INTRODUCTION

It has become evident that southern and northern polar waters share the same basic food web structure (Andersen 1988, Garrison & Buck 1989, Sherr et al. 1997, Edwards et al. 1998). It has also been demonstrated that Arctic and Antarctic marine pelagic food webs show the same complexity as lower latitude ecosystems (von Bröckel 1981, Garrison & Gowing 1993, Nielsen & Hansen 1995, Rivkin et al. 1996). Research now suggests that protozoa are pivotal members of polar marine food webs (von Bröckel 1981, Hewes et al. 1985, Paranjape 1987, 1988, Andersen 1988, Smetacek et al. 1990). The dominant protozoans include mixo- and heterotrophic ciliates and dinoflagellates (e.g. Levinsen et al. 1999). These groups of organisms are able to consume a wide size range of prey; from bacteria to large diatoms.

To quantify the trophic role of protozoans, much effort has been put into growth measurements (e.g. Stoecker et al. 1983). By comparison, few studies have focused on their response to starvation in natural environments. From the perspective of a protozoon, survival during periods of resource limitation is of fundamental importance. At present, most knowledge comes from laboratory experiments (reviewed by Fenchel & Finlay 1983, Caron et al. 1990). These studies describe how some protozoa drastically reduce their metabolism in response to starvation. The polar winter is the most unfavorable environment for marine primary producers and hence for heterotrophic organisms that depend on...
this food source. However, winter field studies in cold oceans have rarely been conducted (Garrison et al. 1993, Scharek et al. 1994, Rysgaard et al. 1999). Protozoan winter dynamics are poorly understood, although ‘winter’ is the prevailing season at high latitudes.

This study describes the annual succession of ciliates and heterotrophic dinoflagellates and their community structure in Disko Bay on the western coast of Greenland. Sampling throughout the winter enables us to discuss how heterotrophic protists overwinter when phytoplankton production is zero/negligible. To our knowledge, this paper is the first on the annual cycle of ciliates and heterotrophic dinoflagellates in a polar marine ecosystem.

METHODS

Study site and program. From April 1996 to June 1997 a 250 m deep station in Disko Bay, West Greenland (~69° N) was sampled 52 times (Fig. 1). During the ice-free period sampling was carried out from RV ‘Porsild’ (Arctic Station in Qeqertasuaq, University of Copenhagen). When the Bay was ice-covered, dogsleds were used for transport and sampling was done through a hole in the sea ice.

Hydrography and other physical measurements. Vertical profiles of salinity, temperature, and chlorophyll fluorescence were obtained from the surface to just above the bottom with a Seabird CTD system. Water density was calculated from salinity and temperature according to Fofonoff & Millard (1983). Arctic Station, University of Copenhagen, provided data on solar irradiance and ice cover.

Water sampling. Water was collected from 5 to 6 depths in the upper 30 to 60 m of the water column with a Niskin bottle. During part of the winter when the upper water column was mixed, water collection was restricted to 2 and 15 m. More intensive sampling was carried out monthly when an additional 4 depths (50, 100, 150, 200 m) were included to improve the vertical resolution.

Ice communities. In spring, water from the ice-water interface and ice cores (triplicates) were collected and processed according to Garrison & Buck (1986) and Grastrup-Hansen (1998) in order to determine the con-

![Fig. 1. Approximate position of the 250 m deep sampling station in Disko Bay](image-url)
centation of algae living within and immediately below the sea ice.

**Phytoplankton.** Chlorophyll *a* (chl *a*) concentrations in water from all depths were measured spectrophotometrically (Strickland & Parsons 1972) after filtering 3 to 15 l of seawater onto GF/F filters and extracting in 96% ethanol over night (Jespersen & Christoffersen 1987). Phytoplankton carbon biomass of all nano- and micro-sized cells from 2 m depth and from the depth of the fluorescence maximum was estimated based on 2% final concentration (v/v) Lugol's-preserved samples analyzed with an inverted microscope (H. A. Thomsen unpubl. data). Autotrophic and heterotrophic flagellates <20 µm, which were considered primary prey for ciliates, were also counted by epifluorescence microscopy of proflavine stained filters (Haas 1982). Samples were preserved with glutaraldehyde to a final concentration of 1%. A volume-specific carbon weight of 0.13 pg C µm⁻³ was used for both Lugol's-preserved and glutaraldehyde-preserved cells (Hansen et al. 1997).

**Ciliates and heterotrophic dinoflagellates.** Water samples of 325 ml were, as phytoplankton, preserved in acid Lugol's solution immediately and stored in the dark until examination with an inverted microscope. Samples (325 ml) from 2 m and the chl *a* fluorescence maximum were also preserved in a buffered aldehyde mixture for subsequent epifluorescence microscopy (see Levinsen et al. 1999).

Biomass was estimated in Lugol's-preserved water from 2 m, the chl *a* fluorescence maximum (or 15 m) and 30 m. Depending on water column structure and chlorophyll fluorescence distribution, additional samples were occasionally analyzed. On the monthly intensive sampling days, the biomass from 8 to 9 depths between 2 and 200 m was determined. Organisms in 50 to 325 ml aliquots were settled and counted. Winter and deep-water samples with low cell concentrations were pre-settled for a week prior to the removal of the supernatant by siphoning and a final sedimentation of the remaining 100 ml. The supernatant contained an insignificant number of small cells (<5% of settled cells). Surface samples were counted within a week, the other samples were processed after 6 mo.

