

Environmental impacts of coastal fish farming: carbon and nitrogen budgets for trout farming in Kaldbaksfjørður (Faroe Islands)

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ABSTRACT: Flow of organic carbon (OC) and nitrogen through a sea cage trout farm was calculated on the basis of detailed studies of the farming operation, water circulation, OC and nutrient transport and recycling processes in sediment. A third of the OC and nitrogen provided by fish food was incorporated into fish biomass, which is more than has been found in previous studies. Most OC input was respired by the fish (52 to 70%), and ~63% of the associated nitrogen was lost as dissolved inorganic nitrogen (DIN), potentially stimulating pelagic primary production. Approx. 6% of carbon and 5% of nitrogen derived from fish food settled on the seabed, where it was either mineralized or accumulated in the sediment. Based on transect measurements of diagenetic activity, the farm footprint was found to cover an area ~10 times the farm area. OC mineralization in the sediment increased linearly with increasing food input; the divergence between carbon efflux and oxygen uptake in sediment likewise increased with increasing food input, reflecting an increasing level of sediment reduction. Directly below the farm, the dissolved organic carbon (DOC) efflux was high (on average 53% of dissolved inorganic carbon efflux), indicating that DOC efflux is an important pathway for benthic carbon release below aquaculture farms. Overall, microbial processes removed 56 and 38% of OC and nitrogen, respectively, that settled to the seabed. During a 39 d break in farming activity, due to the combined effect of mineralization and resuspension of surface sediment, sediment conditions improved considerably.

KEY WORDS: Fish farming · Sediment · Organic enrichment · Nutrient enrichment · Organic matter mineralization · Carbon budget · Sedimentation · Benthic recovery

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INTRODUCTION

Coastal cultivation of fish in floating net cages is a rapidly expanding industry worldwide, now providing a significant portion of the world's market in fish (FAO 2009). Fish are typically cultivated in cages or pens in sheltered coastal areas; they impact the marine environment primarily through release and accumulation of waste products (Pillay 2004). Since these areas often have little water exchange, carrying capacity is limited, and the environment may be heavily affected. The oxygen consumption of the farmed fish and asso-

ciated microbes in the cages may put excessive demand on the dissolved oxygen (DO) supply, reducing it severely (Johansson et al. 2007). Likewise, ammonia excretion from the farmed fish can lead to elevated ammonium concentrations in the cages (Brooks & Mahnken 2003, Islam 2005), which may have negative effects on fish health.

Since nitrogen is the limiting nutrient in many marine environments, the input of ammonium might lead to increased primary production and changes in the plankton community. Such effects are usually not observed at the fish farm itself (Brooks & Mahnken

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2003, Islam 2005, Pitta et al. 2006), but increased nutrient loading from fish farming may cause such effects on a regional scale (Karakassis et al. 2005). Pelagic effects are highly dependent on hydrographic and nutrient status in the farming area.

Models have shown little effect on primary production in a Norwegian fjord (Skogen et al. 2009). However in oligotrophic waters in southern Tasmania, theoretical models indicate significant impact on primary production, especially during summer and autumn when surface waters are otherwise nutrient-depleted (Wild-Allen et al. 2010).

Particulate waste products in the form of fish food and faeces quickly sink to the seafloor (Cromey et al. 2002). The sedimentation of organic carbon (OC) below fish farms has been found to be from 4 to 27 times higher than at unaffected sites, declining rapidly with distance from the farm (Hall et al. 1990, Chamberlain & Stucchi 2007, Holmer et al. 2007, Kutti et al. 2007). The high deposition of waste particles in the sediment near aquaculture operations stimulates metabolic activity in the sediment and hence the consumption of electron acceptors, resulting in changed pathways for carbon mineralization and nutrient regeneration (Christensen et al. 2003, Valdemarsen et al. 2009).

Dissolved inorganic carbon (DIC) efflux and total oxygen uptake (TOU) in sediments below fish farms often is an order of magnitude higher than at non-fish-farm sites (Hall et al. 1990, Findlay & Watling 1997, Holmer et al. 2002, 2003, Nickell et al. 2003, Giles 2008). Elevated ammonium effluxes and nitrogen uptake are likewise common in sediments impacted by fish farms (Hall et al. 1992, Holmer & Kristensen 1992, Christensen et al. 2000, Holmer et al. 2002, Nickell et al. 2003).

Reduced sulphur species accumulate in the sediment, where they can be released to the overlying water, causing local elimination of fauna and hampering the well-being of the farmed fish (Black et al. 1996, Hargrave et al. 2008). In some cases, enhanced methanogenesis may even lead to release of methane gas (Hall et al. 1990, Holmer et al. 2002).

In order to halt serious deterioration of sediment conditions below fish farms, it is necessary to move the cages from time to time, allowing the sediment to recover. However, in many regions of the world, suitable fish farm areas are limited by space constraints and legislation. Many farming operations, laboring under such constraints, must intensify production for economic profitability, putting an even greater strain on the local environment (FAO 2009).

From the perspective of environmentally sustainable estuarine environments, it is essential to evaluate the wider impact of farming wastes on the plankton com-

munity and how this interrelates with carrying capacity. The subject of environmental impact of aquaculture has received a lot of attention worldwide, and an emerging view is that aquaculture operations should be managed using an ecosystem-based approach (Soto et al. 2008). However, the nature of farm impacts are very dependent on local characteristics such as hydrodynamic conditions, temperature and water depth at the site (Kempf et al. 2002), as well as on farm size and husbandry methods (Mente et al. 2006).

In trying to manage a fish farm in an environmental context, a thorough knowledge of the associated waste load is essential—not only the quantity, but also the quality: Is it particulate or dissolved, organic or inorganic (Strain & Hargrave 2005)? Quantification may be approached by carbon and nitrogen mass balances (e.g. Hall et al. 1990, Holmer et al. 2002, Strain & Hargrave 2005). In most studies, however, mass balance calculations are based on information gathered from several studies, and under different environmental and physiological conditions, which in some cases may affect the estimated budgets.

Aquaculture has been rapidly growing in the Faroe Islands; in 2009 the country was ranked the world's fifth-largest producer of Atlantic salmon *Salmo salar*, exporting 47 800 t of farmed salmonid fish that year. Fish farming is licensed in the vast majority of fjords and sounds in the Faroe Islands, and is the major source of anthropogenic OC and nutrients in these coastal areas (Mortensen 1990). For sustainable management of the impacted ecosystems, it is imperative to understand how they function and how they respond to increased loading of nutrients and organic matter.

In the present study we investigated the dispersal of OC and nutrient wastes from a fish farm and their impact not only on the water column at the farm, but also on the seabed in the farm vicinity. We estimated how much of the OC and nitrogen added to the system in feed was incorporated into the farmed fish, how much was lost to the environment in particulate and dissolved form, and the effect of this loss on the benthic diagenetic pathway. Results of these studies were combined into an overall mass balance for carbon and nitrogen flow in the farm area. We furthermore investigated the relative importance of mineralization and resuspension on the short-term local sediment recovery rates following intense farming activity. The present study contributes to the knowledge base of environmental impact of farming activity on high-latitude fjord systems.

MATERIALS AND METHODS

Study site. The study was conducted in Kaldbaksfjørður, Faroe Islands, a fjord 6.6 km long and from 500

to 1700 m wide, with a surface area of 5.41 km². At the entrance is a sill from 30 to 40 m deep; maximum fjord depth is 60 m (Fig. 1). Most of the year the water column is 2-layered, with a fresher and warmer surface layer (salinity from 32.1 to 34.9 and temperature from 5.8 to 11.4°C) receiving freshwater run-off from the 42 km² catchment area, and a deeper water mass in contact with seawater from the sound outside the fjord (salinity from 34.9 to 35.3 and temperature from 6.2 to 10.8°C). The pycnocline is generally situated between 8 and 20 m depth. During summer, however, an additional thermocline periodically appears at ~40 m depth.

The hydrography is highly influenced by wind. Stormy periods with daily mean wind speeds of 15 to 20 m s⁻¹ frequently induce water-column mixing. This may happen in all seasons, although it is more common during winter (Hansen 1990). Most of the catchment area is uncultivated and uninhabited, hence the anthropogenic input of labile OC and nutrients is minor. However, fish farming has been operating periodically in the fjord since 1985, and this is the major source of anthropogenic OC and nitrogen to the system (Mortensen 1990). In 2006, 2 areas — each approx. 0.25 km² in extent — were licensed for fish farming. One was near the fjord mouth, and one relatively far from it. We studied the latter site (Fig. 1).

The study was conducted at farming site FS (Fig. 1), which had been without farming activity for more than 6 mo prior to late April 2006, when juvenile rainbow trout *Oncorhynchus mykiss* were introduced into cages there (Fig. 1). Farming continued uninterrupted until 2 October 2006, at which time the platform from which the cages were suspended was temporarily moved to an outer location in the fjord. Moorings, however, were not removed, and on 9 November 2006 the same plat-

form (with the same fish) was reinstalled to the same moorings, and farming continued until 3 April 2007.

Our sampling period was April 2006 to May 2007.

