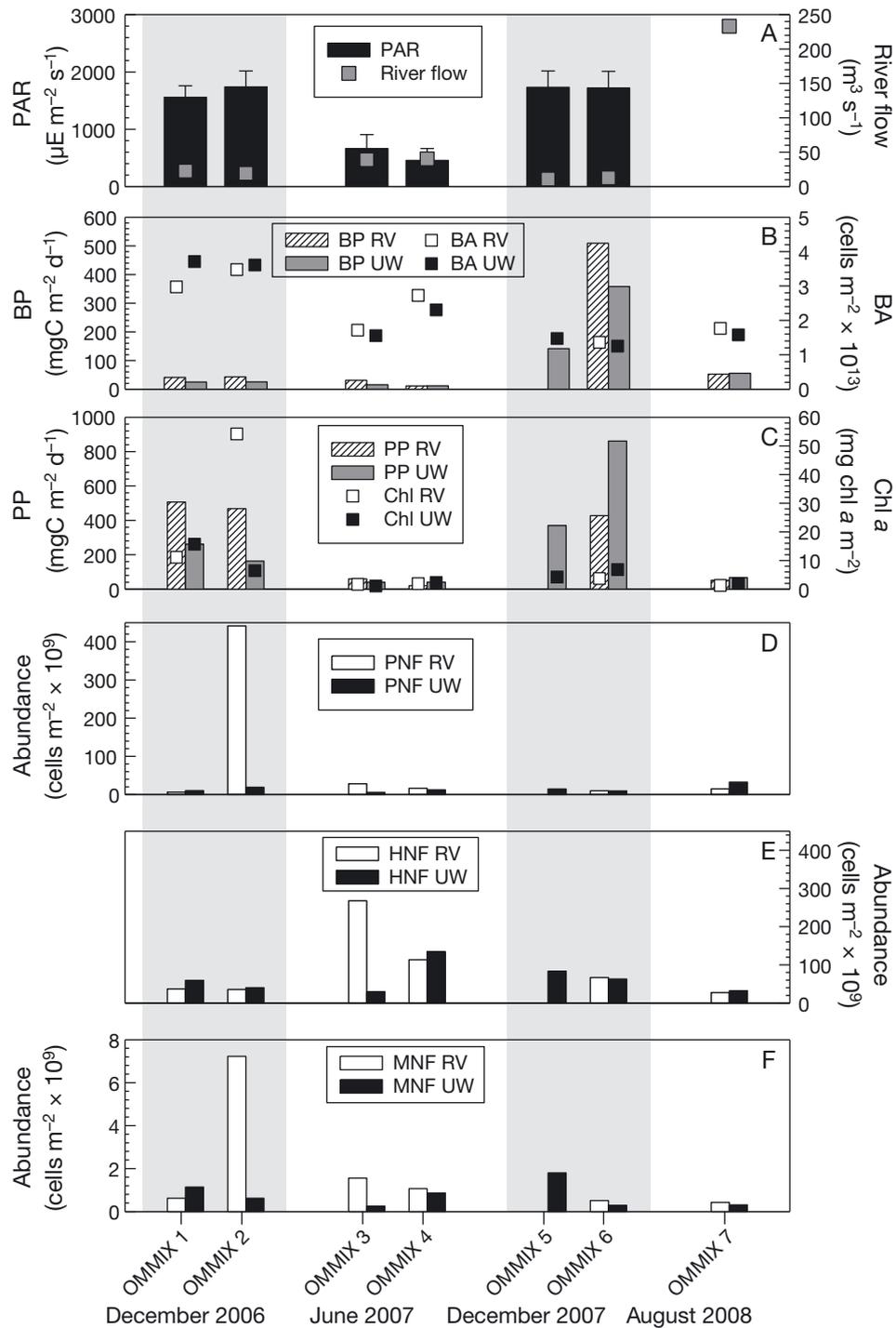


Fig. 3. Seasonal variations (spring vs. winter) of ambient parameters, rates, and abundances evaluated at both sampling stations (Stns RV and UW) during OMMIX cruises 1 to 7: (A) photosynthetically active radiation (PAR) and river flow, (B) upper 20 m depth-integrated bacterial production (BP) and bacterial abundance (BA), (C) nanophytoplankton primary production (PP) and chlorophyll a biomass (Chl a), and (D–F) abundance of phototrophic (PNF), heterotrophic (HNF) and mixotrophic (MNF) nanoflagellates, respectively. Shaded areas represent spring sampling periods

$\text{m}^{-2} \text{s}^{-1}$  in winter (Fig. 3A). CTD profiles indicated a well-defined seasonal thermocline located between 3 and 10 m depth at both sampling stations during the

spring periods (OMMIX 1, 2, 5, 6) (Fig. 1). Salinity profiles at Stn RV (Fig. 1) evidenced the influence of the freshwater flow on the hydrographic characteris-

tics of the surface water column, with a shallow lens of less-saline water located in the upper 5 m, except in spring 2007, when river flow was  $<20 \text{ m}^3 \text{ s}^{-1}$  (Fig. 2A). In spring 2006, the water column showed high thermal and haline stratification, with low nutrient concentrations in the upper 10 m depth. In spring 2007,  $\text{NO}_3^-$  and  $\text{PO}_4^{3-}$  were not depleted in the surface waters ( $>10$  and  $1 \mu\text{M}$ , respectively), and higher values were associated with the river plume (Stn RV). In winter, the vertical distribution of nutrients was relatively homogeneous in the upper 25 m depth, with non-significant differences between stations ( $t$ -test:  $p > 0.05$ ; Fig. 2B,D).

### Bacterioplankton and nanoplankton abundance, biomass and production

Maximum fluorescence depth was  $\sim 10$  m during spring and from 5 to 8 m depth during winter, and non-significant differences were observed between both sampling stations. The mean bacterial cell volume ( $0.3$  to  $0.7 \mu\text{m}^3$ ) did not vary substantially between stations (data not shown). In the upper 20 m, integrated BP ranged between 11 and  $509 \text{ mg C m}^{-2} \text{ d}^{-1}$ . During most of the field surveys, BP was relatively low ( $<60 \text{ mg C m}^{-2} \text{ d}^{-1}$ ), and high values ( $\sim 500 \text{ mg C m}^{-2} \text{ d}^{-1}$ , Fig. 3B) were only observed in December 2007 (OMMIX 6). BP was always higher in the river plume area at Stn RV (Fig. 3B), and the highest values of BP were associated with the FMD in spring and with deeper waters in winter ( $\sim 25$  m depth; Fig. 4). In contrast, the average BP at Stn UW was higher at the base of the thermocline for all field campaigns. In general terms, bacterial abundance was higher at Stn RV, except in spring 2006. A gross estimation of bacterial growth rate (BGR) (calculated as a ratio of BP over biomass) showed that BGR was always higher at the Stn RV ( $t$ -test,  $p < 0.01$ ) (data not shown).

