INTRODUCTION

The 'deep sea' includes the waters and sediments of the ocean interior beneath approximately 200 m depth, representing more than 65% of the Earth’s surface and ca. 95% of the global biosphere (Gage & Tyler 1991). Despite the huge dimensions and ecological importance of deep-sea habitats, our knowledge on deep-sea microbiology and its role in ecosystem functioning is still limited. Current estimates report that only 5% of the deep oceans have been explored in detail so far, and that less than 0.001% has been sampled and described in terms of microbiology and biodiversity (Rex et al. 2006, Wei et al. 2010). Even less is known about the functioning of, and the interactions between, biodiversity and ecological processes in the deep seas (Danovaro et al. 2014, Mora et al. 2014). The development of advanced instrumentation for observing, mapping and sampling the seabed and all of its features has enabled scientists to develop a picture of the deep-sea environment as a highly dynamic geo- and biosphere. Besides the vast deep-sea flat and muddy plains and their subsurface, the ocean seafloor is characterized by high habitat complexity and heterogeneity (Ramirez-Llodra et al. 2010). A variety of highly diverse landscapes have been recently described, including canyons, seamounts, ridges, fur-
rows, deep-water coral reefs, cold seeps, pockmarks, mud volcanoes, carbonate mounds, fractures and trenches that host rich and highly diversified microbial assemblages. Compared with the water column, the ocean seafloor provides a wider area of solid surfaces composed of biogenic, terrigenous, volcanogenic particles, with heterogeneous pore spaces, as well as a higher concentration of organic matter per unit volume (Danovaro et al. 2014).

Along with the export of organic material from the photic zone, geological processes transport chemical energy (hydrogen, methane, hydrogen sulphide, iron) up to the seafloor from the subsurface below. At the sites on the seafloor where chemical energy is particularly high, rich and diverse microbial assemblages are typically found, such as in hydrothermal vents and cold seeps (Cavanaugh 1985, Jørgensen & Boetius 2007). However, microorganisms inhabiting the subsurface of the ocean floor live in conditions of extreme energy limitation, with mean generation times up to thousands of years (Hoehler & Jørgensen 2013), depending only on the small amount of the organic carbon buried in deep-sea sediments (Lutz et al. 2007) and on chemoautotrophic processes (D’Hondt et al. 2002, 2004).

Although the scientific discovery of the ubiquity and enormous abundance of microorganisms dates back to the 1860s, microbes inhabiting the deep sea started to be studied in detail only in the 1950s, along with the appearance of the first historical monographs on marine microbiology (e.g. Benecke 1933). In the 1950s, deep-sea samples that were gathered on the Danish ‘Galathea’ Deep-Sea Expedition from depths of more than 10 000 m were shown to host millions of viable Bacteria per gram of sediment, confirming that pressures up to more than 1000 times higher than atmospheric pressure do not exclude the possibility for life to evolve (ZoBell & Morita 1959). Subsequent studies conducted in the 1980s and 1990s that were aimed at quantifying the microbial assemblages demonstrated their activity in the deep ocean interior (Jannasch & Taylor 1984, Alongi 1990, Deming & Baross 1993, Yayanos 1995). Most deep-sea microorganisms are extremely difficult to cultivate, due to their slow in situ growth rates and their likely strict adaptation to extreme environmental conditions (oligotrophy, low temperatures and high in situ pressures or anoxia). The number of pure cultures that are available is limited to a few hundred species, and the cultivation of microorganisms under high pressure has been achieved only in a limited number of cases (Kato 2011).