The location contained 10 net cages: 4 cages of 20 × 20 m each, mounted below a floating platform covering an area of 2025 m², and 6 circular cages (diameter 25 m) located 125 m away. We studied only the square cages, and farming activity at the circular cages did not affect conditions at the stations sampled from the platform. Sediment time-series samples were collected at regular intervals during the 13 mo study period on the west side of the platform 1.5 m away from the platform-mounted net cages (water depth 48 m; Fig. 1c). Additionally, on 22 June 2006, sediment samples were taken along a transect at 0, 10, 30 and 60 m distance from the farm (t₁; Fig. 1c) and on 7 September 2006 at 0, 20, 50 and 90 m distance (t₂; Fig. 1c). A reference station (R) at 52 m depth and 2.2 km ESE of the fish farming site (Fig. 1) was sampled in parallel on all sampling dates. Sediment characteristics are given in Table 1.

During the sampling period the total production of farmed fish in the fjord was 2440 t (wet wt), and the total feed supply was 2932 t (dry wt).

Measurements in water column. Vertical salinity and temperature profiles were obtained with a pre-calibrated Seabird SBE 37-SM MicroCAT. Current-depth profiles were measured at the fish farm using a bottom-mounted Acoustic Doppler Current Profiler (AADI 600 kHz RDCP600) deployed at 47 m depth during the period 23 March to 25 April 2006. Recordings were made at 20 min intervals; they represent the mean current over 2 m thick bins with the center of the deepest bin at 44 m depth.

Seawater for determination of NO₃⁻, NH₄⁺, DO, chl *a*, primary production, DIC and particulate organic carbon (POC) was sampled at 1, 5, 10, 20 and 45 m depth with a 5 l Niskin sampler. For DO and DIC concentration measurements, the water was transferred to gas-tight glass vials. DO concentration was determined by Winkler titration (Grasshoff et al. 1999), while DIC samples were preserved with 20 µl of HgCl₂ (5% w/v) until later analysis on an infrared gas analyzer (ADC-225-MK3). NO₃⁻ and NH₄⁺ samples were preserved with 3 droplets of chloroform per 20 ml sample. NO₃⁻ was measured on an auto analyzer according to Grasshoff et al. (1999) and ammonium was measured manually by the salicylate-hyperchlorite method, with a 35 saline standard curve (Bower & Holm-Hansen 1980). The detection limit was 0.2 and 0.5 µM for ammonium and nitrate, respectively, and the accu-

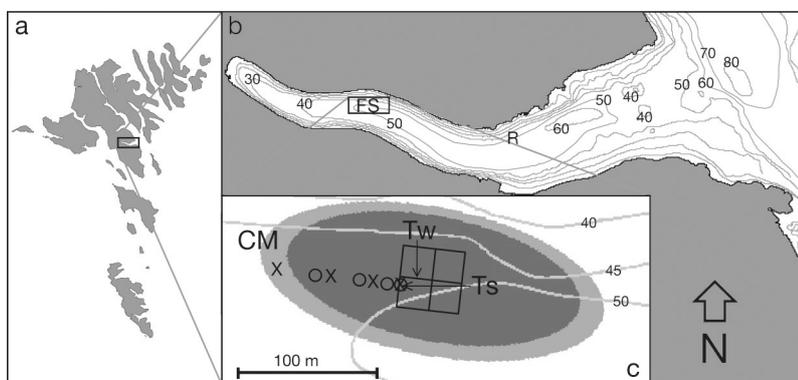


Fig. 1. (a) The Faroe Islands, (b) Kaldbakksfjørður — showing locations of farm site (FS) and reference station (R) — and (c) farm site. Ts indicates location of sediment time-series measurements; Tw indicates location of time-series water and sedimentation samples; CM indicates location of current measurements. Stations sampled on 22 June (O) and on 7 September (X). Light grey in (c) indicates maximum extent of estimated footprint area and dark grey minimum extent of that area

Table 1. Sediment characteristics (\pm SE) at farming site prior to and during farming, 22 June to 29 September, and 5 December 2006. Reference station not statistically different from farming site prior to farming (2-tailed t-test, $p < 0.05$). Numbers in parentheses are total number of samples. **Bold** = significant difference between farming site conditions prior to and during farming; nd = not determined

	Farming site prior to farming	Farming site during farming	Reference station
Grain size (% of total) ^a			
<63 μm	51 \pm 1.6 (2)	52 \pm 2.4 (4)	45 \pm 1.7 (2)
63–250 μm	46 \pm 1.0 (2)	36 \pm 1.9 (4)	48 \pm 3.9 (2)
>250 μm	2 \pm 0.3 (2)	10 \pm 1.0 (4)	6 \pm 5.0 (2)
Density (g cm^{-3}) ^a	1.29 \pm 0.003 (3)	1.07 \pm 0.027 (20)	1.29 \pm 0.02 (27)
Porosity ^a	0.72 \pm 0.004 (3)	0.75 \pm 0.016 (20)	0.71 \pm 0.01 (27)
Organic carbon (mmol g^{-1} dry wt) ^a	1.88 \pm 0.02 (5)	5.55 \pm 0.37 (24)	1.78 \pm 0.07 (31)
Nitrogen (mmol g^{-1} dry wt) ^a	0.2 \pm 0.01 (5)	0.41 \pm 0.02 (24)	0.2 \pm 0.01 (31)
Macrofauna abundance (ind. m^{-2})	nd	2192 \pm 875 (10)	6411 \pm 641 (20)
Most abundant species (% of total) ^b			
<i>Abra nitida</i>	nd	0 (5)	48.0 \pm 5.7 (20)
<i>Thyasira flexuosa</i>	nd	0 (5)	30.2 \pm 3.4 (20)
<i>Scalibregma inflatum</i>	nd	9.2 \pm 12.3 (5)	14.7 \pm 5.3 (20)
<i>Capitella capitata</i>	nd	59.4 \pm 6.0 (5)	0 (20)
^a Mean values from 0 to 4 cm depth			
^b Samples without fauna excluded			

racy was better than 0.2 μM . Chl *a* was measured spectrophotometrically according to Parsons et al. (1984), after filtration of 2 l seawater, and was calculated according to Jeffrey & Humphrey (1975), with a detection limit and precision of 0.2 and $>0.05 \mu\text{g l}^{-1}$ respectively. For POC determination, 1 l of seawater was filtered through pre-combusted (475°C) Whatman GF/F filters, and analyzed on a CE 440 Elemental analyzer (precision $>99\%$) after being fumed with concentrated HCl.

Primary production was measured in samples recovered from 5 and 20 m water depth using the $\text{H}^{14}\text{CO}_3^-$ incubation approach (Steemann-Nielsen 1952) in an incubator exposed to a light gradient (see paragraph below). The 2 sampling depths generally represent phytoplankton communities with potentially different light adaptation from well mixed water bodies above and below the pycnocline. Samples were transferred to 40 ml plastic bottles and stored dark and cold ($\sim 5^\circ\text{C}$) until arrival at the laboratory (2 to 3 h later). After addition of 1.2 μCi of $\text{H}^{14}\text{CO}_3^-$, the bottles were placed on a rotating wheel and incubated at *in situ* temperature $\pm 1^\circ\text{C}$ for 2 h. Philips TLD 15W/33 lamps were used as the light source, providing each bottle with 1 of 11 light intensities from 210 to 16 $\mu\text{E m}^{-2} \text{s}^{-1}$. After incubation, the content of each bottle was filtered on Whatman GF/F filters, and the radioactivity was measured on a Packard Tri-Carb 1500 liquid scintillation counter after addition of scintillation liquid. For each of the 2 depths, a production versus irradiance (PI) curve was constructed based on 12 light bottles and corrected for dark-fixation (2 dark bottles).

The light intensity range during incubation was within the *in situ* light intensity range, which varies

with season, and no photoinhibition was observed. The chl *a*-normalized primary production at depth *z* was calculated according to Sakshaug et al. (1992):

$$P_z^B = P_{\text{max}}^B \times \left(1 - e^{-\left(\frac{\alpha^B \times I_z}{P_{\text{max}}^B}\right)} \right) \quad (1)$$

where P_{max}^B is the chl *a*-normalized light-saturated primary production, α^B is the photosynthetic efficiency and I_z is the irradiance at depth *z*. Daily primary production was calculated from the PI curve, taking into account the *in situ* irradiance and the chl *a* content at 1, 5, 10, 20, and 50 m depth, with linear extrapolation between the measuring points.

Sedimentation. Duplicate sedimentation traps (KC-Denmark; Lundsgaard et al. 1999) were moored to the platform at 20, 40 and 45 m depths, and also at the reference station at 20, 40 and 50 m depths. On 7 September 2006, additional traps were moored 20, 50 and 90 m away from the platform, at 20 and 40 m depth. The flux in the traps at 20 m depth was assumed representative of vertical settling of material from the euphotic zone, while the general elevation in flux in the deeper traps was ascribed to resuspension.

The traps were replaced approximately every second week. One trap from each station + depth combination was kept unpreserved while the other was preserved with formaldehyde added to dense seawater (salinity ~ 60). After recovery of traps, subsamples of the content were filtered on pre-combusted (475°C) and pre-weighed Whatman GF/F filters. The filters were flushed with artificial saline water in order to remove dissolved OC and nitrogen. The flux of total particulate material was determined by the weight

gain of the filter after filtration and drying at 60°C. POC and particulate nitrogen content was determined (as % dry wt) from sub-samples of the dried filters, after fuming with HCl.