Upper 20 m depth-integrated PP in the nanoplankton size fraction ranged from 20 to  $862 \text{ mg C m}^{-2} \text{ d}^{-1}$ , with the lowest values recorded in winter 2007 and the highest in spring 2007 (Fig. 3C). A trend of higher nanophytoplankton PP in the river plume was also observed during the study period, except in spring 2007 (OMMIX 6). Despite extremely high chl  $a$  in December 2006 (OMMIX 2) at Stn RV ( $\sim 55 \text{ mg chl } a \text{ m}^{-2}$ ; Fig. 3C), significant differences were not found at the 2 sampling stations over time ( $t$ -test;  $p > 0.01$ ). A randomized paired  $t$ -test comparing depth-integrated PP and BP at the 2 sampling stations showed that only BP was substantially different at Stn RV and UW in spring, being higher in the river plume.

Fig. 3D–F shows the integrated abundance of PNF, MNF, and HNF in the upper 20 m depth. Flagellates  $<5 \mu\text{m}$  in size were much more abundant than those between 5 and  $20 \mu\text{m}$ , whereas heterotrophic forms were more abundant than strictly plastidic ones during most winter periods. PNF abundance varied widely, from  $6 \times 10^9$  to  $411 \times 10^9 \text{ cells m}^{-2}$ . However, with the exception of a large peak in abundance of small PNF during OMMIX 2 in December 2006 at Stn RV, abundances were commonly  $<35 \times 10^9 \text{ cells m}^{-2}$  (Fig. 3D). There was no clear seasonal trend in the abundance of PNF at either sampling site. The HNF abundance ranged between  $27 \times 10^9$  and  $267 \times 10^9 \text{ cells m}^{-2}$ , and was highest in winter (July) 2007 at the Stn RV (Fig. 3E). The MNF abundance was usually low, with a maximum of  $\sim 7 \times 10^9 \text{ cells m}^{-2}$ . Moreover, it broadly reflected the temporal pattern of the PNF abundance, which made the greatest contribution during the peak of PNF in December 2006 (OMMIX 2; Fig. 3F). On average, in spring, the highest abundance of PNF, MNF, and HNF was associated with the maximum fluorescence depth (PNF =  $7.8$  to  $8.8 \text{ cells ml}^{-1}$ , MNF =  $0.06$  to  $0.2 \text{ cells ml}^{-1}$ , and HNF =  $2.4$  to  $3.5 \text{ cells ml}^{-1}$ ), whereas in winter, these maxima were concentrated toward the bottom of the thermo-

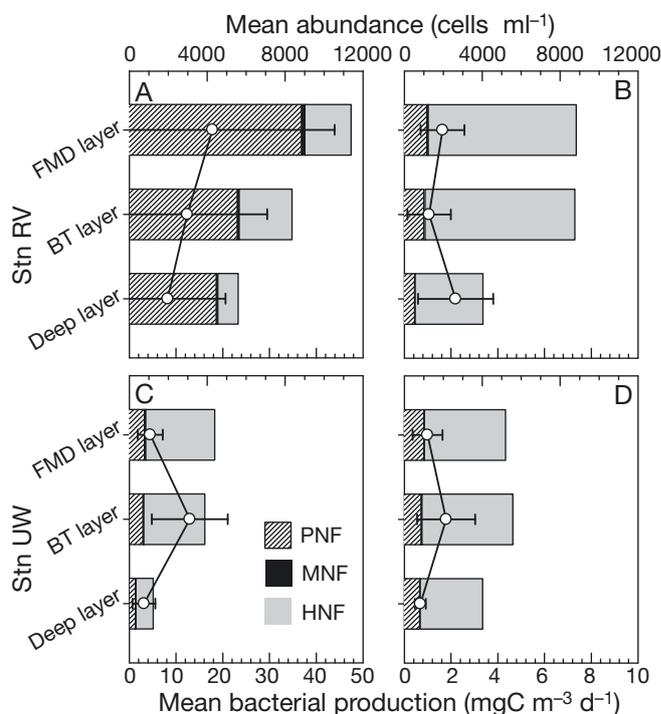


Fig. 4. Vertical distribution of mean ( $\pm$ SD) bacterial production ( $\text{mg C m}^{-3} \text{ d}^{-1}$ ) and nanoflagellate abundance ( $\text{cells ml}^{-1}$ ) for the 3 categories analyzed (PNF, MNF, HNF): (A,B) in the river plume area, Stn RV, and (C,D) off Coliumo Bay, Stn UW, averaged for spring (A,C) and winter (B,D) conditions

cline (Fig. 4). We did not observe significant differences in integrated biomasses between stations for the different groups of flagellates or for bacteria (*t*-test;  $p > 0.01$ ; Table 1, Fig. 5a), despite the peak in

Table 1. Results of randomized paired *t*-tests comparing depth-integrated biomass and size-fractionated primary production and bacterial production between stations within each season. Significant differences ( $\alpha = 0.05$ ) are shown in **bold**. Bonferroni-corrected significance levels for multiple biomass and primary production comparisons were  $\alpha_b = 0.0056$  and  $\alpha_p = 0.0167$ , respectively. PNF, MNF and HNF: phototrophic, mixotrophic and heterotrophic nanoflagellates, respectively