Cells were identified and measured using the criteria outlined in Levinsen et al. (1999) and converted from biovolume to units of carbon using a conversion factor of 0.13 pg C µm⁻³ (Hansen et al. 1997). The aldehyde-preserved samples intended for trophic separation of dinoflagellates were settled in the dark in a refrigerator and counted on an inverted microscope equipped with a 100 W mercury lamp within 2 d of sampling. A ratio between autotrophic (pigmented) and heterotrophic dinoflagellates was calculated for unidentifiable spherical to subspherical gymnodinoid forms in size groups of 10 µm starting with an equiva-
lent spherical diameter (ESD) of 20 to 30 µm. At least 75 cells were inspected. Heterotrophic dinoflagellate biomass was estimated by multiplication of the autotrophic/heterotrophic cell ratio with the cell concentrations from similar size groups determined from Lugol's counts. During winter, when the cells were scarce, a mean of ratios obtained from 2 sampling dates in the late autumn was applied. The method did not allow for separation between mixotrophic and autotrophic dinoflagellates. Possibly, some cells categorized as autotrophic were mixotrophic. However, the epifluorescence samples were used to distinguish mixotrophic from heterotrophic ciliates. Uniformly red-orange fluorescent cells were considered mixotrophic and green cells heterotrophic. Cells were divided in size classes of 10 µm based on length. At least 75 ciliates were counted per sample. For both ciliates and dinoflagellates it was assumed that heterotrophic, mixotrophic and autotrophic cells were preserved equally well.

By ‘oligotrichous ciliates’ we refer to the subclass Oligotrichia that include the ciliate orders Choreo-trichida and Oligotrichida (Lynn & Corliss 1991). Biomass per m² was calculated by trapezoidal integration over the depth strata 0 to 200 m (Nielsen & Bresta 1984). A linear interpolation from the temporally nearest corresponding depths was used to estimate the biomass at 50, 100 and 200 m when these depths were not sampled. Because the proportion of unidentified autotrophic gymnodinoids was not established at all depths, we present integrated heterotrophic dinoflagellates as ‘total’ values, excluding only the autotrophic species recognized from the Lugol’s counts. This was preferred over the use of interpolation of the established autotrophic/heterotrophic dinoflagellate ratios between samples as these ratios were found to vary unpredictably with depth and time. Thus, depth-integrated values are slightly overestimated and strict heterotrophic dinoflagellates are presented for the surface water and fluorescence maximum depth only.

**Growth and mortality.** Net growth and loss rates were calculated for spring, summer and autumn-winter communities as the slope of the regression line fitted to ln transformed biomass values versus time assuming exponential biomass changes. Data points for the regressions were selected using stepwise analysis to get the best fit to the model; at each step a Pearson correlation coefficient was calculated and the significance of the correlation was determined by a *t*-test.

**Diversity.** Margalef’s index (Margalef 1951) was used to show the seasonal pattern in ciliate and heterotrophic dinoflagellate diversity. This index (*I*ₘ) suggests a simple logarithmic relationship between number of species (*S*) and individuals (*N*): *I*ₘ = *(S – 1)/log*N*. We used the identified morphospecies as the ‘species concept’ to calculate the index.
RESULTS

Hydrography, ice and meteorological conditions

High seasonal variation in solar irradiance characterizes the study site. Polar darkness occurs from late November to late January (Fig. 2a). Snow-covered annual sea ice further lengthens the period with reduced light penetration into the water (Fig. 2b). In 1996, the sea ice broke up on April 26. In 1997, ice formation started in February but due to strong wind episodes, fast ice first developed in March. The fast ice cover lasted until late April. Thereafter it broke up and drifted in Disko Bay for several weeks as in the previous year. Sub-zero temperatures characterized the surface water from late November until May-June (Fig. 2b). An annual temperature maximum of ~6°C was recorded in August.

The water column experienced an annual pattern in structure and stability essentially consisting of 2 elements corresponding to the productive summer period and the unproductive winter period (Fig. 2c). After ice-breakup melt water and solar heating formed a warmer less-dense surface layer. This stratification strengthened and deepened during summer. In the beginning of May 1996 the 33.0 PSU isoline was located near the surface; in August it was at 30 ± 1 m (mean ± SD of 5 station visits). The mean temperature in the upper 5 m during this period increased from <–1.7 to 5.9 ± 0.4°C. Following the stratified summer conditions there was a short transition to mixed winter conditions in September (Fig. 2c). Increased wind and reduced daylength resulted in the development of an upper water column with a more-or-less uniform distribution of salinity and temperature and hence density. The mixed water column conditions lasted until the ice breakup in spring 1997.

Annual succession of auto- and heterotrophs

We sampled a complete annual succession of plankton from April 1996 to April 1997 and the beginning of a second succession cycle with the biomass build-up in May through June 1997.