In the present study we used results only from the preserved traps, although they might have been compromised by migrating zooplankton caught in the traps. The unpreserved traps might, for their part, have been affected by zooplankton grazing and microbial degradation (Gundersen & Wassmann 1990). At 20 m depth at the fish farm, the carbon content in the preserved traps was 1.4 ± 0.2 times as high as in the unpreserved traps at the same depth. The preserved:unpreserved carbon content ratio at the reference station was 1.7 ± 0.3 . At the fish farm the C:N ratio of the sinking particles in the preserved and unpreserved traps was 11.2 ± 1.1 and 12.2 ± 1.3 , respectively, and at the reference station it was 7.1 ± 0.3 and 7.8 ± 0.3 , respectively. For assessment of pelagic–benthic coupling, the trap data were complemented with detailed benthic mapping of OC accumulation and mineralization rates within the footprint (see 'Sediment characteristics').

Sediment characteristics. Sediment was retrieved with a HAPS bottom corer (KC-Denmark; Kannevorf & Nicolaisen 1973). Only cores with a clear water phase were used. For further analysis sub-cores were collected in Plexiglas tubes (internal diameter = 5.6 cm). Cores were kept dark and at bottom-water temperature during transport to the laboratory, which was reached within 4 h of sampling. For determination of porosity, total organic carbon (TOC) and total nitrogen (TN) content, 2 cores from each station were sectioned into 1 cm intervals down to 4 cm depth, and 2 cm intervals from 4 to 12 cm depth. Sediment porosity was determined from density and water content measured as the weight loss after drying at 70°C for ~48 h. TOC and TN were measured on a CE 440 Elemental analyzer after the sediment had been homogenized, acidified (with 4 to 5% H_2SO_4), and dried. For macrofauna identification, sediment cores were sieved (1 mm mesh size) and preserved with formalin until later counting and identification down to lowest possible taxonomic level.

Sediment water fluxes. Upon return to the laboratory 3 sediment cores from each station were submerged in an incubation tank, holding bottom water at *in situ* temperature $\pm 0.5^\circ\text{C}$. On most sampling dates, the *in situ* bottom-water DO concentrations corresponded to 90% to 100% air saturated; water in the incubation tank was flushed with air, maintaining the ambient water at 100% air saturation. However, on 2 sampling dates, 22 August and 7 September 2006, the *in situ* DO concentration of the bottom water was only ~50% of air saturation. On 22 August, the incubation water was flushed with a mixture of nitrogen and air, keeping the oxygen concentration at 60% air saturation, but on

7 September (the date of the second transect measurements), the incubation water was mistakenly flushed with air only (no added nitrogen) and these data must therefore be treated with caution. In order to ensure well mixed conditions in the core liners, small Teflon-coated magnets were attached to the inner wall of the cores, their momentum imparted by an externally rotating magnet (Rasmussen & Jørgensen 1992). The stirring resulted in a measured effective diffusive boundary layer (DBL) thickness of $364 \pm 13 \mu\text{m}$ (data not shown).

After pre-incubation of 12 to 24 h the cores were closed, leaving an internal water height of approximately 8 cm. Oxygen concentration in the cores was monitored with a Clark-type oxygen minielectrode with a tip diameter of 500 μm , 90% response time of ~15 s and stirring sensitivity of <2% (Revsbech 1989, Glud et al. 1995a). Seawater samples for measurements of DO, DIC, DOC, NO_3^- and NH_4^+ were taken at the start of the incubation and when the oxygen concentration inside the respective cores had decreased 15 to 20%. Samples were collected with a gas-tight syringe equipped with a Tygon tube. Samples for NO_3^- , NH_4^+ , DO and DIC were stored and analyzed in the same way as the water column measurements. Samples for DOC measurements were filtered through pre-combusted (475°C) Whatman GF/F filters, and acidified with 1.2 M HCl for later analysis on a TOC analyzer (T5000 Shimadzu). Sediment water fluxes were calculated linearly from the change in solute concentration during incubation, accounting for incubation time and enclosed water volume. A linear decline in the DO concentration was confirmed from the continuous recording of the minielectrode. Sediment uptake is defined as a negative flux while sediment release is defined as a positive flux.

Oxygen microprofiles. Oxygen profiles were measured in the same cores 12 h after the flux measurements had been terminated by cap removal. Profiles were obtained with a Clark-type microelectrode equipped with a guard cathode and an internal reference (Revsbech 1989). The tip diameter was ~10 μm , stirring sensitivity <1% and the 90% response time was <1 s (Gundersen et al. 1998, Glud et al. 2000). Three profiles were measured in each core, adding up to a total of 9 microprofiles at each measuring event. The microelectrodes were positioned by a motor-driven micromanipulator and profiles were measured at a depth resolution of 50 or 100 μm . The sensor current was measured by a picoammeter connected to an analogue–digital converter, which transferred the signal to a PC (Revsbech & Jørgensen 1986). The microelectrode was calibrated by 2-point calibration from the signal in the well mixed air-saturated water and the signal of the anoxic sediment.

The oxygen penetration depth obtained from the oxygen microprofiles was defined as the depth between the sediment surface and the onset of the constant anoxic signal. The position of the sediment surface was estimated from a shift in the linear concentration gradient in the DBL of the individual profiles (Glud et al. 1995b). The diffusive oxygen uptake (DOU) was calculated from the calibrated microprofiles by $DOU = -D_0 \delta C / \delta z$, where D_0 is the temperature-corrected molecular diffusion coefficient and C the oxygen concentration at depth z within the DBL (Jørgensen & Revsbech 1985). Higher total oxygen uptake (TOU) than DOU is generally ascribed to fauna-mediated oxygen uptake in the form of respiration and irrigation (Archer & Devol 1992, Glud 2008).

Farming activity. Information on fish biomass and food usage was acquired from the farmer. The farming operation was controlled by a farm control monitoring system with automatic feeding (<http://www.akvagroup.com>).

Total carbon, nitrogen and water content in relation to fish size were determined on 14 rainbow trout *Oncorhynchus mykiss* (weight from 200 to 1700 g each) from the same tribe, fed the same food as the fish farmed in Kaldbaksfjörður. Individual weight and fork length were measured and the gut content removed, prior to mincing in a Bear Varimixer Type AR 10 (A/S Wodschow & Co, Denmark). Three subsamples from each fish were dried at 70°C until reaching constant weight. The water content was determined as the weight loss, and total carbon and nitrogen content was measured after the dried samples had been ground up and homogenized by hand. Dry wt, organic carbon and nitrogen content were also measured on both food pellets and faeces from juvenile (~90 g) and adult fish (1.5 to 2.5 kg). Faecal pellets were collected from sedimentation traps deployed at the farming site for 24 h, and dry wt, TOC and nitrogen content was determined on 6 individual pellets. Food samples used for the 2 fish sizes were obtained from the farmer.

RESULTS

Farming activity and sedimentation

Farming activity at this site started in late April, when ~770 000 juvenile rainbow trout *Oncorhynchus mykiss* (100 to 105 g each) were released into cages at the farming site. Approximately 35% of the fish were placed in the 4 square cages attached to the platform (Fig. 1c). The amount of feed usage at the platform gradually increased from this point into early July, after which the feed supply was held relatively constant at ~470 gC m⁻² d⁻¹ (Fig. 2a). From 2 October to 9 November there was no farming activity at the site.

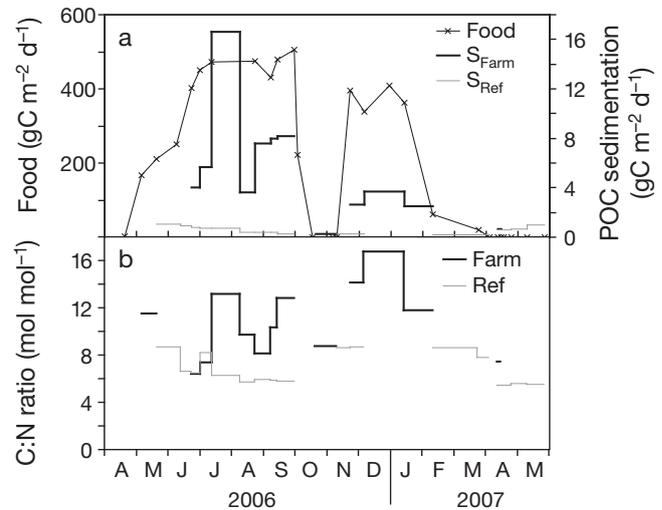


Fig. 2. (a) Daily carbon supply from food and particulate organic carbon (POC) sedimentation at fish farm (S_{Farm}), also showing POC data for reference station (S_{Ref}); and (b) C:N ratio of sedimenting material for both sites. All measurements made at 20 m depth in water column. Width of intervals given for sedimentation represents deployment time of traps

After the platform and cages containing fish were reinstalled, farming continued, with a feeding rate of ~380 gC m⁻² d⁻¹ until February 2007, when the majority of the fish stock was permanently removed. Approx. 61 500 specimens of 0.9 kg trout were left in the cages until April 2007, at a feeding rate of ~40 gC m⁻² d⁻¹. They were removed on 3 April 2007.