**ECOLOGICAL ROLE OF MICROBES IN THE DEEP SEAФLOOR**

In recent years, enormous progress has been made in the investigation of deep-sea microbes and their functional role, and an increasing portion of the deep seafloor has been sampled (Fig. 1). The analysis of these samples has revealed that deep
surface sediments, at all depths, contain 10- to 10 000-fold more cells per unit volume than productive ocean-surface waters (Whitman et al. 1998). It is thus now evident that prokaryotes represent the largest pool of biomass in the world oceans (Rex et al. 2006, Wei et al. 2010), and that they contribute to the biogeochemical cycles in several fundamental ways, thus playing a key role in deep-sea ecosystem functioning (Nealson 1997, Mason et al. 2009, Arndt et al. 2013). Besides prokaryotes, other microscopic biological entities, viruses, are also known to play a major role in nutrient cycling and transfer of energy to higher trophic levels, both in the water column and sediments (Suttle 2005, 2007, Danovaro et al. 2008a,b). Not being able to self-replicate, viruses invade other organisms and use their cells’ machinery to propagate (Suttle 2005). They can infect all known life forms in the oceans, from the smallest marine microbes to the largest mammals (Rohwer & Thurber 2009). Since prokaryotes are the predominant life forms in the deep oceans in terms of biomass, the interactions between viruses and members of the 2 prokaryotic domains (i.e. Bacteria and Archaea) are considered among the most relevant processes driving ecosystem functioning on the global scale (Proctor & Fuhrman 1990, Fuhrman 1999). The prokaryotic biomass represents a potentially enormous food source for higher trophic levels in deep-sea food webs (Dell’Anno & Danovaro 2005). However, since heterotrophic consumers belonging to meio-, macro- and mega-fauna decrease exponentially with increasing water depth (Rex et al. 2006), the potential role of viruses in controlling microbial components increases in the deep ocean interior (Danovaro et al. 2008b). Viral ecology is a crucial field in deep-sea microbiology, largely neglected thus far, and is only now starting to be included in global models of carbon, nitrogen and phosphorus cycling and nutrient fluxes (Jover et al. 2014, Dell’Anno et al. 2015). By killing their hosts, viruses can transform the living biomass into organic detritus (mostly dissolved organic matter, DOM); this can then be used again by other microbes, stimulating their growth. This process has been termed ‘viral shunt’ (Suttle 2005, 2007) and, depending on the balance between the abatement of living cells and the stimulation of the metabolism of the survivors, it can either fuel prokaryotic heterotrophic and autotrophic production by releasing DOM and enhancing nutrient regeneration pathways, or it can also decrease the efficiency of the carbon transfer to higher trophic levels. The debate on this topic is still open among marine ecologists.

**MICROBIAL ABUNDANCE, BIOMASS AND DISTRIBUTION IN THE OCEAN SEAFLOOR**

Whitman et al. (1998) made a global estimate of the number of prokaryotic cells that are present in the oceans, from the sea surface to the subsurface sediments. They concluded that deep-sea microbes represented the ‘hidden majority’ of all life forms, comprising between 50 and ~80% of the Earth’s total microbial biomass and between 10 and 30% of the Earth’s total living biomass (Teske 2005). Studies conducted on sediment cores, within the frame of the Integrated Ocean Drilling Program expeditions, provided information on the number, biomass and activity of prokaryotes below the sediment surface, revealing the presence of metabolically active microorganisms several hundred metres below the ocean floor (Cragg et al. 1990, Parkes et al. 2000, D’Hondt et al. 2002, 2004, Teske 2005). A more recent meta-analysis on the available information from the ocean surface to the deep-sea subsurface sediments indicates that the average cell numbers and activities decrease exponentially with depth (both in the water column and in the sediment; Lloyd et al. 2013). Archaea have been assumed for a long time to be a minor portion of the total microbial assemblages in the oceans and in the seafloor. Subsequent works highlighted that archaeal abundance, not only in deep-sea waters beneath 1000 m depth but also in subsurface sediments, can equal bacterial abundance or even dominate the assemblages (Karner et al. 2001, Schippers et al. 2005, Biddle et al. 2006, Schippers & Neretin 2006, Lipp et al. 2008, Lloyd et al. 2013, Xie et al. 2013).

However, information on the relative importance of Bacteria and Archaea and their biomass in marine sediments is still limited (Kallmeyer et al. 2012, Xie et al. 2013), and even less is known about the contribution of viruses to the total microbial benthic biomass (Anderson et al. 2013, Engelhardt et al. 2014). The first investigations of global trends in the distribution and abundance of microbial cells showed that the abundance of cells in surface sediments is usually related to the input of fresh detritus, rather than to any other environmental variable (Deming & Baross 1993, Smith et al. 1997, Boetius & Damm 1998). It was generally assumed that the abundance and biomass of the microbial components are controlled by the quantity and quality of the available organic substrates. This applies to most biological components in the deep sea (e.g. meiofauna macrofauna, megafauna), and since the availability of food and resources typically decreases with increasing water depth (Danovaro et al. 1998), the biomass of all fau-
nal components decreases exponentially (Rex et al. 2006). However, in the same work, Rex et al. (2006) demonstrated that benthic Bacteria only barely decreased with increasing water depth.

Viruses inhabiting deep-sea ecosystems are quantitatively far more important than was assumed only 2 decades ago. Now, we know that viruses are the most abundant biological entities of the oceans, outnumbering prokaryotes by at least 1 order of magnitude (Suttle 2005, 2007). Current estimates are in the order of $10^{30}–10^{31}$ viruses in the water column (Suttle 2005, 2007), and available information demonstrates that viruses are also highly abundant and active in deep-sea sediments (Middelboe et al. 2006, 2011, Danovaro et al. 2008b).