Phytoplankton

Prior to May 4 the phytoplankton biomass was low (Fig. 2c). The chl a in the upper 30 m of the water column was 0.2 ± 0.1 µg l–1. Between May 4 and the subsequent sampling on May 20, a spring bloom developed and chl a increased to 5.9 ± 1.2 µg l–1. *Phaeocystis* sp., chain-forming centric diatoms of the genus *Thalassiosira*, *Detonula confervaceae*, and pennate diatoms dominated this bloom. After the bloom the phytoplankton surface community changed, and nanoflagellates became more abundant than diatoms. In the pycnocline, diatoms (*Chaetoceros* spp.) remained dominant most of the summer, where they formed a subsurface bloom. The annual development of integrated phytoplankton biomass showed increasingly higher levels after the main spring bloom and following summer minimum, reflecting this summer subsurface bloom (Fig. 3a). The bloom persisted until August, when a second phytoplankton peak occurred. During the autumn-winter transition chl a gradually declined to a minimum of 0.11 ± 0.07 µg l–1 in the upper 30 m between November and March. Winter chl a concentrations corresponded to deep water levels during summer.

Following the break-up of the sea ice, the phytoplankton developed again during late April and peaked in mid-May. The integrated phytoplankton spring bloom biomass recorded in 1997 was nearly twice the biomass of 1996. This reflects the fact that the bloom peaked between 2 samplings in May 1996, as
verified by the depletion in the concentration of nitrate. The dominant bloom forming phytoplankton species in 1997 were similar to those observed in 1996 (H. A. Thomsen unpubl. data).

Ice algae

Pennate diatoms accumulating at the bottom of the sea ice prior to the spring phytoplankton bloom dramatically increased the algal biomass in April (Fig. 3a). The mean net ice algal growth rate in 1997 was 0.11 d\(^{-1}\). On April 15 the chl \(a\) concentration peaked at 184 ± 76 \(\mu g\) l\(^{-1}\) in the bottom ~1 cm of the ice \((n = 3)\). The ice algae biomass declined during the last 2 sampling occasions before ice breakup. At the ice-water interface, chl \(a\) displayed a similar trend, delayed a few days in comparison with the ice algae (data not shown). Here the mean growth rate was 0.17 d\(^{-1}\). Maximum chl \(a\) concentration at the ice-water interface was 13 \(\mu g\) l\(^{-1}\), which is comparable to the phytoplankton bloom maximum of 14 \(\mu g\) chl \(a\) l\(^{-1}\) recorded a month later. However, the integrated ice algal biomass was only ~10% of the pelagic phytoplankton biomass.

Protozoa

The bimodal seasonality of the primary producers was reflected in the biomass patterns of protozoa, which ranged over 3 orders of magnitude from winter to summer (Fig. 3b,c). In late winter the integrated community biomass \((0 \text{ to } 200 \text{ m})\) for both ciliates and heterotrophic dinoflagellates was ~10 mg C m\(^{-2}\). The spring biomass increase followed the ice-breakup-induced phytoplankton bloom initiated in April-May with only a short time lag, on the order of days. A typical summer community was established by the end of May. In August the ciliate and heterotrophic dinoflagellate biomasses reached a second peak corresponding to annual maxima of 620 ± 206 and 910 ± 347 mg C m\(^{-2}\), respectively (mean ± SD of 4 samplings). Heterotrophic dinoflagellate biomass subsequently declined at a constant rate; ciliate biomass declined 3 wk later. The annual minimum was reached in January. Winter levels of protozoan biomass remained low until April.

| Table 1. Net community growth and loss (mortality + respiration) rates \((\mu d^{-1})\) at 2 m depth of chlorophyll \(a\), phytoplankton, total ciliates (excluding Myrionecta rubra) and heterotrophic dinoflagellates assuming exponential growth. First number in parenthesis: \(r^2\) of the regression line fitted to the data points representing the exponential model; second number in parenthesis: n, the number of data points included in the model; see 'Methods: Growth and mortality'; –: No data |
|-----------------|-----------------|-----------------|-----------------|-----------------|-----------------|
| Chlorophyll \(a\) | 0.09            | 0.01            | −0.05           | 0.05            | 0.09            | 0.09            |
| (0.80;11)       | (0.73;18)       | (0.91;7)        | (0.90;8)        |                  |
| Phytoplankton   | −               | 0.00            | −0.04           | 0.11            |                 |
| (0.16;10)       | (0.89;7)        | (0.94;6)        |                  |                  |
| Ciliates        | 0.15            | 0.00            | −0.04\(^a\)     | 0.07            |                  |
| (0.88;12)       | (0.11;17)       | (0.84;11)       | (0.79;10)       |                  |
| Heterotrophic dinoflagellates | 0.12            | 0.01            | −0.05           | 0.08            |                  |
| (0.87;12)       | (0.24;8)        | (0.94;15)       | (0.94;10)       |                  |

\(^a\)This value covers 2 periods with different mortality rates. From October 18 to November 18, 1996, \(\mu = −0.1\) \((r^2 = 0.98, n = 4)\); from November 27, 1996 to February 3, 1997, \(\mu = −0.03\) \((r^2 = 0.83, n = 7)\).
Similar seasonal abundance and biomass patterns were observed in all depth strata investigated (Fig. 4a–j). At 2 m and the depth of fluorescence maximum, or 15 m, both ciliate and heterotrophic dinoflagellate biomasses were on the order of 0.1 µg C l⁻¹ during winter and 10 µg C l⁻¹ during summer. Cell concentrations were <<1 × 10³ and 1–10 × 10³ l⁻¹, respectively, during these seasons. Below the euphotic zone, the abundance and biomass levels were approximately an order of magnitude lower; here the seasonal amplitude was dampened but otherwise followed the same pattern as at the surface (Fig. 4e–j).