Group	Spring		Winter	
	<i>t</i>	<i>p</i>	<i>t</i>	<i>p</i>
Biomass (mg C m <sup>-2</sup> )				
Bacteria	-1.13	0.757	<b>3.53</b>	<b>&lt;0.001</b>
PNF	0.97	0.246	<b>13.96</b>	<b>&lt;0.001</b>
MNF	1.11	0.121	0.02	0.373
HNF	0.71	0.243	1.39	0.115
Primary production (mg C m <sup>-2</sup> d <sup>-1</sup> )				
Size class 2–20 μm	0.14	0.249	1.55	0.130
Bacterial production (mg C m <sup>-2</sup> d <sup>-1</sup> )				
Bacterioplankton	1.21	<b>0.024</b>	0.05	0.950

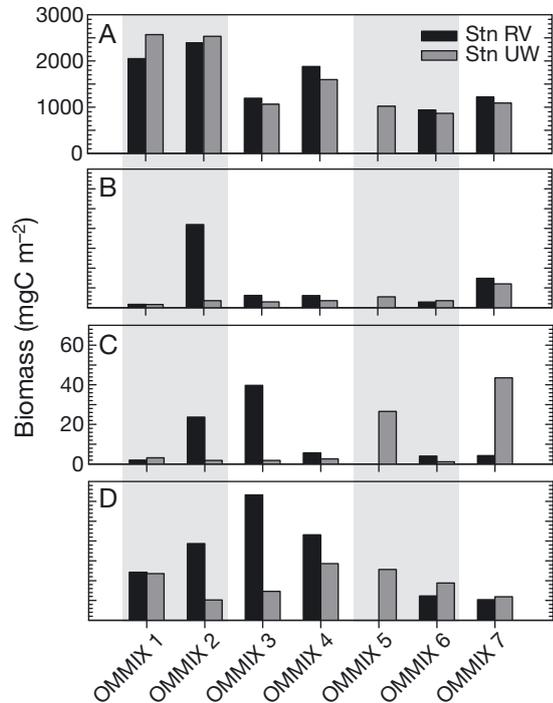


Fig. 5. Upper 20 m depth-integrated biomass of (A) bacterioplankton, (B) phototrophic, (C) mixotrophic, and (D) heterotrophic nanoflagellates during each sampling campaign (OMMIX 1–7) in the river plume area (Stn RV) and off Coliumo Bay (Stn UW). Shaded areas represent spring and white areas winter sampling periods

PNF biomass observed during OMMIX 2 (Fig. 5B). In winter, bacteria and PNF biomass were significantly higher at Stn RV (*t*-test,  $p < 0.01$ ; Table 1). In winter 2007, biomass values were higher for both MNF and HNF around the river plume (Stn RV; Fig. 5C,D).

### Ingestion rates

There was always a background of small HNF for which FLB ingestion was rarely observed (80 to 90%). Average flagellate ingestion rates determined during the different sampling periods varied between spring and winter (Fig. 6). MNF ingestion rates oscillated between 7.3 and 30.7 bact flag<sup>-1</sup> h<sup>-1</sup>, whereas HNF ingestion ranged from 2 to 7.5 bact flag<sup>-1</sup> h<sup>-1</sup>, and ingestion rate showed a better correlation with bacterial abundance for small mixotrophic

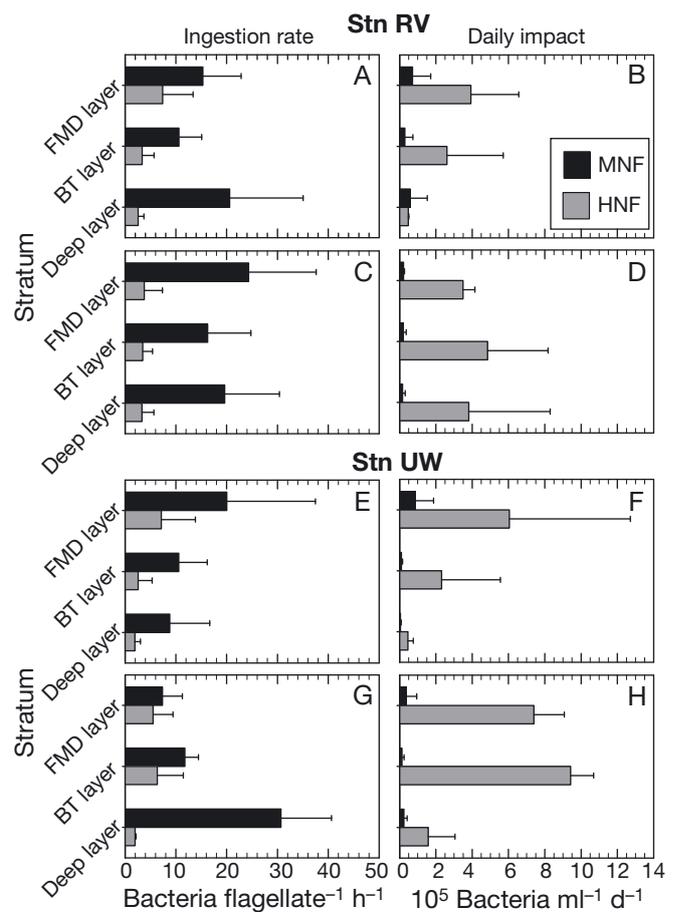


Fig. 6. Mean ( $\pm$  SD) ingestion rates (bacteria flagellate<sup>-1</sup> h<sup>-1</sup>) and daily grazing impact on bacterioplankton (bacteria ml<sup>-1</sup> d<sup>-1</sup>) determined through short uptake experiments of fluorescently labeled bacteria averaged for (A,B,E,F) spring and (C,D,G,H) winter at 2 sampling stations (Stns RV and UW). MNF (HNF): mixotrophic (heterotrophic) nanoflagellates