POC sedimentation was highly variable. During the period of relatively constant feeding rate, ~470 gC m⁻² d⁻¹, POC sedimentation varied between 3.6 and 16.6 gC m⁻² d⁻¹, while in the later farming period with a feeding rate of ~380 gC m⁻² d⁻¹, POC sedimentation was only between 2.5 and 3.7 gC m⁻² d⁻¹ (Fig. 2a). Thus no apparent relation existed between feeding rate and POC sedimentation directly below the farm (Fig. 2a). This was possibly due to temporally variable local hydrodynamics affecting the distribution of the sinking material, or possibly to a variable feeding rate. Sedimentation at the farm was nevertheless much higher than at the reference station. From July to September average sedimentation at the farming site was 9.7 ± 2.6 gC m⁻² d⁻¹ while at the reference station during the same period it only was 0.50 ± 0.09 gC m⁻² d⁻¹. Sedimentation was, moreover, markedly reduced at the farming site after the fish farm was removed (Fig. 2a).

During farming activity the C:N ratio of the settling material was clearly higher at the farm (12.2 ± 0.9) than at the reference station (7.1 ± 0.3 ; Fig. 2b). For comparison, the C:N ratio of fish feed used for juveniles was 7.6 ± 0.2 , and for adults (>1.3 kg), 10.2 ± 0.3 . The C:N ratio of faecal pellets of juveniles and adult trout was 13.2 ± 1.6 and 18.1 ± 0.6 , respectively. This indicates

that the majority of particulate wastes settling below and around the farm was faecal rather than feed pellets. Only occasionally was the C:N ratio relatively low, indicating trapped fish feed (Fig. 2b).

Since hydrography, primary production, chl *a* content and algal composition were highly similar between the farm and the reference station, it is reasonable to assume that natural sedimentation at the farm and reference station were similar as well. On the basis of this assumption, from 22 June to 29 September 2006 (the period of continuous sedimentation measurements at the farm site; Fig. 2a), of the total POC sedimentation of 914 gC m⁻², 5.6% was natural material. Using this same assumption, of total sedimentation of nitrogen during this period (101 gN m⁻²), 9.4% was of natural origin.

Rainbow trout carbon and nitrogen content

The water content of farmed rainbow trout was 70.5 ± 0.6%, a percentage that did not change with fish weight. However the weight % of C content showed a logarithmic increase with increasing fish weight, while the weight % of N showed a logarithmic decrease with increasing fish weight (Fig. 3). Overall the C:N ratio gradually increased from 5.0 for juveniles to 8.7 for adult fish (~1.5 kg).

Effects of fish farming on water column

Current velocity and direction was relatively uniform throughout the water column. From 23 March to 25 April 2006 current velocity ranged between 0.2 and 18.2 cm s⁻¹ at 6 m depth, with a mean (±SD) of 5.7 ± 5.7 cm s⁻¹, while at 3 m above the seabed (44 m depth)

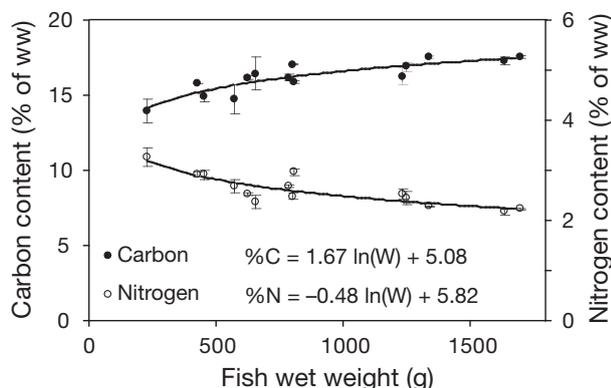


Fig. 3. Carbon and nitrogen content of rainbow trout relative to fish total weight including gut content (W). Values presented as mean ± SE of 3 subsamples from each individual. r² = 0.779 for carbon and 0.783 for nitrogen; p < 0.0005 for both relations

it ranged between 0.1 and 14.2 cm s⁻¹, with an average (±SD) of 4.2 ± 4.4 cm s⁻¹. The dominant current direction at the farming site was into the fjord from the sound (Table 2).

DO at 5 m depth at the farm site was usually supersaturated and within ±8% of the DO concentration at the reference station (Fig. 4a). However, from July to mid-September 2006 the oxygen concentration at the farm site was 11 to 26% lower than levels at the reference station. This coincided with the highest measured ammonium concentration in the surface water at the fish farm (Fig. 4b). While both the farm site and the reference station showed similar ammonium concentrations when there was no farming, in periods with active farming, the ammonium concentration at the fish farm was on average 3.9 ± 1.2 μM higher than at the reference station. Generally, however, the relationship was weak (r² = 0.28, linear regression) between feeding activity and ammonium concentration in the upper water column. Nitrate concentrations in the surface water did not differ significantly between the 2 sites (Fig. 4c).

In bottom waters, DO, ammonium and nitrate concentrations were similar between the 2 sites (Fig. 4). During summer 2006, a weak thermocline was found at ~40 m depth, and consequently the DO concentration in the bottom water at both stations decreased to a minimum of 48% of air saturation and the ammonium concentration increased (Fig. 4d,e). The decrease in bottom water nitrate concentration during summer (Fig. 4f) was due to decreased nitrate content in inflowing water from the sound outside the fjord (data not shown).

Despite higher nutrient availability at the farm site, chl *a* concentration, algal composition and primary production showed no local productivity increase (Fig. 5). Annual primary production at the farming site was 352 gC m⁻², while the corresponding value for the reference station was 335 gC m⁻².

Sediment description

Prior to initiating farming activities, the sediment at the 2 sampling sites was sampled and found to be similar in physical and chemical characteristics (Table 1).

Table 2. Mean, maximum and residual current velocities and their direction at 3 depths close to the farming site

Depth (m)	Velocity (cm s ⁻¹)			Residual dir. (°)	
	Mean	Max.	SD		
6	5.7	18.2	5.7	4.9	287
26	4.3	12.2	4.4	3.5	270
44	4.2	14.2	4.4	2.3	290

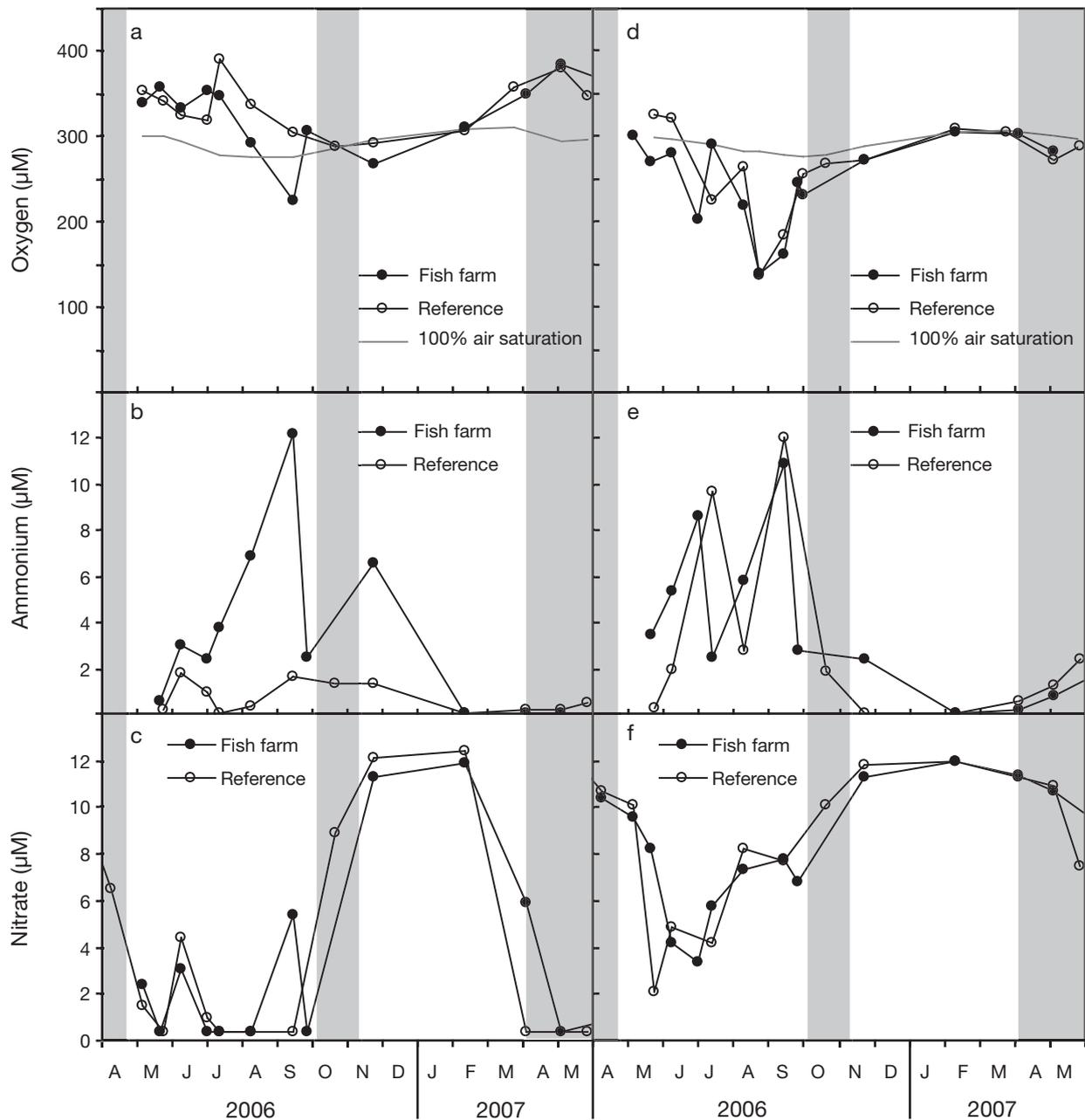


Fig. 4. (a,d) Oxygen, (b,e) ammonium and (c,f) nitrate concentration in (a,b,c) surface waters, and (d,e,f) bottom waters at fish farm and reference station. Grey lines in (a,d) represent oxygen concentration of 100% air-saturated water at given salinities and temperatures. Vertical grey zones indicate periods with no farming activity. SD on duplicate measurements not included since error bars in most cases were smaller than symbol size. Data on reference station from Gaard et al. (2011)