In spring, net community growth rates at 2 m for the 2 protozoan groups of 0.07 to 0.08 d⁻¹ were only slightly lower than the growth rates of their phytoplankton prey, which was 0.09 to 0.11 d⁻¹ (Table 1). The net mortality calculated for the protozoan communities at 2 m during autumn 1996 corresponded to a loss of 4 to 5% of the biomass per day equal to that of phytoplankton (Table 1).

**Species diversity and succession**

Naked oligotrichous ciliates (*Strombidium*, *Strobilidium*, *Lohmaniella*, *Laboea*, *Tontonia* spp.) and athecate, gymnodinoid dinoflagellates (*Gymnodinium*, *Gyrodinium* spp.) were the most abundant protozoans. Tintinnids and thecate dinoflagellates were less numerous. Usually <10% of the total integrated ciliate biomass was tintinnids. Thecate heterotrophic dinoflagellates, mostly *Protoperidinium* spp. and species of the ‘*Diplopsalis* group’, also contributed <10% of the heterotrophic dinoflagellates during summer, but they accounted for between 10 and 60% of the biomass during the autumn-winter period. In addition to oligotrichous ciliates and tintinnids, other ciliates were present (e.g. *Didinium* sp.), but they comprised an insignificant fraction of total ciliate abundance and biomass. An exception was *Balanion* sp., which at times were abundant, obtaining maximum concentrations of 2.5 × 10³ l⁻¹ (35% of total number of ciliates). In terms of biomass this corresponds to 10% of the total ciliates.

The annual successions of selected protozoans showed that each species had 1 or 2 peaks of occurrence of different amplitude and duration (Fig. 5). Peaks ranged from ~5 mg C m⁻² for the small dinoflagellate *Amphidinium sphenoides* to ~900 mg C m⁻² for the large dinoflagellate *Gyrodinium spirale*. Among the ciliates *Laboea strobila* obtained the highest bio-

Fig. 4. Annual patterns of ciliates (excluding *Myrionecta rubra*) (left panels) and heterotrophic dinoflagellates (right panels) at (a, b) 2 m, (c, d) the fluorescence maximum depth, (e, f) 50 m, (g, h) 100 m and (i, j) 200 m. (●) Biomass; (---) concentration. Data from April 22 to May 5, 1996, are from Grastrup-Hansen (1998). Note logarithmic scales.
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mass, 235 mg C m\(^{-2}\). The autotrophic ciliate *Myrionecta rubra* had a maximum of 110 mg C m\(^{-2}\), which corresponds to ~20% of the total integrated ciliate biomass. On average this species contributed 5%.

Protozoan species richness was lowest during the winter and highest in the summer (Fig. 6a,c), but when corrected for differences in sample size the seasonality in species diversity was less pronounced (Fig. 6b,d). Winter diversity did not change significantly from that in summer (ANOVA, p > 0.05), although heterotrophic dinoflagellates showed a tendency to decline during the late autumn.

The seasonal distribution of cell size of the 2 protozoan communities differed (Table 2). While the mean ciliate volume did not change, winter dinoflagellates were smaller than summer cells. In deep-water samples, protozoans were small throughout the year.

*Gyrodinium spirale* showed the typical size pattern for dinoflagellates (Fig. 7c,d). Small individuals occurred during winter and large cells during the summer, associated with the diatom subsurface bloom (Fig. 7d). *Myrionecta rubra* also demonstrated a pronounced seasonal variation in volume (Fig. 7a,b). Thus, on the species level the trend of this ciliate contrasted the temporal size variation on the community level.

Mixotrophic ciliates were recorded even in winter and during summer they dominated, accounting for 45 to 85% of the total oligotrichous ciliate numbers at 2 m (Fig. 8). Division of oligotrichs into the size groups ‘small’ (<30 µm),

<table>
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<th>Depth (m)</th>
<th>Ciliates Summer</th>
<th>Winter</th>
<th>Dinoflagellates Summer</th>
<th>Winter</th>
</tr>
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<tbody>
<tr>
<td>2</td>
<td>13.9 ± 6.3</td>
<td>12.3 ± 5.7</td>
<td>13.8 ± 8.4</td>
<td>5.2 ± 1.9</td>
</tr>
<tr>
<td>FM</td>
<td>12.1 ± 9.7</td>
<td>10.7 ± 5.3</td>
<td>13.1 ± 12.9</td>
<td>5.8 ± 2.4</td>
</tr>
<tr>
<td>50</td>
<td>7.3 ± 3.4</td>
<td>10.1 ± 5.5</td>
<td>6.5 ± 2.7</td>
<td>4.8 ± 2.9</td>
</tr>
<tr>
<td>100</td>
<td>5.8 ± 3.3</td>
<td>6.9 ± 2.3</td>
<td>6.4 ± 1.6</td>
<td>9.2 ± 5.7</td>
</tr>
<tr>
<td>200</td>
<td>5.4 ± 2.6</td>
<td>5.1 ± 2.0</td>
<td>5.7 ± 1.7</td>
<td>3.5 ± 1.4</td>
</tr>
</tbody>
</table>

**Table 2. Mean summer and winter volumes of ciliates and heterotrophic dinoflagellates (µm\(^3\)×10\(^3\) ±SD). FM = depth of fluorescence maximum, numbers in parenthesis refer to samples (n)**
'medium' (30 to 50 µm) and 'large' (>50 µm) showed that mixotrophy dominated in all size classes (ANOVA, p < 0.05). Thus, the probability that an oligotrich was mixotrophic was larger than the probability that it was heterotrophic. However, all ciliates <20 µm were heterotrophic. The same seasonality as observed at 2 m depth was found in subsurface waters, although the overall contribution of mixotrophs was lower.