Analysis of the spatial patterns of prokaryotic abundance and biomass suggests that they do not decrease significantly with increasing water depth as observed for mega-, macro- and meiofauna, (Rex et al. 2006). The meta-analysis we conducted on available benthic data confirms that prokaryotic and viral abundances are high at all depths in the deep sea, from the shelf-break down to the abyssal depths (Danovaro et al. 2008a,b, Siem-Jørgensen et al. 2008). Results presented here indicate that while the abundance and biomass of Bacteria showed a decreasing pattern (negative slope coefficients) with increasing water depth, the abundance and biomass of viruses and Archaea did not decrease (Fig. 2A,B, Table 1). In contrast, a decreasing pattern along the vertical profile of sediment was observed for all microbial components investigated (Fig. 2C–E). If the available data provide the first insights on the bathymetric patterns of benthic microbial abundance and biomass, more effort is needed to analyse these trends in more detail in each water depth realm (e.g. meso-, bathy- and abyssopelagic).

Indeed, our data suggest a significant increase in microbial abundance and biomass (especially for Archaea) at depths from 4000 to 6000 m (Fig. 2A,B). However, the number of analysed samples and level of replication in this depth range remain insufficient to draw significant and definitive conclusions.

![Fig. 2. Bathymetric patterns for (A) abundance and (B) biomass of Bacteria, Archaea and viruses in surface deep-sea sediments (top 1 cm), and patterns for (C) bacterial, (D) archaeal and (E) viral abundance along the vertical profile of deep-sea sediments. Cell and virus counts were carried out by epifluorescence microscopy (i.e. CARD-FISH for Bacteria and Archaea, Ishii et al. 2004, Molari & Manini 2012, SYBR Green I staining for viruses, Danovaro et al. 2001, Noble & Fuhrman 1998). For the determination of the prokaryotic biomass, prokaryotic size was estimated using a micrometer ocular (as maximal length and width), assigning prokaryotic cells into different size classes; prokaryotic bio-volumes were then calculated after inter-calibration with measurements conducted with scanning electron microscopy and converted into carbon content assuming 310 fg C µm$^{-3}$ as a conversion factor (Fry 1988). For viral biomass, the conversion factor of 0.2 fg C virus$^{-1}$ was used (Wilhelm & Suttle 1999).](image-url)
Multiplying the average value of bacterial, archaeal and viral abundance or biomass obtained within a discrete bathymetric range (using 200 m intervals) by the portion of global surface corresponding to that bathymetric interval, we found a marked increase in the abundance and biomass of all microbial components with increasing water depth (Fig. 3A,B), indicating that the deep seas host the vast majority of the benthic biosphere.

Viruses represent the most abundant biological entities in deep-sea surface sediments, at all water depths investigated so far, on average accounting for 67% of the overall microbial abundance (Fig. 4A). However, the small size of viruses makes these biological components far less important in terms of biomass than that of Bacteria and Archaea, contributing only 0.5 to 4% (on average 1.2%) to the total microbial biomass (Fig. 4B) if the conversion factor of 0.2 fg C virus$^{-1}$ is used (Wilhelm & Suttle 1999). Based on a biophysical model and different assumptions, Jover et al. (2014) suggested that the carbon content and thus the contribution of marine viruses to the total microbial biomass could be up to 10 times lower. As such, current research perspectives on marine viruses mainly focus on virus–host interactions influencing global biogeochemistry and virus–host evolution, rather than standing stocks of viral biomass per se (Rohwer & Thurber 2009, Jover et al. 2014, Dell’Anno et al. 2015). The analysis of the relative contribution of Bacteria, Archaea and viruses to the total microbial abundance and biomass in each bathymetric range did not show clear patterns with increasing water depth (Fig. 5A,B). By integrating the values reported here of abundance and biomass of the 3 microbial components from the top 1 cm of the sediment down to 50 cm depth, we estimated the abundance and biomass of Bacteria, Archaea and viruses in surface deep-sea sediments on a global scale. Overall, surface deep-sea sediments host $3.5 \times 10^{28}$ bacterial cells, $1.4 \times 10^{28}$ archaeal cells and $9.8 \times 10^{28}$ viruses, corresponding to 1.29, 0.35 and 0.02 Pg C, respectively.