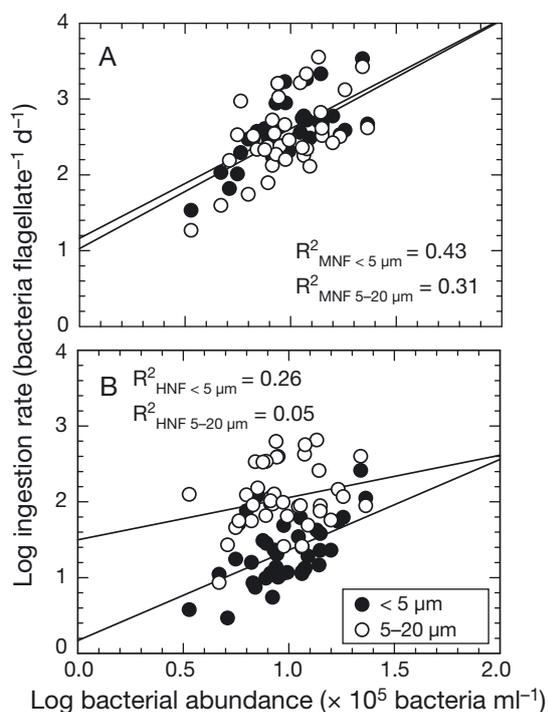


Fig. 7. Relationship between log bacterial abundance ( $\times 10^5$  cells  $\text{ml}^{-1}$ ) and log ingestion rates (bacteria flagellate $^{-1}$  d $^{-1}$ ) of (A) mixotrophic and (B) heterotrophic nanoflagellates in 2 size categories:  $< 5 \mu\text{m}$  and  $5\text{--}20 \mu\text{m}$

forms than those for heterotrophs (Fig. 7). However, in terms of daily impact, the trend was just the opposite due to the numerical dominance of HNF in the water column (Fig. 3E); in other words, grazing by HNF had a stronger effect. The highest ingestion rates (up to  $\sim 30$  bact flag $^{-1}$  h $^{-1}$ ) for MNF were found in winter. The daily grazing impact of MNF and HNF on the bacterioplankton communities ranged from  $0.04 \times 10^5$  to  $1 \times 10^5$  and from  $0.5$  to  $9.4 \times 10^5$  bacteria  $\text{ml}^{-1}$  d $^{-1}$ , respectively. Bacterivory by MNF had the greatest impact on the standing stock in spring in the FMD layer at both sampling stations. On the other hand, the highest daily impact of HNF occurred in winter, mostly associated with the base of the thermocline. We also evaluated the degree of prey saturation by plotting the log-values of ingestion rates for phagotrophic nanoflagellates against the log-values of bacterial abundances (Fig. 7). A linear function described the relations better than an exponential model, and the best fits—according to the coefficient of determination ( $R^2$ )—were found for MNF. A Spearman's rank correlation analysis between nanoflagellate ingestion and environmental variables, including bacterial abundance and production, showed that MNF ingestion by both size classes was significantly correlated with bacterial abun-

dance, whereas ingestion by small HNF was also correlated with bacterial abundance but also negatively correlated with N:P ratio (HNF  $< 5 \mu\text{m}$ ) (Table 2).

Carbon balances were made using ingestion rate data (Fig. 8A), assuming a gross growth efficiency of 32% (Straile 1997). Mixotrophs were assumed to be only 52% as efficient at fixing inorganic carbon as strict autotrophs (Hansen et al. 2000, Adolf et al. 2006), as determined by calculating their share of the total nanoplankton PP. Estimates of individual photosynthetic rates (Fig. 8B) showed that phototrophy made a more important contribution to the carbon acquisition of MNF in spring (Fig. 8C). In spring 2006 (OMMIX 2), when PNF and MNF peaked around the river plume, photosynthesis contributed little to the cell carbon balance in this area ( $< 1 \text{ pg C cell}^{-1} \text{ h}^{-1}$ ) (Fig. 8B). Heterotrophic carbon acquisition by cell ingestion varied widely throughout the study period, with the highest ingestion rates in spring 2006 and winter 2007 during OMMIX 1 and 4, respectively ( $\sim 2$  to  $4 \text{ pg C cell}^{-1} \text{ h}^{-1}$ ) (Fig. 8A). However, on average, in winter, heterotrophy by MNF accounted for more than 80% of the carbon acquisition (Fig. 8C).

Integrated carbon ingestion by mixotrophic and heterotrophic forms showed that the highest carbon removal by mixotrophs occurred in spring 2006 at Stn RV ( $27 \text{ mg C m}^{-2} \text{ d}^{-1}$ ) and in spring 2007 at Stn UW ( $15 \text{ mg C m}^{-2} \text{ d}^{-1}$ ). Similarly, heterotrophic forms also showed maximum carbon removal in spring. Nonetheless, the trend was the opposite, with maxima observed in spring 2007 at Stn RV ( $123 \text{ mg C m}^{-2} \text{ d}^{-1}$ ) and spring 2006 at Stn UW ( $133 \text{ mg C m}^{-2} \text{ d}^{-1}$ ) (Table 3).

Table 2. Spearman's rank correlation analyses between ingestion by mixotrophic (MNF) and heterotrophic (HNF) nanoflagellates and environmental variables. PAR: photosynthetically active radiation. \*: highly significant correlation. Assigned weights  $> 0.3$  are marked in **bold**.  $n = 44$ . Data were  $\log(x + 1)$  transformed

Physical/ chemical variable	Trophic status/size class			
	MNF $< 5 \mu\text{m}$	MNF $5\text{--}20 \mu\text{m}$	HNF $< 5 \mu\text{m}$	HNF $5\text{--}20 \mu\text{m}$
Temperature	0.28	0.18	0.18	0.14
Salinity	-0.002	-0.15	0.11	-0.21
Oxygen	-0.12	-0.13	-0.07	0.18
PAR	-0.02	-0.27	0.26	0.06
$\text{NO}_3^-$	-0.07	-0.08	0.08	-0.24
$\text{NO}_2^-$	0.14	0.20	-0.02	0.11
$\text{PO}_4^{3-}$	-0.09	0.03	-0.21	-0.18
N:P	-0.10	0.10	<b>-0.36*</b>	-0.05
Bacterial abundance	<b>0.65*</b>	<b>0.50*</b>	<b>0.47*</b>	0.18
Bacterial production	<b>-0.38*</b>	-0.28	0.003	-0.08
River flow	0.03	0.12	-0.14	0.22