**DISCUSSION**

This study emphasizes that pelagic protozoans are essential components of the Arctic pelagic food web during the summer period. It also provides new information on winter plankton dynamics. A diverse but dilute protozoan community was present throughout the winter. This observation is corroborated by extensive sampling and it enables us to hypothesize on pelagic heterotrophic protistan overwintering strategies in a system where phytoplankton production is low or zero during several months each year.

**Annual dynamics and composition**

Disko Bay experiences great fluctuations in the degree of ice cover from year to year. In the winter of 1996/97, sea ice was formed late due to mild weather conditions. Seasonal changes in the water column structure are mainly determined by differences in the extent of melt water. Stratification is established when Disko Bay receives a large freshwater input of melting snow and ice in early summer. These melting processes result in surface water that flows out of the Bay (e.g. Andersen 1981). Advection thus influences our sampling station. Nonetheless, during horizontal surveys in Disko Bay in June-July 1996 and June 1997 we observed that the general pattern in plankton succession was similar on a larger scale (P. Munk et al. unpubl. data). For the purpose of this discussion, we assume that our sampling station represents the annual plankton succession in the Bay.

Fig. 6. Species richness ($S$) normalized to maximum, and Margalef’s species diversity index ($I_m$) of (a,b) ciliates and (c,d) heterotrophic dinoflagellates at 2 m depth. [$I_m = (S - 1)/\ln N$]

Fig. 7. Seasonal variations in the mean cell volume of (a,b) *Myrionecta rubra* and (c,d) *Gyrodinium spirale*, (a,c) at 2 m depth and (b,d) at the depth of the fluorescence maximum.
The protozoan community in Disko Bay showed a strong seasonal signal in biomass. The winter values are among the lowest ever reported for marine waters (Garrison et al. 1993, Scharek et al. 1994, Boyd et al. 1995, Sherr et al. 1997). The summer protozoan concentrations, in contrast, were as high as those reported for neritic, temperate marine waters (Hansen 1991, Pierce & Turner 1992). Key species were the ciliate Laboea strobila and the heterotrophic dinoflagellate Gyrodinium spirale. This agrees with data from the same sampling station collected during summer 1994 (Levinsen et al. 1999).

Ciliates and heterotrophic dinoflagellates revealed a similar temporal pattern with one peak in late May-early June and another in August. Such a pattern is different from the unimodal pattern found for ciliates in Antarctic coastal waters (Leakey et al. 1994), although the difference may reflect a lower temporal and spatial sampling frequency in the Antarctic study. Bimodality in protozoan biomass is found also in temperate waters; here peaks co-occur with the spring and autumn phytoplankton blooms (Smetacek 1981, Nielsen & Kierboe 1994). Thus, on an annual scale, these patterns from different latitudes demonstrate a tight trophic coupling between phytoplankton and protozoan grazers. The pattern of protozoans from Disko Bay, however, contrasts with the classical idea that the primary productivity cycle in the Arctic consists of a single peak connected to the ice-break (Smith & Sakshaug 1990).

Heterotrophic dinoflagellates, primarily athecate species, frequently dominated spring and summer as reported for temperate (Smetacek 1981), tropical (Verity et al. 1996), Antarctic (Bjørnsen & Kuparinen 1991) and Artic marine waters (Nielsen & Hansen 1995). Compared to mesozooplankton heterotrophic dinoflagellates can respond quickly to episodic diatom blooms. Obviously, carbon flow studies that do not include this group of organisms are not complete.

The majority of the ciliates were mixotrophic. This confirms several other studies where plastidic oligotrichs have been found to be ubiquitous and important components of surface plankton communities (Stoecker et al. 1989, Putt 1990, Bernard & Rassoulzadegan 1994). Unexpectedly, mixotrophs were also observed in high proportions during winter (18%). Even Laboea strobila, which is considered an obligate mixotrophic ciliate (Stoecker et al. 1988), was recorded during the dark winter months, as was the autotrophic ciliate Myrionecta rubra (Fig. 9). Apparently, the dependence on photosynthesis for growth does not prevent individuals of these species from surviving periods without light.

**Overwintering strategy**

Dormancy and overwintering in copepods have been studied extensively (Conover & Siferd 1993, Dahms 1995, Hirche 1996) and dark survival strategies in...
phytoplankton have also received attention (Fryxell 1983, Jochem 1999). Much less is known about the strategies used by protozoans to cope with adverse food conditions. Laboratory studies report rapid mortality of starving ciliates, where all cells die during a few days (Fenchel & Jonsson 1988, Nielsen et al. 1990, Stoecker & Silver 1990, Wickham et al. 1993, Montagnes 1996, Jacobsen & Hansen 1997). This is in stark contrast to the observations of protozoa in Disko Bay throughout the year. Evidently, a winter protozoan community survived several months despite the fact that food concentrations were below the threshold where the growth rate equals zero (Fig. 10). In the following, we focus on this paradox and discuss 4 overwintering strategies in order to answer the question ‘How do protozoans survive the arctic winter?’