Farming activity appeared immediately to change this. First, distinct faecal pellets (and occasional feed pellets) were observed, along with patches of the sulphur bacteria *Beggiatoa* spp. (Jørgensen & Revsbech 1983). Later the whole sediment surface at the farm site became covered with mats of *Beggiatoa* spp., and the underlying sediment turned black. The black zone, at first noted at the top of the sediment, progressed

downward as farming activity continued, with faecal pellets spreading sporadically throughout the black layer. In August the sediment was black to the maximum sampling depth of >18 cm, and gas bubbles were observed from ~6 cm on down. In early November 2006, following 39 d of no farming activity, the sediment exhibited signs of recovery, becoming light grey to a depth of ~1 cm, but remaining black in the deeper

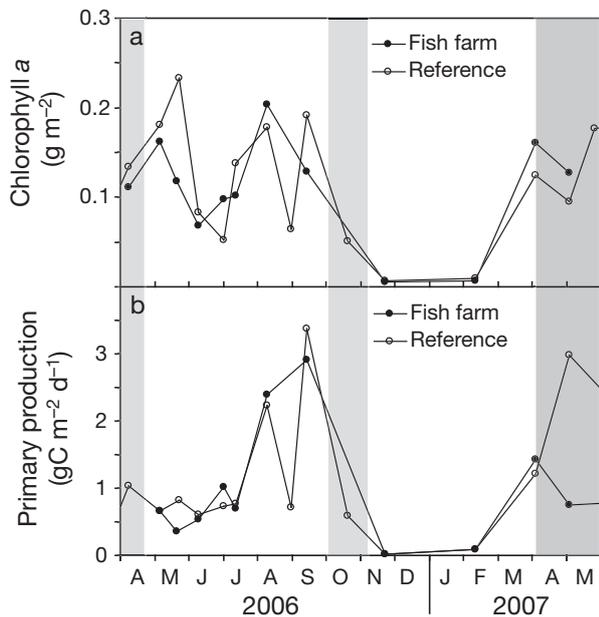


Fig. 5. (a) Chl *a* content (from 0–20 m depth), and (b) primary production at fish farm and reference station. Vertical grey zones indicate periods with no farming activity. Data on reference station from Gaard et al. (2011)

layers. Burrowing fauna, dominated by the H₂S-tolerant polychaete *Capitella capitata*, was observed in the cores, but the burrow sides still hosted *Beggiatoa* spp. On 5 December 2006, <1 mo after farming activity was resumed, the sediment again appeared highly impacted. The surface was again covered by *Beggiatoa* spp., the sediment was black down to 12 cm, and gas bubbles reappeared from ~6 cm depth downwards. In April and May 2007, after 2 mo with a low feeding rate and then cessation of farming activity, the sediment showed improvement similar to what had been observed in November 2006.

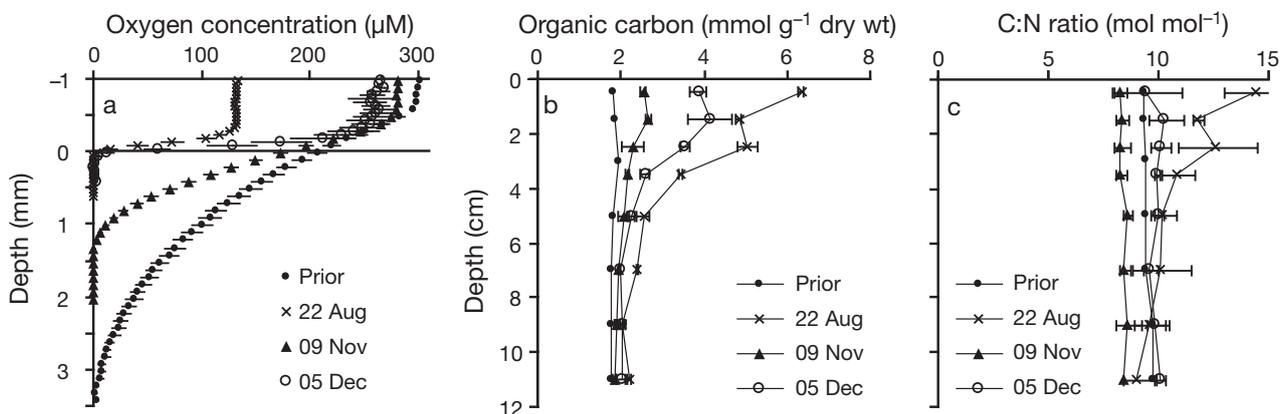


Fig. 6. (a) Oxygen microprofiles as mean (\pm SE) of 7–9 profiles in 3 cores, (b) carbon content as mean (\pm SD) of 2 cores, and (c) C:N ratio as mean (\pm SD) of 2 cores for 4 sampling dates in 2006: prior to farming, after ~4 mo of active farming (22 August), after 39 d without farming activity (9 November), and 1 mo after re-establishment of farming (5 December). Points lacking error bars indicate non-replicated samples

Prior to commencement of farming activity the oxygen penetration depth into the sediment varied, but the average penetration depth was >3 mm (Fig. 6a). During farming, DO microprofiles became less variable, with oxygen penetration depth <0.2 mm (Fig. 6a). However, during October 2006, when no farming activities were being conducted, conditions improved substantially, and after 39 d without farming the oxygen penetration had increased to 1.3 ± 0.1 mm (Figs. 6a & 7a). Some increase of the oxygen penetration depth was likewise observed in April and May 2007 when farming activity had stopped, although not at the same magnitude as during the recovery event in October–November 2006 (Fig. 7a).

Sediment OC content also exhibited significant changes in relation to farming activity. Prior to initiation of farming activity, the average OC content was 1.8 mmolC g^{-1} (dry wt, top 12 cm), a value which held relatively constant down through the sediment (Fig. 6b). After 4 mo of fish farming, the sediment was highly enriched with OC. The average OC content in the upper 4 cm at this time was $4.9 \pm 1.2 \text{ mmolC g}^{-1}$ (dry wt), and the sediment remained enriched in OC during farming activity. After 39 d without farming activity, the OC content of the sediment decreased substantially, and elevated OC content was observed only in the top 2 cm of the sediment layer, but when farming was re-established OC content began to increase again. The C:N ratio showed a similar trend, with elevated values during farming periods and declining values in periods with reduced or no farming activity (Fig. 6c).

Total and diffusive oxygen uptakes at the farm site generally reflected a similar pattern and were of similar magnitude at all times (Fig. 7b), indicating that fauna irrigation contributed only marginally to benthic solute exchange. Oxygen uptake increased with increasing

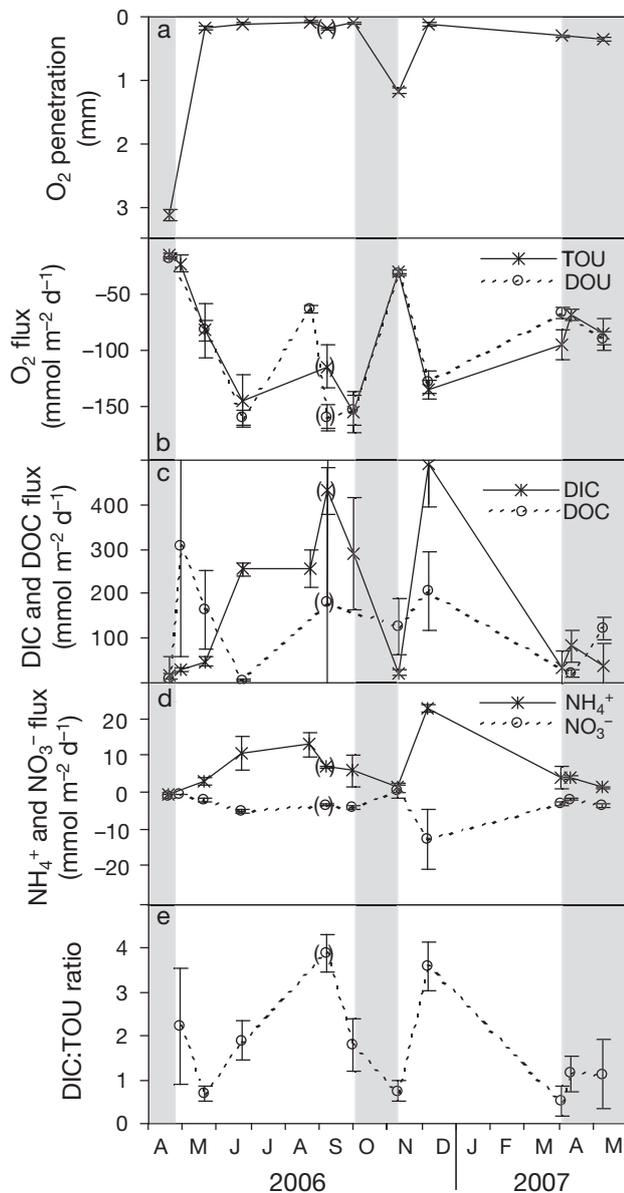


Fig. 7. Time series of (a) oxygen penetration depth into sediment, (b,c,d) sediment water exchange rates for several variables (DOC: dissolved organic carbon), and (e) ratio between dissolved inorganic carbon (DIC) efflux and total oxygen uptake (TOU) at farming site from April 2006 to May 2007. Negative fluxes represent sediment uptake. Error bars indicate SE. For oxygen penetration depth and diffusive oxygen uptake (DOU), $n = 9$; for other measurements, $n = 3$. Fluxes measured on September 7 (indicated with parentheses) were incubated with 100% air-saturated seawater although, *in situ*, bottom water was only 50% saturated. Vertical grey zones indicate periods with no farming activity