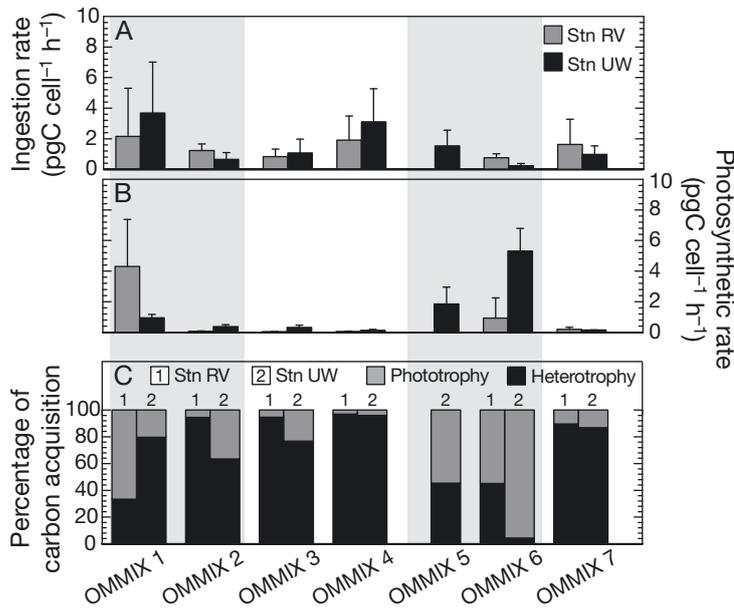


Fig. 8. (A) Individual ingestion rate (mean  $\pm$  SD; pg C cell<sup>-1</sup> h<sup>-1</sup>), (B) individual photosynthetic rate (mean  $\pm$  SD; pg C cell<sup>-1</sup> h<sup>-1</sup>), and (C) percentage of carbon acquisition, through phototrophy vs. heterotrophy during each sampling campaign (OMMIX 1–7) in the river plume area (Stn RV) and off Coliumo Bay (Stn UW). Shaded areas represent spring and white areas winter sampling periods

Table 3. Integrated carbon ingestion (mg C m<sup>-2</sup> d<sup>-1</sup>) by mixotrophic (MNF) and heterotrophic (HNF) nanoflagellates in the upper 20 m depth layer of the study area in spring and winter at both sampling stations

	— Mixotrophs —			— Heterotrophs —		
	MNF < 5 $\mu$ m	MNF 5–20 $\mu$ m	Total	HNF < 5 $\mu$ m	HNF 5–20 $\mu$ m	Total
<b>Stn RV</b>						
Spring 2006	25	2	27	28	36	64
Winter 2007	4	3	7	40	14	54
Spring 2007	1	2	3	111	12	123
Winter 2008	5	3	8	18	16	34
<b>Stn UW</b>						
Spring 2006	4	3	7	85	48	133
Winter 2007	7	5	12	23	20	43
Spring 2007	4	11	15	12	60	72
Winter 2008	1	1	2	18	14	32

## DISCUSSION

### Relative importance of different nanoflagellate forms under contrasting seasons

Estimations of the contribution made by MNF bacterivores, in comparison with those of strictly heterotrophic nanoflagellates were, for years, an unknown 'black box' in river-influenced continental shelves

and coastal upwelling carbon and food web models. The present study offers the first evaluation of the role of PNF, MNF, and HNF communities in a river-influenced eutrophic coastal ecosystem. Although limited in terms of spatial coverage, we herein show the spatial heterogeneity between 2 sampling sites influenced by different oceanographic processes (i.e. river runoff and coastal upwelling). Bacterial production at this coastal upwelling area off Central Chile measured in other studies and from other coastal upwelling areas, fit well with our estimations, e.g. from 19.2 to 33.6 mg C m<sup>-3</sup> d<sup>-1</sup> in a coastal upwelling site off Central Chile (McManus & Peterson 1988), 4 to 92 mg C m<sup>-3</sup> d<sup>-1</sup> in Benguela (Lucas et al. 1984), and 2.7 to 22.2 mg C m<sup>-3</sup> d<sup>-1</sup> in the NW Iberian margin (Barbosa et al. 2001). Typically, bacteria have been shown to rapidly respond to the pulsed production cycles typical of upwelling areas (McManus & Peterson, 1988). However, many studies have also shown enhanced bacterial heterotrophic activities in river plume areas worldwide. For instance, Albright (1983) observed the stimulation of bacterial activity when river waters entered the salty surface waters of the Strait of Georgia in British Columbia, Canada. Carlsson et al. (1995) noted that humic substances from river discharge may enhance the abundance of bacterial grazers in river mouths, as reflected in increased bacterial production.

High abundances of pico- and nanophytoplankton have also been reported in river plume areas. The present study evidenced higher nanophytoplankton PP in the river plume area. Similarly, high abundances of picophytoplankton, picoeukaryotes (Liu et al. 2004), and nanophytoplankton (Jochem 2003) have been observed in river plumes. Recently, Christaki et al. (2009) showed that the PNF abundance was 3 to 5 fold higher in the Rhone River plume waters compared to marine waters. Abundances of PNF and HNF were within the ranges reported in other studies of oligotrophic, mesotrophic, and eutrophic marine ecosystems, mostly from 0.1 to 30 cells ml<sup>-1</sup>  $\times 10^3$  (Table 4 and references in there). We did not observe a clear upwelling vs. non-upwelling trend for PNF abundance or biomass, contrasting with the summer peaks of PNF abundance observed by Böttjer & Morales (2007) at a coastal upwelling site off central Chile. In a study of the annual cycle of nanoflagellates in the central Cantabrian Sea, Granda & Anadón (2008) also reported the highest abundance during summer upwelling periods. Although the PNF biomass in our study was substantially higher in spring 2006 (OMMIX 2), it was relatively low in spring 2007. In

Table 4. Abundance of phototrophic (PNF), mixotrophic (MNF), and heterotrophic (HNF) nanoflagellates and bacterivory accounted for by heterotrophic and mixotrophic forms in different environments. FLM: fluorescently labeled micicells; FLB: fluorescently labeled bacteria; SS: standing stock; BP: bacterial production