(1) Allochthonous sources

Advection of plankton from southern areas can lead to erroneous conclusions on winter surviving protozoa. Disko Bay occasionally receives Irminger Water of Atlantic origin, which flows northwards as a coastal stream. However, the stream occurs at depths much below productive layers, and further, assuming typical winter flow velocities of 20 to 25 cm s⁻¹ (E. Buch pers. comm.), it takes almost 2 mo for a parcel of water to move the ~1000 km up the west Greenland coast and enter the Bay. Adverted protozoa should therefore, like the protozoa in Disko Bay, be able to starve for long periods.

Sea ice is another potential source of allochthonous material. Garrison et al. (1993) suggested that input of organisms and detritus from the sea ice served as an alternative carbon source for planktonic heterotrophs in the ice-edge zone of the Weddell and Scotia Seas during the early austral winter. It is likely that, in the late spring, input from sea ice also occurred in Disko Bay, but during winter, when production in the ice was negligible, such input could not be important for sustaining pelagic protozoan communities. We conclude that allochthonous input is of temporally restricted or local importance and not a mechanism able to explain the general paradox.

(2) Cyst formation

Fenchel (1987) hypothesized that responses of protozoa reflect the time scale of fluctuations in food resources. Cells will stay active if periods of starvation are short and will encyst under extended periods of starvation. As the polar winter certainly has an extended period with low food supply, encystment would be expected. However, cysts were rarely recorded in the present study and then only of tintinnid and thecate dinoflagellate origin. The dominant naked oligotrichous ciliates and gymnodinoid dinoflagellates were never found to produce cysts. The apparent lack of resting stages might reflect (1) difficulties in detection and identification of cysts, or (2) that many pelagic protozoans do not usually produce cysts. Only 3 non-tintinnid, marine oligotrichous ciliates and ~5% of all dinoflagellates are known to produce resting cysts (Fauré-Fremiet 1948, Reid 1987, Kim & Taniguchi 1995, Sonneman & Hill 1997).

In Disko Bay steep coastal slopes limit resuspension of potential cysts from shallow water. Rather than acting as a near-surface ‘inoculum’ coastal cysts would be advected to deep water (up to >800 m), and likely become lost from the pelagial, if sedimenting there. At least, a fast responding protozoan community on the spring phytoplankton increase, such as the observed, would not be expected. The adaptive significance of cyst formation, therefore, seems to be small for the pelagial protozoan community of Disko Bay. We hypothesize that the most favorable overwintering strategy for planktonic protozoa involves them remaining motile in surface waters until conditions become favorable for growth.
(3) Food patchiness

Protozoans in Disko Bay may regularly, by chance, have encountered and exploited prey patches. A generally stable field ciliate abundance in the range of 1 to \(10 \times 10^3\) cells l\(^{-1}\) may be used as an indication of the existence of prey patches because ‘average’ prey biomass levels often are near or below the ciliate threshold concentrations (Montagnes 1996). For >3 mo during winter, when ciliate and heterotrophic dinoflagellate concentrations in the Bay were \(<1 \times 10^3\) cells l\(^{-1}\), potential food levels were much below the average threshold value for growth (Fig. 10). This suggests that if the protozoan populations survived this period based on growing individuals they relied on concentrated prey (e.g. on marine aggregates) with distances between patches that correspond to no more than the distance displaced within a protozoan lifetime. The prey enrichment should exceed the gap between the measured bulk phytoplankton concentration and the threshold concentration depicted in Fig. 10 (triangles) because many protozoans only feed on prey within a certain size spectrum (Jonsson 1986, Kivi & Setälä 1995). Thus, not all phytoplankton cells in such a food estimate may be suitable prey. Food patchiness was not examined in this study. However, we expect that rapid aggregate degradation by presumably carbon-limited bacteria populations resulted in a depleted prey habitat within a few weeks after aggregate formation during the spring-summer period. New aggregate formations during winter would be severely limited due to low plankton concentrations.

(4) Reduced winter metabolism

The heterogeneous environment suggests that protozoa seldom experience balanced growth. Naturally, therefore, they must have developed metabolic adaptations to variations in food resources in addition to encystment. From laboratory measurements, some protozoans are known to reduce their respiration substantially at resource levels below that sustaining minimum growth (Fenchel & Finlay 1983). If ciliates and heterotrophic dinoflagellates typical for the pelagial can similarly respond to the onset of the starvation period by reducing their metabolism, it will greatly prolong their survival.

How long will it take starving, non-feeding ciliates to respire their cell carbon at –1.5°C?

This can be estimated using \(O_2\) consumption rates for starved ciliate and flagellate species reviewed by Fenchel & Finlay (1983) and the following assumptions: cells survive 80% shrinkage (Caron et al. 1990); carbon content = 0.13 pg C \(\mu m^3\) and \(Q_{O_2} = 2.8\) (Hansen et al. 1997); respiratory quotient (RQ) = 0.8; 1 l \(O_2 = 1.428\) g \(O_2\) (the density of \(O_2\) gas).