farming activity, in June 2006 reaching a maximum of $160 \pm 7 \text{ mmol m}^{-2} \text{ d}^{-1}$. In August the oxygen uptake transiently decreased, presumably due to reduced DO (50% of air saturation) in the bottom water, limiting the potential high oxygen uptake rate of the carbon-

enriched sediment. After the October recovery period the oxygen uptake was only $30 \pm 4.5 \text{ mmol m}^{-2} \text{ d}^{-1}$, increasing to the former level when farming was re-established (Fig. 7b). Similar trends were also observed in DIC flux (Fig. 7c). DIC efflux was, however, markedly higher than oxygen uptake in the sediment, and this difference increased with increasing exchange rates (Fig. 7e). During the entire measuring period of 387 d (linear integration between measuring points), the total oxygen uptake was 40.6 mol m^{-2} , while the DIC efflux was twice as high, viz. $81.7 \text{ mol C m}^{-2}$. DOC flux was highly variable, both in time and between replicate measurements (Fig. 7c), and did not show similar development over time as the other parameters, but in all measurements there was an efflux of DOC from the sediment, which, integrated over the entire study period, amounted to $43.4 \text{ mol C m}^{-2}$.

An NH_4^+ efflux from the sediment was evident at all times, increasing with increasing DIC release rate (Fig. 7). The C:N ratio of the measured effluxing solutes was 70 ± 20 , while the C:N ratio of the settling material was 11 ± 1 . This apparent mismatch could be related to (1) denitrification/anammox activity, (2) release of dissolved organic nitrogen (DON) or (3) preferential nitrogen incorporation into bacterial biomass. There was almost always a sediment uptake of NO_3^- (from 0.46 to $12.2 \text{ mmol m}^{-2} \text{ d}^{-1}$). Two exceptions were during the October 2006 recovery period, with a release of $0.86 \pm 0.09 \text{ mmol m}^{-2} \text{ d}^{-1}$, and at the very beginning of the farming season when the release was $0.22 \pm 0.03 \text{ mmol m}^{-2} \text{ d}^{-1}$.

Benthic footprint of farming activity

Sediment was visibly impacted 30 m from the farm along the transects. Within this distance, the sediment was black and the surface was covered by *Beggiatoa* spp. At 60 m distance, the only visible farming impact was that the top 3 cm of the sediment were slightly darker than at the reference station. Sediment at both sites contained macrofauna, although the composition differed between sites: the dominant fauna at the farm site was *Capitella capitata*, while the reference station was dominated by the bivalves *Abra nitida* and *Thyasira flexuosa*. At 90 m distance from the farm no signs of farming activity were observed.

The oxygen concentration in the sediment reflected the organic load. Within 20 m of the fish farm, the oxygen concentration changed rapidly in the diffusive boundary layer, and at the sediment surface the oxygen concentration was $<10 \mu\text{M}$ (Fig. 8a,b). The oxygen penetration into the sediment increased with increasing distance from the farm, from $<0.2 \text{ mm}$ below the farm to $1.7 \pm 0.1 \text{ mm}$ at 90 m distance. At the reference

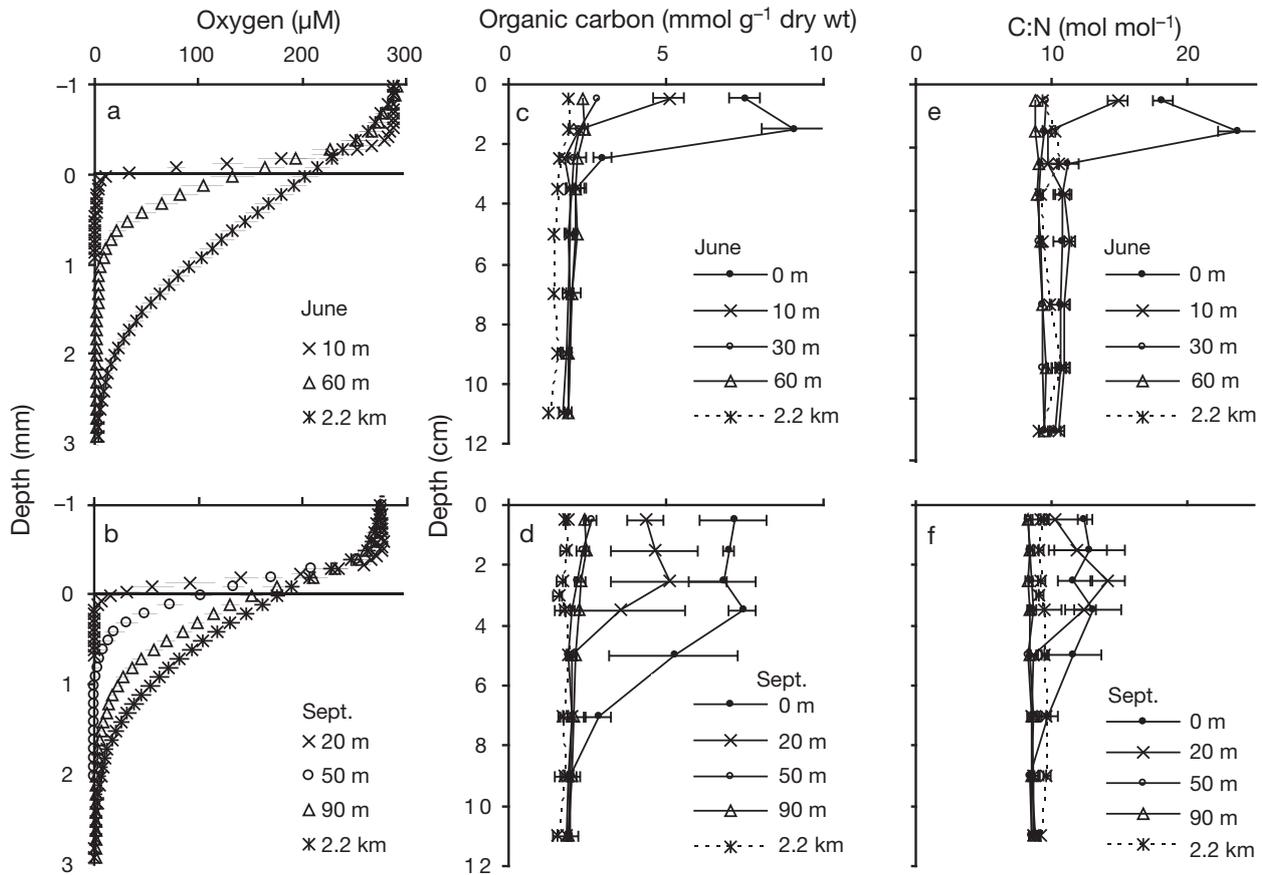


Fig. 8. Sediment characteristics at various distances from fish farm on (a,c,e) 22 June, and (b,d,f) 7 September 2006. Oxygen microprofiles presented as mean \pm SE of 7–9 profiles in 3 cores, carbon content as mean of 2 cores \pm SD, and C:N ratio as mean \pm SD of 2 cores. Points lacking error bars indicate non-replicated samples

station oxygen reached more than 2.5 mm into the sediment (Fig. 8b).

The OC content of the sediment was highest at the surface and decreased with sediment depth and distance from the fish farm (Fig. 8c,d). On 22 June 2006 the average OC of the top 3 cm of the sediment was 6.5 mmolC g^{-1} (dry wt), while the content of the deeper sediment was 2.0 mmolC g^{-1} (dry wt). At 10 m from the farm the upper 2 cm showed an average OC content of 3.7 mmolC g^{-1} (dry wt) (Fig. 8c), and at 30 m distance only the upper 1 cm exhibited elevated levels of OC (2.8 mmolC g^{-1} dry wt). At a distance of 60 m the OC content was fairly constant with depth and resembled conditions at the reference station. However, during the 77 d that separated measurements along the 2 transects the profiles of OC content changed considerably, reflecting a continuous accumulation of OC near the fish farm (Fig. 8c,d). The C:N ratio of the sediment showed a similar pattern, with an elevated C:N ratio in sediment affected by farming activity (Fig. 8e,f).