**IMPLICATIONS FOR GLOBAL C CYCLING AND PROVISION OF GOODS AND SERVICES**

Photoautotrophic production in the open ocean accounts for over 80% of the global oceanic primary production (del Giorgio & Duarte 2002). Organic particles originating from the surface waters are exported down to
the deep sea at a rate of approximately 10 Pg C yr\(^{-1}\) (Duce et al. 2008), but only a small fraction of the particulate organic carbon produced in the ocean surface reaches the ocean seafloor (Muller-Karger et al. 2005, Dunne et al. 2007), mostly composed of organic particles resistant to degradation (Middelburg & Meysman 2007). Indeed, within the organic matter pool settling on the seafloor, only a minor portion is represented by compounds immediately bioavailable for heterotrophic consumers (Fabiano et al. 1993, 2001, Danovaro et al. 1999). With the average prokaryotic (i.e. bacterial plus archaeal) abundance reported here for marine sediments (approximately 5 \(\times\) 10\(^8\) g\(^{-1}\), range: 2 \(\times\) 10\(^8\) to 2 \(\times\) 10\(^9\) g\(^{-1}\)) and considering the amounts of total organic carbon generally observed in open ocean sediments (ca. 5 mg C g\(^{-1}\), Pedersen & Calvert 1990, Arndt et al. 2013), bacterial and archaeal biomass together contributes, on average, ca. 2\% to the organic carbon pool in surface sediments. Estimates of pools based on CARD-FISH analyses are used conservatively, as these values are generally underestimated when compared to results on total prokaryotic abundance based on SYBR Green staining (Molari et al. 2013).

Prokaryotes are known to duplicate faster than eukaryotic components, resulting in a turnover of the prokaryotic carbon biomass faster than that of other benthic organisms. Based on present estimates of the overall bacterial and archaeal biomass, and assuming an average turnover time of 20 d for prokaryotes inhabiting the deep ocean interior (which is a modal value of estimates available in the literature, where the range is 2−100 d; Herndl et al. 2005, Danovaro et al. 2008b, Molari et al. 2013), the biomass production in surface deep-sea sediments of the world oceans is equivalent to 0.88 Pg C yr\(^{-1}\) for \textit{Bacteria} and 0.24 Pg C yr\(^{-1}\) for \textit{Archaea}. Hence, if these estimates are correct, the global organic carbon inputs to the ocean seafloor (~1 Pg C yr\(^{-1}\); Muller-Karger et al. 2005, Dunne et al. 2007) are apparently insufficient to sustain microbial metabolism and growth since a large fraction of carbon is also utilized for respiration processes (50–80\% of the total carbon consumed). Nevertheless, a fraction of organic carbon (approximately 1–10 kg C m\(^{-3}\) at the seafloor; Mayer 1993, Arndt et al. 2013), not consumed by heterotrophic organisms (Hedges & Keil 1995), is buried in the sediments. This apparent paradox of the imbalance between the organic carbon supply and its consumption can be partly explained taking into account the microbial recycling of carbon (Strayer 1988), or with a potential underestimation of the allochthonous inputs of organic carbon (Arístegui et al. 2009, Burd et al. 2010) and/or of the additional sources of organic carbon that are produced \textit{in situ} through chemosynthetic processes (Herndl et al. 2008, Reinthaler et al. 2010, Middelburg 2011, Molari et al. 2013, Herndl & Reinthaler 2013). However, using the model proposed by Strayer (1988), we calculated that the observed overall carbon production can be sustained only for ecological (carbon conversion) efficiency >45\%. This value appears very high for oligotrophic ecosystems, such as the deep seas (del Giorgio & Cole 1998, Dunne et al. 2007). Moreover, recent evidence suggests that intracellular dissolved material released by viral lysis can be an additional important bioavailable source of organic carbon for sustaining prokaryotic metabolism (Corinaldesi et al. 2007, Danovaro et al. 2008b, Corinaldesi et al. 2014). All of these processes have a profound influence on the budgetary considerations, and are still difficult to model in deep-sea ecosystems.

Prokaryotic heterotrophic production and viral production in surface deep-sea sediments were demonstrated to be high at all
depressions and were significantly correlated (Danovaro et al. 2008b). This is in line with the theoretical dependence of viral replication on host abundance and metabolism. The impact of viruses on benthic prokaryotes led to the estimation of a virus-induced abatement of 80% of the total prokaryotic heterotrophic production in deep-sea sediments, increasing with water depth, from 16 to 63% in coastal surface sediments up to >80% in deep-sea surface sediments (at depths >1000 m), indicating that viruses are the main agents of mortality for prokaryotes in deep-sea sediments worldwide (Danovaro et al. 2008b).