Location	Trophic status	Season	Trophic mode	Abundance (cells ml <sup>-1</sup> × 10 <sup>3</sup> )	Tracer	Ingestion (cells flag <sup>-1</sup> h <sup>-1</sup> )	% Bacterivory	Source
Aegean Sea	Oligotrophic	Summer	PNF	0.4–1.0				Christaki et al. (1999)
<b>Aegean Sea</b>	<b>Oligotrophic</b>	<b>Summer</b>	<b>MNF</b>	<b>0.2–1.6</b>	<b>FLM</b>	<b>&lt;2</b>	<b>5 % BP</b>	<b>Christaki et al. (1999)</b>
Aegean Sea	Oligotrophic	Summer	HNF	0.05–0.12	FLM	1–4	40 % BP	Christaki et al. (1999)
Mediterranean Sea	Oligotrophic	Summer	PNF	0.6–2.7				Christaki et al. (2001)
Mediterranean Sea	Oligotrophic	Summer	HNF	0.5–1.2	FLB	0.9–4.3		Christaki et al. (2001)
Rhone River plume, Mediterranean Sea	Oligotrophic	Spring	PNF	0.1–29.2				Christaki et al. (2009)
Rhone River plume, Mediterranean Sea	Oligotrophic	Spring	HNF	0.2–2.6	FLB	1.6–3.5	25 % BP	Christaki et al. (2009)
Convergence, East New Zealand	Oligomesotrophic	Summer	PNF					Safi & Hall (1999)
<b>Convergence, East New Zealand</b>	<b>Oligomesotrophic</b>	<b>Summer</b>	<b>MNF</b>	<b>0.5–1.2</b>	<b>FLB</b>	<b>0.1–0.9</b>	<b>40–55 % SS</b>	<b>Safi &amp; Hall (1999)</b>
Convergence, East New Zealand	Oligomesotrophic	Summer	HNF	0.4–0.7	FLB			Safi & Hall (1999)
<b>Bay of Aarhus, Denmark, surface</b>	<b>Mesotrophic</b>	<b>Spring</b>	<b>MNF</b>		<b>FLB</b>	<b>0.1–0.9</b>	<b>86 % SS</b>	<b>Havskum &amp; Riemann (1996)</b>
Uranouchi-Wan, Japan	Eutrophic	Summer	HNF	0.5–73	FLB	4.8–16.9		Fukami et al. (1996)
East Antarctica	Mesotrophic	Summer	HNF	1.6–4.2	FLB	0.06–1.03	9–33 % BP	Leakey et al. (1996)
East Sea, Korea	Oligotrophic	Annual	HNF	0.4–7.5	FLB	1.5–8		Cho et al. (2000)
<b>Ross Sea, Antarctica</b>	<b>Mesotrophic</b>	<b>Spring</b>	<b>MNF</b>	<b>&lt;0.2</b>	<b>FLB</b>			<b>Moorthi et al. (2009)</b>
<b>Blanes Bay, NW Mediterranean Sea</b>	<b>Oligotrophic</b>	<b>Annual</b>	<b>MNF</b>	<b>1.5–3</b>	<b>FLB</b>	<b>1–1.5</b>	<b>35–65 % SS</b>	<b>Unrein et al. (2007)</b>
Cantabrian Sea	Mesotrophic	Annual	PNF	0.2–4.2				Granda & Anadón (2008)
<b>Cantabrian Sea</b>	<b>Mesotrophic</b>	<b>Annual</b>	<b>MNF</b>	<b>0–0.14</b>				<b>Granda &amp; Anadón (2008)</b>
Cantabrian Sea	Mesotrophic	Annual	HNF	0.04–0.9				Granda & Anadón (2008)
River-influenced upwelling area, Chile	Eutrophic	Winter	PNF	0.5–1.1				Present study
River-influenced upwelling area, Chile	Eutrophic	Spring	PNF	0.7–8.8				Present study
<b>River-influenced upwelling area, Chile</b>	<b>Eutrophic</b>	<b>Winter</b>	<b>MNF</b>	<b>0.02–0.06</b>	<b>FLB</b>		<b>23–51 % BP</b>	<b>Present study</b>
<b>River-influenced upwelling area, Chile</b>	<b>Eutrophic</b>	<b>Spring</b>	<b>MNF</b>	<b>0.02–0.2</b>	<b>FLB</b>		<b>1–6 % BP</b>	<b>Present study</b>
River-influenced upwelling area, Chile	Eutrophic	Winter	HNF	3.2–7.8	FLB		>100 % BP	Present study
River-influenced upwelling area, Chile	Eutrophic	Spring	HNF	0.8–3.5	FLB		24–29 % BP	Present study

contrast, nanophytoplankton PP always showed a clear seasonal trend, being higher in spring. The lack of a clear seasonal trend in PNF abundance and biomass implies that, despite higher growth rates in spring–summer (i.e. high PP), the PNF abundance was tightly controlled by grazers (e.g. ciliates and dinoflagellates grazing was estimated during the same field surveys by Vargas & Martínez 2009). These grazers fail to reveal a contrasting seasonal trend between winter and spring.

During the present study, average MNF abundances were <200 cells ml<sup>-1</sup>, a close fit to those values reported by Granda & Anadón (2008) for the Cantabrian Sea and by Moorthi et al. (2009) for the Ross Sea, Antarctica, both mesotrophic environments (Table 4). However, the use of similar FLB uptake measurements in more oligotrophic environments (Aegean Sea, NW Mediterranean Sea, East New Zealand) has revealed larger MNF abundances, from 200 to 3000 cells ml<sup>-1</sup> (e.g. Christaki et al. 1999, Safi & Hall 1999, Unrein et al. 2007).

An important preliminary implication of these results is that MNF seem to be lower in terms of abundance in highly productive coastal environments, as compared to more mesotrophic and/or oligotrophic environments (e.g. Mediterranean Sea). The relative abundance of MNF as a percentage of total bacterivorous and photosynthetic nanoflagellates in our study averaged from 1 to 6% and from 2 to 6%, respectively, which contrasts with the higher contribution found in other studies. For instance, Moorthi et al. (2009) reported that MNF comprised 8 to 42% of bacterivorous nanoflagellates in the water column of Ross Sea. However, if we consider the feeding-inactive