Sedimentation measurements were conducted only along the second transect, although the feeding rates

were fairly similar when each transect was sampled (Fig. 2a). During measurements along the second transect the feeding rate was $430 \text{ gC m}^{-2} \text{ d}^{-1}$, and the vertical carbon flux (POC) below the farm at 20 and 40 m depth was 4.3 and $4.9 \text{ g m}^{-2} \text{ d}^{-1}$, respectively (Fig. 9a). The POC flux decreased rapidly within 20 m of the farming site. At this distance from the site, at 20 and 40 m depth, it was 0.4 and $1.2 \text{ g m}^{-2} \text{ d}^{-1}$, respectively. At 90 m from the farm site it was only slightly higher than at the reference station. The C:N ratio of the sinking material showed a similar trend, with decreasing ratio with distance from the farm; at 90 m distance it was only slightly higher than at the reference station.

Like oxygen penetration depth (Figs. 8a,d & 9b), DOU decreased rapidly with distance from the farm, while TOU remained fairly constant within 90 m of the fish farm. Thus the difference between TOU and DOU increased with distance, reflecting higher faunal activity further away from the farm (Fig. 9c). Oxygen fluxes and penetration depth were similar at the 2 transects (Fig. 9b,c). However, the DIC flux in the vicinity of the farm increased considerably between the 2 sampling

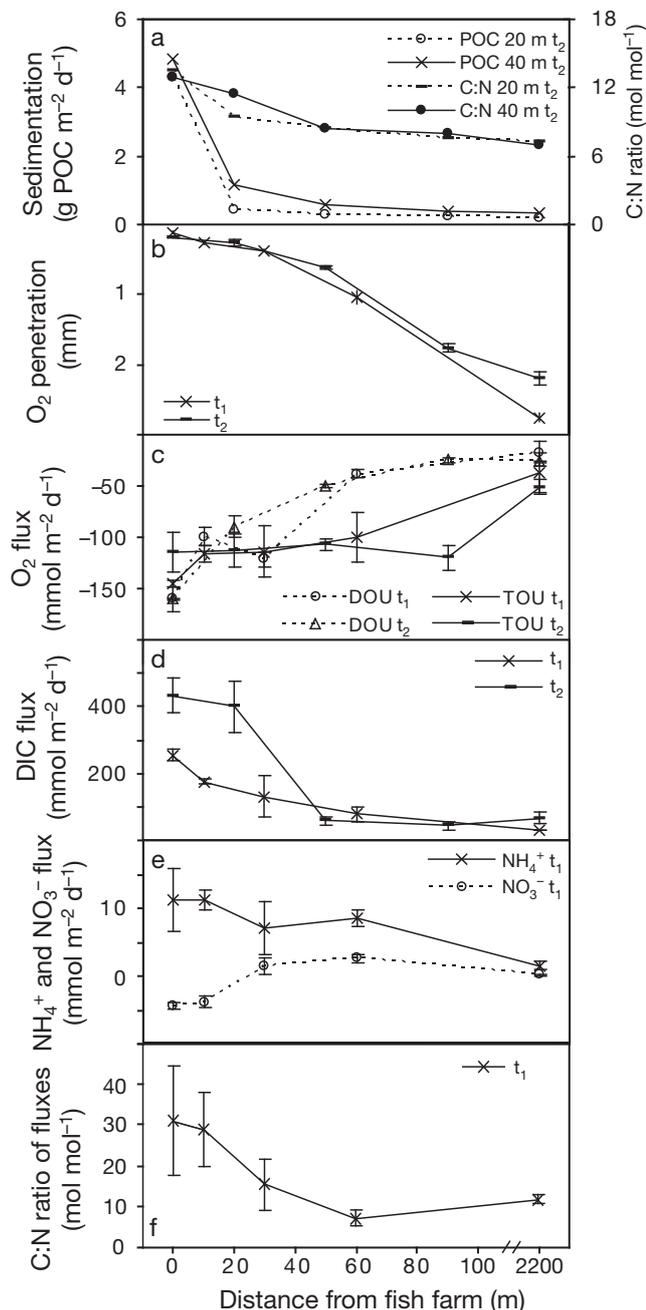


Fig. 9. Measurements along transect of (a) POC sedimentation and C:N ratio of sedimenting material, (b) oxygen penetration depth into sediment, (c,d,e) sediment water exchange rates for several variables, and (f) ratio between DIC and DIN effluxes at various distances from the fish farm on 22 June 2006 (t_1) and 7 September 2006 (t_2). Abbreviations as in Fig. 7. Negative fluxes represent sediment uptake. Error bars indicate SE. For oxygen penetration depth and DOU, $n = 9$; for other measurements, $n = 3$

dates (Fig. 9d). On 22 June 2006 the DIC flux was $255 \pm 16 \text{ mmol m}^{-2} \text{ d}^{-1}$ at the farming site, and it decreased exponentially with distance. Seventy-seven days later DIC was $432 \pm 52 \text{ mmol m}^{-2} \text{ d}^{-1}$ at the farm, decreasing

to $399 \pm 79 \text{ mmol m}^{-2} \text{ d}^{-1}$ 20 m from the farm, and dropping to $58 \pm 12 \text{ mmol m}^{-2} \text{ d}^{-1}$ 50 m from the farm. The ratio between DIC efflux and TOU at the farming site increased from 1.8 to 3.8 in the 77 days between the 2 transect measurements. On 22 June 2006 the ammonium efflux decreased slightly with distance, from $11.3 \pm 4.6 \text{ mmol m}^{-2} \text{ d}^{-1}$ at the farm to $8.6 \pm 1.2 \text{ mmol m}^{-2} \text{ d}^{-1}$ 60 m away; sediment ammonium release at the reference station was $1.5 \pm 0.6 \text{ mmol m}^{-2} \text{ d}^{-1}$ (Fig. 9e). The nitrate flux changed from a sediment uptake of $4.4 \pm 0.5 \text{ mmol m}^{-2} \text{ d}^{-1}$ at the farming site to an efflux of $2.6 \pm 0.5 \text{ mmol m}^{-2} \text{ d}^{-1}$ 60 m away. At the reference station there was a small nitrate efflux of $0.2 \pm 0.1 \text{ mmol m}^{-2} \text{ d}^{-1}$. The C:N ratio of the DIC and DIN effluxes decreased rapidly from 30 ± 17 at the farm to 6.9 ± 1.8 , 60 m from the farm. These values are comparable to the C:N ratio of the sedimenting material (Fig. 9).

DISCUSSION

Effect of farming activity on water column

From July to mid September the DO concentration at the farming site was 11 to 26% lower than at the reference station (Fig. 4a). This offset was not caused by differences in primary production, which was the same at the 2 sites (Fig. 5), but may rather be ascribed to localized respiration of the fish. An estimate based on temperature-corrected respiration rates for medium-sized *Salmo salar* swimming 0.75 bl s^{-1} (Grøttum & Sigholt 1998) amounted to an oxygen requirement of the fish stock on the order of 360 and 740 mol h^{-1} for June and September, respectively. In order to maintain the measured DO concentration in the cages, ambient current velocities in the range of 0.2 to 0.5 cm s^{-1} were required, assuming that current velocity is reduced by 10 to 35% during passage through the net (Pillay 2004, Patursson et al. 2010). Current velocities below 2 cm s^{-1} were measured only 6.5% of the time during March and April 2006. During summer, lower current velocities might prevail but apparently they remained sufficient to counteract any severe DO depletion within or around the farm (Fig 4a).

Ammonium is the main dissolved nutrient excreted from farmed fish; concentrations of other nutrients (e.g. nitrate, phosphate) apparently remain unaffected by farming activity (Brooks & Mahnken 2003, Sander et al. 2008). In the present study, the average ammonium concentration during farming activity was $4.0 \pm 1.3 \text{ } \mu\text{M}$, which is 4.3 times higher than at the reference station. The highest ammonium concentration ($12.4 \text{ } \mu\text{M}$) was observed at the same time as the minimum DO concentration ($224.6 \text{ } \mu\text{M}$, 81% saturation), indicating a transiently lower water exchange

during the measurements. Generally, ammonium concentrations at fish farms reach values of 2 to 10 μM (Pitta et al. 2006, 2009, Sanderson et al. 2008), even at sites with relatively efficient water exchange (Merceron et al. 2002). However, in agreement with most studies (Brooks & Mahnken 2003, Islam 2005, Pitta et al. 2006), we did not see any localized stimulation of primary production (Fig. 5).

The flushing time through the farm was about 6 min, while the specific primary production at the site implies an average phytoplankton doubling time of 2 d (Gaard et al. 2011). Thus no local effects would be expected. This supports the conclusion by Brooks & Mahnken (2003) that the time required for phytoplankton cells to divide implies that the cells will be transported away from the nutrient point source before any local growth response can be detected. This makes it difficult to detect the possible effect on primary production from point sources in open-water bodies. However, on regional scales, farming may affect primary production depending on e.g. hydrographic features and nutrient status in the euphotic zone (see 'Carbon and nitrogen budget for the fish farm').

Effects of farming activity on benthic mineralization rates

There was an obvious accumulation of particulate wastes below the fish farm (Fig. 6), leading to reduced oxygen penetration and elevated benthic solute exchange rates, both reflecting the stimulated mineralization rates (Fig. 7).