Temperature-corrected cell-specific \(O_2\) consumption rates, expressed in nl \(O_2\) cell\(^{-1}\) h\(^{-1}\), from 21 measurements reported in the review were converted to daily respiration rates in carbon units (pg C d\(^{-1}\)). For each species, the cell carbon weight, \(W\) (pg C), and respiration rate, \(r\), were inserted in the allometric equation with an exponent of 0.75 (Fenchel & Finlay 1983)

\[
r = aW^{0.75}
\]

The respiration rate was transformed to a specific respiration rate by division with \(W\)

\[
\frac{r}{W} = aW^{-0.25}
\]

and the constant \(a\) isolated. As there was no difference between ciliates and flagellates (Student’s \(t\)-test, \(p > 0.05\)) a mean constant (± SD) was calculated from all 21 calculations, which gives \(a = 0.091 ± 0.056\) pg C pg\(^{-1}\) C d\(^{-1}\). Inserting \(a\) and the carbon content of a mean sized ciliate of 15 000 \(\mu m^3\) into Eq. (2) yields an initial weight-specific respiration rate during starvation of 1.4% d\(^{-1}\) (0.5 to 2.2%).

The weight loss over time due to respiration can be expressed by the differential equation

\[
\frac{dW}{dt} = -aW^{0.75}
\]

assuming that the allometric relationship documented at the interspecific level (Eq. 1) is applicable also for a shrinking individual. Inherent in this assumption is that the weight-specific respiration rate of the starving individual increases with time, leading to a conservative estimate of survival time compared to applying a constant specific rate. Eq. (3) can be solved to

\[
W_t = (-0.25at + W_0^{0.25})^4
\]

where \(W_0\) is the initial weight and \(W_t\) is the weight at time \(t\). Inserting \(W_t = 0.2 \times W_0\) to Eq. (4) and isolating \(t\) gives the survival time \(t_{20\%}\)

\[
t_{20\%} = \frac{4(1 - 0.2^{0.25})W_0^{0.25}}{a}
\]

Using Eq. (5), a calculation for the mean sized, 15 000 \(\mu m^3\), ciliate resulted in a survival period of 3 mo (60 to 252 d). The survival periods of the largest (350 000 \(\mu m^3\)) and the smallest (1500 \(\mu m^3\)) protozoans in Disko Bay with initial specific respiration rates of 0.6 and 2.4% d\(^{-1}\), respectively, correspond to 213 and 54 d. Thus, as evident from Eq. (5), survival time relates to the size of the initial cell making it more profitable to
be large when starvation begins (‘survival of the fattest’). The observed size range of Laboea strobila indicates that the selected minimum cell size of 20% of \( W_0 \) is reasonable. The smallest \( L. \) strobila cells were 16% of the largest (Fig. 9). Indeed, reduced metabolic activity may explain survival of at least the largest protozoans even if they were not feeding at all.

The observed loss rate of the protozoans was 4 to 5% of the community biomass per day (Table 1). It is unknown how much of this loss was due to respiration and how much was due to mortality, and to what degree the cells were still feeding and growing. However, this loss rate is much closer to the loss rates calculated above (assuming reduced metabolism) than to the mortality rates found for starved protozoans in the laboratory, where rates of >50% d\(^{-1}\) are usually reported (e.g. Crawford & Stoecker 1996).

Based on the discussion of overwintering strategies we suggest that the protozoans survived the winter primarily due to their ability to reduce the maintenance metabolism. We speculate that pelagic protozoans have developed metabolic adaptations to variations in food resources in addition to encystment and hypothesize that there exists a continuum of metabolic stages between exponentially growing cells and cysts. Because of the suggested starvation mode, which was influenced by low temperature, they only needed very rare prey encounters to survive. A simple model of prey encounter in a turbulent environment showed that the protozoans encountered ~3 prey particles d\(^{-1}\) (calculated according to Kjørboe 1997). Prey encounters of this order of magnitude corresponds to a specific ingestion less than a few percent of the maximum ingestion measured in the laboratory at 12 to 22°C (Hansen et al. 1997). Inclusion of ciliates and heterotrophic dinoflagellates in the prey category for the possibility of a heterotrophic population ‘feeding on itself’ does not change this figure.

On average, the protozoan biomass during the dark months of November to February was ~70 mg C m\(^{-2}\) (ciliates 40, dinoflagellates 30). Assuming a metabolic rate of 1.4% d\(^{-1}\) protozoans would require a primary production of ~1 mg C m\(^{-2}\) d\(^{-1}\) (70/100 x 1.4). The average measured production in November to February was about 0.6 mg C m\(^{-2}\) d\(^{-1}\) (T. G. Nielsen et al. unpubl.). Therefore, despite the fact that primary production is negligible, it (on average) might be large enough to support the populations of protozoa in starvation mode.

We further expect the threshold concentration for growth to show temperature dependency, which would lead to a lowering of reported threshold values. This seems reasonable as respiration is found to scale with temperature for various zooplankton organisms (e.g. Hansen et al. 1997).

Finally, the majority of the ciliates in Disko Bay and an unknown proportion of the dinoflagellates were mixotrophic. They potentially have the additional ability to digest retained plastids during starvation where autophagy occurs, thus having an advantage over strict heterotrophs.

### Ice-water interactions and the winter-spring transition

Prior to the spring phytoplankton blooms ice algae developed in the bottom sea ice. Microscopic evidence suggests that these algae did not contribute significantly to the pelagic community as proposed by Smith & Nelson (1985) and Michel et al. (1993). Rather, they sediments quickly out of the euphotic zone and formed a transient food resource for pelagic grazers, when released from the melting ice. This is in accordance with studies, which suggest that ice algae tend to form aggregates with high sinking rates (Carey 1987, Riebesell et al. 1991).