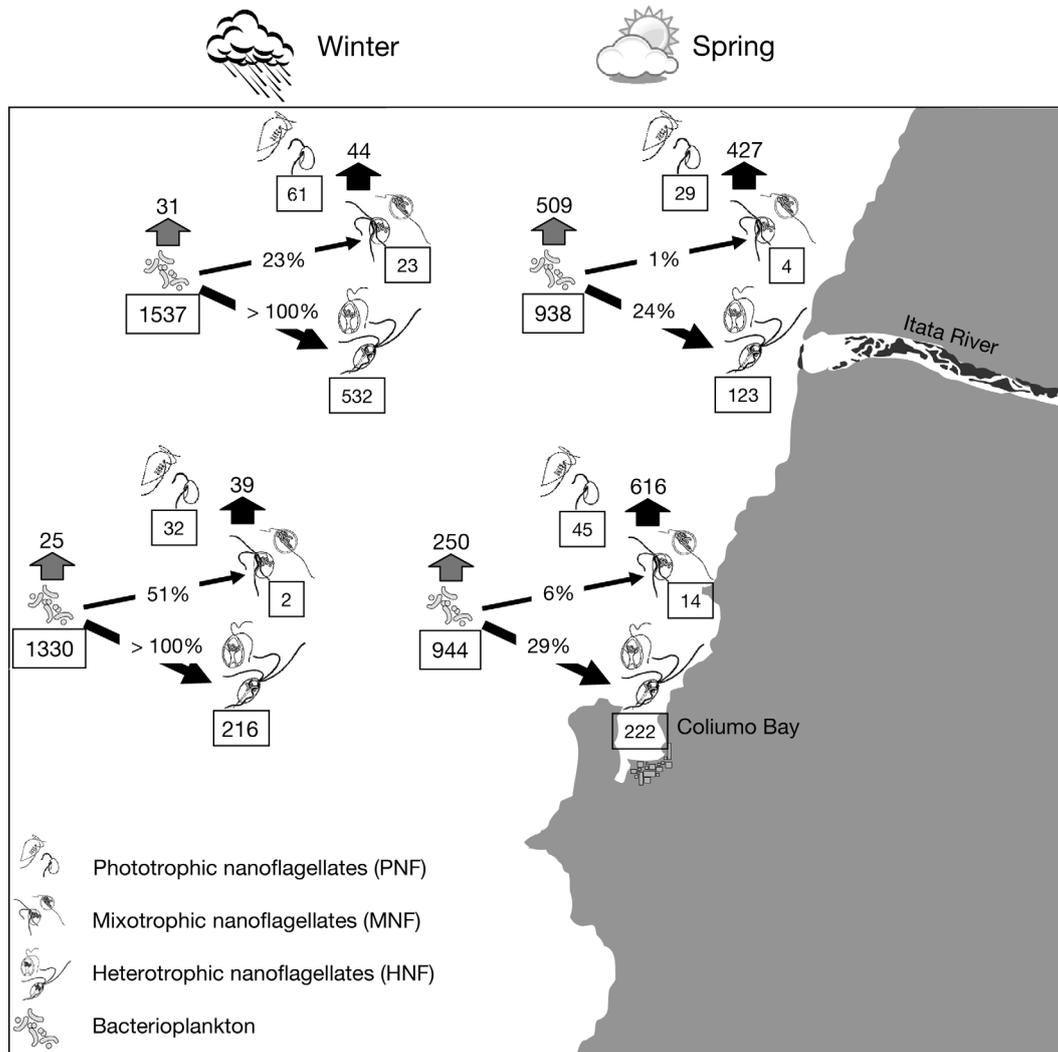


Fig. 9. Carbon fluxes in winter and spring 2007 between bacterioplankton and nanoflagellates (PNF, MNF, HNF), and for the Itata River plume area (Stn RV) and off Coliumo Bay (Stn UW). Numbers inside the boxes show upper 20 m depth-integrated biomass for each group ( $\text{mg C m}^{-2}$ ); numbers above the wide grey and black arrows show bacterial production (BP) and nanophytoplankton primary production (PP) (both in  $\text{mg C m}^{-2} \text{ d}^{-1}$ ), respectively. The percentage in parentheses (%) on the thin arrows represents the percentage of BP consumed per day

cryptophytes in the mixotrophic compartment (i.e. actively grazing HNF are defined as those ingesting any bacteria during uptake experiments; Cho et al. 2000), we may have misclassified a large fraction of the actual MNF as PNF and affected our ability to identify seasonal trends in mixotrophic behavior. In fact, inactive grazers are a common phenomenon in feeding experiments (e.g. Schmidtke et al. 2006), leading to underestimations of mixotroph abundance and overestimations of phototrophic forms. Furthermore, the mixotrophs were identified based on their FLB uptake, and many studies in the literature report discrimination against FLB compared to

living bacteria (e.g. Landry et al. 1991), and in particular against motile bacteria (González et al. 1993). Our data also showed that HNF reached a higher biomass than strictly plastidic flagellates and were numerically dominant in terms of biomass in winter around the river plume. Relatively higher HNF abundance and biomass in river plumes may have been conditioned by higher bacterial production occurring there. Consequently, the dominance of river-induced conditions in winter (e.g. stratification, nutrients, DOM fluxes) may also lead to a sustained high heterotrophic and pico- and nanophytoplankton biomass.

### Environmental factors affecting nanoflagellate abundance and bacterivory

Carbon balances were calculated by assuming that mixotrophs were only 50% as efficient at fixing inorganic carbon as strict autotrophs (Skovgaard 1999, Adolf et al. 2006). This assumption was based on an average from values reported in studies dealing with individual photosynthetic measurements in dinoflagellates, since to our knowledge, there are no reports on individual photosynthetic rates measured simultaneously for both MNF and PNF. For instance, Hansen et al. (2000) found that the mixotrophic dinoflagellate *Fragilidium subglobosum* can reduce its photosynthetic performance to about 40% of that found in phototrophic cultures. Photosynthetic performance in *Karlodinium micrum* was around 52% lower during mixotrophic growth than during autotrophic growth (Adolf et al. 2006). Finally, Skovgaard (1999) also found that photosynthetic rate was reduced by approximately 50% in mixotrophic dinoflagellates as compared with cells growing strictly photosynthetically. Marshall & Laybourn-Parry (2002) also calculated individual photosynthetic rates and grazing and photosynthesis contributions to the carbon balance in ambient mixotrophic cryptophytes communities by using an approach similar to ours. However, Marshall & Laybourn-Parry (2002) did not report any information regarding how primary production was shared between MNF and PNF, and the photosynthetic efficiency they considered for photosynthesis estimates in mixotrophic forms. Consequently, we assume that small flagellates could reduce their autotrophic efficiency at least for an average of 50% as done by larger flagellate organisms. Of course, this is a gross estimation assuming a similar efficiency for small flagellates, but additionally, we have not considered that contributions of photosynthesis and phagotrophy also depend on the type of mixotrophy (Stickney et al. 2000). Indeed these estimates should be an important issue for research on the role of mixotrophic nanoflagellate communities in marine ecosystems. Based on these assumptions, carbon balances for mixotrophs showed that, in winter, heterotrophy could contribute most of carbon acquisition (from ~75 to 95% heterotrophic carbon gain). It also makes sense that phototrophy would become less important in river plume waters in winter, as these waters typically carry sediments from river runoff in this region (Araneda et al. 2009), with the subsequent reduction in light availability at depth. Therefore, although dissolved  $\text{PO}_4^{3-}$  concentrations were usually low in the surface layer, we as-