Indeed, viruses are responsible for the abatement of an important fraction of the benthic prokaryotic biomass at fast rates (Danovaro et al. 2008b), releasing, on a global scale, 0.4 to 0.6 Pg C yr⁻¹. The highly bioavailable organic carbon supplied by viral shunting is estimated to sustain 35% of the total benthic prokaryotic metabolism, with important implications for carbon cycling and nutrient regeneration processes (including nitrogen and phosphorus associated with prokaryotic biomass, Danovaro et al. 2008b). Viral infections thus play important roles in the functioning of the benthic ecosystem, by controlling benthic prokaryotic biomass (top down, predatory control), by stimulating prokaryotic metabolism (bottom-up mechanism) and by accelerating biogeochemical processes. Despite the recognized role of viruses in the deep seafloor, only scattered information is available on the composition of the active fraction of standing viruses able to infect Bacteria and Archaea, as well as on the host range of the viruses able to replicate in deep-sea sediments (Yoshida et al. 2013).

The dark portions of the oceans represent the most remote biome on Earth, and they are characterized by a huge, yet-to-be-described microbial diversity (Sogin et al. 2006, Zinger et al. 2011, Mengerink et al. 2014), representing one of the last frontiers in scientific research. Deep-sea microbes play crucial roles in the ecosystem at the global scale, significantly contributing to biogeochemical cycles and the regeneration of organic matter and nutrients which can support food webs and keep the ocean healthy and productive. Moreover, the significance of the value of these almost invisible but essential components of the biosphere extends way beyond ecological aspects. Indeed, deep-sea ecosystems have an enormous potential in the discovery and development of new products of microbial origin, such as pharmaceuticals, molecular probes, enzymes, cosmeceuticals, nutritional supplements and agrochemicals. Deep-sea microbes, and particularly those inhabiting extreme ecosystems such as hydrothermal vents, have been fundamental for the discovery of new bioactive molecules subsequently employed for technological and industrial purposes (Ohta & Hatada 2006, Wu et al. 2006, Cavicchioli et al. 2011). Different enzymes such as DNA polymerases, enzymes involved in ethanol production or others with antioxidant properties, derived from deep-sea Bacteria, are already commercially available. Several secondary metabolites with novel structures and specific bioactivities, which are promising for medical therapy, have been discovered from deep-sea microbes, including alkaloids as therapeutic agents, bone-healing materials and exopolysaccharides with anticoagulant properties (Collec Jouault et al. 2001, Zanchetta et al. 2003, Yang et al. 2013). Evidence is also accumulating on the importance of microbes on the biomineralisation and transformation of deep-sea polymetallic nodules and metal crusts, whose economic value is currently of great industrial and technological interest (Wang et al. 2011). Thus, the knowledge of deep-sea ecosystems and their microbial components is important not only because of their key role in biogeochemical cycles, but also because they provide unique goods and services which can be exploited for human socio-economic development (Thurber et al. 2014). The enormous value of deep-sea ecosystems in terms of exploitable resources highlights the need to establish shared criteria governing their sustainable use and preservation (Mengerink et al. 2014, Van Dover et al. 2014). Human activities such as trawling, deep-sea mining, drilling and dumping (Ramirez-Llodra et al. 2013, Yanagawa et al. 2013, Pusceddu et al. 2014) are profoundly threatening deep-sea ecosystems. Moreover, we must take into account that, if not properly regulated, recent research fields such as deep-sea bio-prospecting (i.e. the screening of deep-sea microbes in search of new products of possible economical interest) could potentially generate impacts, similarly to any other kind of exploitation of the deep sea. To date, a legal framework for the management, sharing of the benefits and access to the resources stored in areas beyond national jurisdictions does not exist (Barbier et al. 2014), but this appears to be of fundamental importance to preserve the deep seas, their astonishing microbial biodiversity and their key role in global ecosystem functioning.

Acknowledgements. This research was conducted within the frame of the EU Project MIDAS (Managing Impacts of Deep-seA reSource exploitation), and supported by the FIRB Project EXPLODIVE and by the Flagship Project RIT-
MARE of the National Research Council of the Italian Ministry Research (MIUR). We thank 3 anonymous reviewers for their useful comments which improved the quality of this manuscript.

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Editorial responsibility: Fereidoun Rassoulzadegan, Villefranche-sur-Mer, France

Submitted: November 17, 2014; Accepted: February 17, 2015
Proofs received from author(s): April 30, 2015