However, before sedimentation and the ice-breakup induced phytoplankton bloom the ice communities may have provided starved pelagic protozoans with a concentrated food source. The chl \( a \) concentration measured at the ice-water interface was several times higher compared to at 2 and 15 m in the last weeks prior to ice breakup (not shown). Grazing on this algal layer would pre-acclimate the winter populations of protozoa, so that they were not severely starved or even were able to resume growth rapidly after initiation of the phytoplankton bloom. Protozoans need to rebuild their cell machinery before they can resume growth after a starvation period (Fenchel 1982). The short time lag on the spring bloom supports this hypothesis.

The winter populations of ciliates and heterotrophic dinoflagellates acted thus as a ‘reservoir’ of starvation-resistant individuals, which could multiply quickly when favored by the right conditions. Net community growth of protozoans was comparable to that of their phytoplankton prey, and they responded exponentially to the spring phytoplankton bloom almost instantaneously, increasing 100-fold from winter to summer. The exponential net increase occurred despite sub-zero temperatures and the potential predation impact from a large biomass of Calanus spp. that ascended from their deep-water hibernation in April (Madsen et al. unpubl.).

For shallow-water diatoms, 1% survival of the cells which sink to the bottom in spring is sufficient to seed the water column the following year (Jewson et al. 1981). A corresponding scenario for Disko Bay protozoans, but assuming a pelagic inoculum, gives a seeding concentration equal to 5% of the summer concen-
Thus, a winter ‘inoculum’ of the recorded 262 ± 181 cells l⁻¹ is sufficient to initiate the succession of protozoa the following year and no other seeding mechanisms are necessary to explain the observed biomass buildup.

Succession and diversity

Ciliates and heterotrophic dinoflagellates showed a highly diverse succession pattern (Fig. 11). Coexisting species that presumably occupied slightly different niches peaked successively during the summer. The small ciliate *Lohmaniella oviformis* was one of the earliest species. It increased exponentially from mid-April and reached a peak at around June 1 in both years. In September, it had a second but less pronounced peak. Small bacterivorous ciliates, on the other hand, peaked rather late (October 1) and in between the peaks of these 2 species were peaks of several others. Such seasonal patterns did presumably also exist for the remaining morphospecies that were grouped together in size classes.

At least for ciliates the ‘morphospecies’ concept appears to be as robust as any other species concept for microbial groups (Finlay 1998). The most striking result of the analysis of diversity is a relatively diverse although dilute winter community. Generally, the morphospecies richness was smaller during winter, but it seems only to be a matter of processing more water to get a higher morphospecies number. In winter, the number of individuals counted was low and 325 ml of water was clearly not enough to find the rare morphospecies. Even when counting 1500 individuals from a single water sample some morphospecies were represented by only 1 to 3 specimens (Fig. 12). Considering that many fewer cells were counted during winter, it is likely that an unknown number of morphospecies was not recorded.

If, for convenience, ciliates and heterotrophic dinoflagellates are divided into 6 functionally and taxonomically different groups (*Myrionecta rubra*, tintinnids, *Strombidium* sp., *Strobilidium spiralis*, *Laboea strobila*, bacterivorous ciliates, (1) *Gyrodinium* sp., (2) *Protoperidinium bipes*, (3) thecate dinoflagellates, (4) *Gyrodinium spirale*).
Conclusion

Compared to in other marine areas, the importance of microbial populations in polar food web dynamics remains poorly documented and does yet not have general acceptance. Ecosystem structure and function in cold waters such as Disko Bay is, however, not different from that at lower latitudes, where protozoan grazers are accepted to play a pivotal role. The present study, through a detailed analysis of the annual pattern of ciliates and heterotrophic dinoflagellates, emphasizes the existence of a microbial community with a high complexity. This complex community of protozoans was maintained even during winter, when the 'obligate' oligotrichous ciliate Laboea strobila and other mixotrophic ciliates were observed. The study stresses that particularly oligotrichous ciliates and heterotrophic dinoflagellates may contribute substantially to the pelagic protozoan biomass in the Arctic. Finally, the winter sampling showed that a small community of ciliates and heterotrophic dinoflagellates could maintain itself in a potentially active state during the dark arctic winter despite living in what seems to be a well-below-food-threshold ecosystem. We suggest that this was accomplished primarily through the ability of the protozoans to reduce their metabolism.

Acknowledgements. We warmly thank the crew on RV 'Porsild' for help during the many cruises. Particularly, we are grateful to Jørgen Broberg and Johannes Filimonsen for their support when the weather and sea were (too) rough. Thanks to the late Leif Skytte for logistical support and to Birgit Soborg for technical assistance. The board of Arctic Station, University of Copenhagen, is acknowledged for letting us use their laboratory premises. Special thanks to Joan Harrison for valuable discussion and insights in mathematics. Finally, we thank Kurt Buck, David W. Crawford, David Montagnes, Diane Stoecker, Helge A. Thomsen, Peter Tiselius, Jefferson T. Turner, Thomas Kierboe, Peter Verity, Per Juel Hansen and 2 anonymous reviewers for valuable comments on earlier drafts of the manuscript. The investigation was supported by The Danish National Research Council project no. 9501038.

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Editorial responsibility: Otto Kinne (Editor), Oldendorf/Luhe, Germany

Submitted: September 20, 1999; Accepted: May 4, 2000
Proofs received from author(s): September 25, 2000