sumed that the main benefit of mixotrophy (observed in winter) was carbon and energy acquisition in an environment of reduced PAR. This agrees with other reports of MNF from other coastal ecosystems with reduced PAR (e.g. Marshall & Laybourn-Parry 2002, Hammer & Pitchford 2006, Czypionka et al. 2010).

Water temperature and prey abundance are usually some of the most important factors regulating HNF phagotrophic activity (Choi 1994, Vaqué et al. 1994). Positive correlation between bacterial abundance and MNF and HNF ingestion evidence this strong dependence (Table 2). However, bacterial abundance is not the best measure of prey availability since it is well known that HNF may use alternative food resources other than heterotrophic bacteria, e.g. picocyanobacteria, picoeukaryotes, or detritus, although they usually do so at lower rates (Kuosa 1991, Sanders et al. 2000, Sherr & Sherr 2002). Ingestion by small HNF was also negatively correlated with N:P ratio, which suggests an increase in bacterivory activity under  $\text{PO}_4^{3-}$  low conditions. Support for this explanation is also found in a study by Nygaard & Tobiesen (1993) who also found higher HNF ingestion rates in P-limited conditions. In addition to bottom-up factors, HNF abundance could also be controlled by top-down processes. For instance, Vargas & Martinez (2009), evaluated the grazing impact of ciliates and dinoflagellates on nanoflagellate communities during the same experimental and field surveys as in the present study. Those authors clearly showed maximum carbon ingestion by dinoflagellates and ciliates while grazing on nanoflagellates in spring. Furthermore, the ingestion rate of small copepods evaluated during the same surveys evidenced that HNF may constitute an important component of copepod diet in spring (C. A. Vargas unpubl. data), resulting in a strong top-down control on HNF populations, similar to that found in other coastal ecosystems (Goericke 2002).

### Mixotrophic nanoflagellates in carbon and food web models of coastal ecosystem

To the best of our knowledge, this is the first comparative measurement of bacterivory rates by the HNF and MNF inhabiting one of the world's most productive, river-influenced, coastal upwelling ecosystems. Since the photic layer ranged between 20 and 25 m depth during our study, biomass and rate measurements were integrated between 0 and 20 m depth to construct a mean carbon budget for each sampling season in 2007 and to evaluate the role that

mixotrophic forms may play in the microbial food web of a productive, river-influenced coastal upwelling site. Calculations of grazing pressure showed that MNF could be capable of removing between 23 and 51 % BP d<sup>-1</sup> in winter and from 1 to 6 % BP d<sup>-1</sup> in spring, with a higher grazing impact at Stn UW, since BP and bacterial biomass were substantially lower at this site than in the river plume area, which is commonly subsidized by terrestrial organic sources (Vargas & Martínez 2009; Table 1). Consequently, in the microbial food web, HNF controlled BP by removing a major fraction of BP in spring (24 to 29 % BP d<sup>-1</sup>) and >100 % BP d<sup>-1</sup> in winter, mainly due to the low BP at that time (Fig. 9). Finally, phototrophic and mixotrophic nanoflagellate forms contributed 44 mg C m<sup>-2</sup> d<sup>-1</sup> of PP in winter at Stn RV, as compared to only 39 mg C m<sup>-2</sup> d<sup>-1</sup> at Stn UW. In winter 2007, nanoplankton PP contributed between 20 and 40 % of the total PP around the river plume (Iriarte et al. in press).

The specific grazing rates estimated in the present study are in the same range as those reported for other coastal areas worldwide (see Table 4 and references therein), where heterotrophic and mixotrophic nanoflagellates accounted for a substantial proportion of total bacterivory. For instance, Leakey et al. (1996) reported that 9 to 33 % of the BP was grazed daily by HNF in the Ross Sea, Antarctica. Safi & Hall (1999) estimated that 40 to 55 % of the bacterioplankton standing stock was grazed by MNF, whereas Christaki et al. (1999) reported that ~40 % BP d<sup>-1</sup> was grazed by HNF and only 5 % BP d<sup>-1</sup> by MNF. Similar to these reports, bacterivory in our study also appeared to be dominated by HNF rather than MNF. Previous investigations in coastal upwelling areas have only considered the top-down effect of HNF grazing. Vargas et al. (2007) constructed a conceptual model of the carbon flow from the estimated PP on the continental shelf off central Chile. With this model, the authors clearly showed that HNF were able to graze >90 % BP d<sup>-1</sup> in winter, whereas in summer, HNF were able to graze ~70 % BP d<sup>-1</sup>.

It is clear from many studies worldwide that grazing by HNF and MNF can control bacterioplankton biomass by removing a substantial fraction of BP (Zubkov & Tarran 2008). Moreover, nanoflagellates are widely recognized as important prey items for protozooplankton, including dinoflagellates and ciliates, in coastal areas (e.g. Vargas & Martínez 2009), which makes these protozoans direct competitors with copepods and other mesozooplankton (Vargas et al. 2007). Through this link, bacterivory by MNF and HNF can make autochthonous organic carbon

from algal exudates and allochthonous organic carbon from riverine DOC indirectly available to zooplankton and upper trophic levels (see Calbet & Saiz 2005). MNF might also account for ~1 to 5 % of the nanoplankton PP, a percentage that needs to be incorporated into food web models of productive coastal ecosystems. Consequently, our results reinforce the idea that, through grazing by MNF and HNF, the microbial pathway may increase yields of both terrestrial (detrital DOC) and photosynthetically fixed organic carbon, supporting higher secondary production of larger metazoans than would be expected from a simple herbivore-dominated food chain in coastal areas.

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