The benthic mineralization rate below the fish farm was controlled by the input of particulate organic wastes from farming activity, and apparently did not reach a saturation level, as has been observed in studies with an even higher carbon load (Holmer et al. 2003). The ratio between DIC efflux and TOU increased dramatically with the organic loading. This reflected increasing accumulation of anaerobic metabolites (viz. FeS , FeS_2 and even H_2S and CH_4 in the form of bubbles) that are not being oxidized by the limited oxygen supply (Hall et al. 1990, Findlay & Watling 1997, Holmer et al. 2002, 2003). We cannot exclude a minor contribution to DIC from enhanced dissolution of shell debris in the reduced sediments. During active farming, the sediment was covered by *Beggiatoa* spp., and the oxygen penetration depth was $<200 \mu\text{m}$. Thus at this point most sediment oxygen uptake was presumably for oxidation of reduced chemical species produced during anaerobic bacterial reduction (e.g. sulphide) rather than for aerobic heterotrophic activity (Jørgensen 1982). The similarity between DOU and TOU (Fig. 7b) below the farm implies little active

macrofauna; it also implies that oxygen is supplied to the sediment mainly by diffusion (Glud et al. 2003).

On average, the DOC efflux amounted to 53 % of the DIC release rate, which is higher than the 10 to 20 % generally encountered in coastal and shelf sediments (Hall et al. 1990, Viollier et al. 2003). However, the importance of DOC release has been found to increase with the reducing state of the sediment, mainly as a result of incomplete oxidation, as respiration processes struggle to keep up with hydrolysis and fermentation (Hansen & Blackburn 1991). DOC efflux has also been found to increase in importance when the material is degraded close to the sediment surface (Blackburn & Blackburn 1993, Blackburn et al. 1996, Hulth et al. 1997, Fenchel et al. 2000). In the present study, the sediment was highly reduced and enriched with OC, and consequently macrofauna was absent or scarce (Table 1). Thus a potentially large portion of the POM was mineralized close to the sediment surface, leading to high DOC effluxes. In fact DOC release may be a more important pathway for benthic carbon release below aquaculture farms than previously thought.

Nitrate uptake of the sediment increased linearly with increasing mineralization rates ($r^2 = 0.62$, $p < 0.005$). Since the carbon-enriched sediment was covered with *Beggiatoa* spp. during active farming, some of this nitrate uptake may be attributable to intercellular bacterial storage (McHatton et al. 1996). However, stimulated denitrification or dissimilatory nitrate reduction to ammonia (DNRA) in carbon-enriched, oxygen-depleted sediment below fish farms has also been documented (Christensen et al. 2000, Holmer et al. 2003). The strong correlation ($r^2 = 0.76$, $p < 0.0005$) between nitrate uptake and ammonium release could indeed indicate a dominant role for DNRA (Nishio et al. 1983, Christensen et al. 2000, Gardner & McCarthy 2009). Overall, the sediment served as a source of dissolved nitrogen, but our nitrogen release rates do not reflect the total mineralization of nitrogen, which could also be released as DON or N_2 , pathways we did not investigate in the present study. The high C:N ratio (70 ± 20) of measured effluxing solutes below the fish farm indicates that some other form of nitrogen release might be quantitatively important. Indeed, DON appears to be the dominant form of nitrogen released from organic-enriched sediments (Hall et al. 1992, Blackburn & Blackburn 1993, Sloth et al. 1995, Blackburn et al. 1996, Fenchel et al. 2000).

Carbon and nitrogen budget for the fish farm

To establish an overall budget for carbon and nitrogen flow through the farming area we focused on the 77 d between the 2 transect measurements (from

22 June to 7 September 2006). During this period sediment traps were deployed 5 times, and the data were complemented by detailed mapping of carbon and nitrogen enrichments in the sediment.

The total carbon and nitrogen input with fish feed in the period of interest was 62.7 t carbon and 9.5 t nitrogen (Fig. 10). The average fish weight increased from 240–280 to 600–950 g, with a total fish biomass increase (including dead fish) of 123 t, amounting to a carbon and nitrogen incorporation of 20.8 and 3 t, respectively.

The inorganic carbon excretion from the fish was estimated from oxygen consumption of *Salmo salar* correlated to weight, temperature and swimming speed (Grøttum & Sigholt 1998), and from the respiration quotient of 0.8 presented by Forsberg (1997). Assuming an average swimming speed of 0.5 to 1.0 bl s⁻¹ (Juell 1995, Cooke et al. 2000), the total inorganic carbon release in the 77 d was from 33 to 44 t (Fig. 10). Most probably there is an additional carbon release to the water column in the form of organic solutes lost from sinking food and faecal pellets (Chen et al. 1999, Fernandez-Jover et al. 2007).

The impacted sediment area was estimated from the sediment exhibiting elevated diagenetic activity along the transects, and from current speed and direction. An oval benthic footprint was suggested, with the longer axis lying parallel to the transect (Fig. 1, Table 2). The oxygen penetration depth was approximated by an exponential increase with distance from

the farm along the 2 transects ($r^2 = 0.95$ and 0.99); the DOU was also approximated well with an exponential decline ($r^2 = 0.88$ and 0.99). Likewise, DIC fluxes decreased exponentially with distance in the first transect ($r^2 = 0.98$; Fig. 9). These exponential extrapolations predicted that reference conditions were reached at a distance of 85 to 93 m and 99 to 103 m from the farm along transects t_1 and t_2 , respectively.

The accumulation of farm-derived OC and nitrogen in the sediment was estimated from the OC and nitrogen that had accumulated in sediment during the period between the 2 transect measurements. The area-based amount of OC and nitrogen (mmol m⁻²) was obtained from the content per unit dry wt of sediment (Fig. 8), and the density and water content of the sediment. The total accumulation in the footprint area was subsequently determined, assuming that the content in the sediment below the cage was represented by the measurement at the edge of the cage, while measurements along the transect represented content at various distances from the farm.

Approx. 32% of the OC and nitrogen that had accumulated in the sediment was found directly below the ~2000 m² fish farm, while the remaining OC and nitrogen was located elsewhere in the footprint area. Correspondingly 31% of the DIC efflux from mineralization of fish farm wastes occurred directly below the fish cages. Due to lack of DOC measurements along the transects, it was assumed that DOC was proportional to DIC release rates, amounting to 53% of DIC efflux

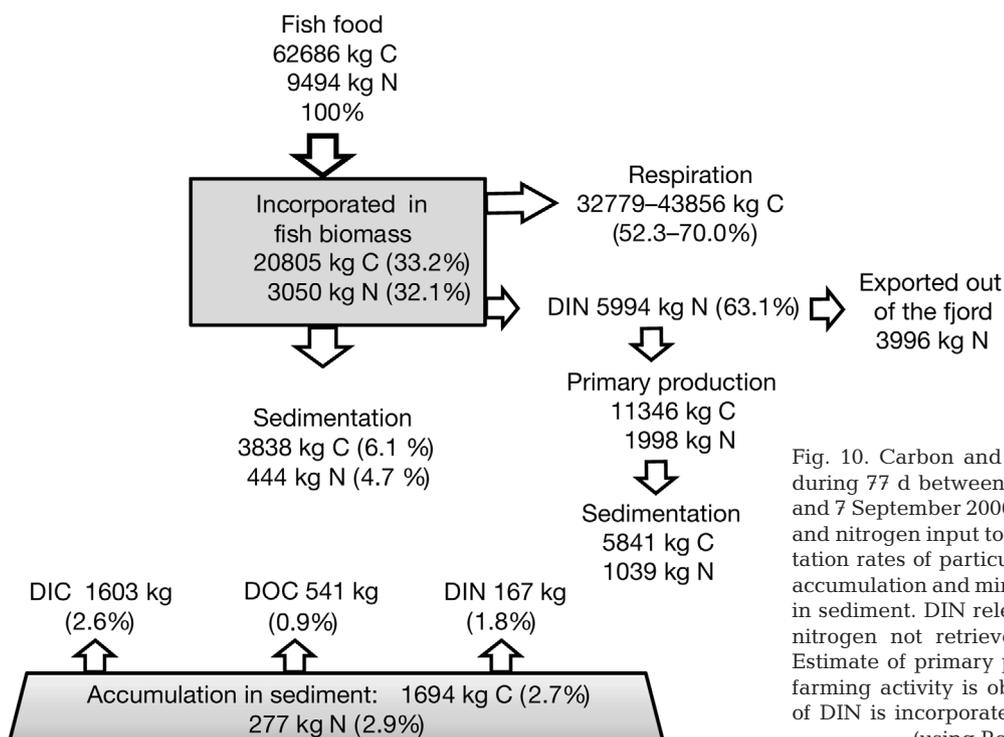


Fig. 10. Carbon and nitrogen budget for fish farm during 77 d between sampling 2 transects (22 June and 7 September 2006). Percentages are % of carbon and nitrogen input to farm in form of feed. Sedimentation rates of particulate farm wastes are based on accumulation and mineralization of OC and nitrogen in sediment. DIN release from farm is calculated as nitrogen not retrieved in biomass and sediment. Estimate of primary production based on DIN from farming activity is obtained by assuming one-third of DIN is incorporated into phytoplankton biomass (using Redfield ratio of C:N)

as measured directly below the farm. Only 10% of total DIN efflux occurred directly below the fish cages, with a much larger percentage occurring over the rest of the footprint. This was to a large extent due to the change from high-nitrate sediment uptake at the fish farm to nitrate release at some distance from the farm (Fig. 9e). It must be emphasized, however, that the benthic nitrogen budget does not account for DON exchange, denitrification or anammox.

During the 77 d between transect measurements, total OC accumulation in the footprint area was 1694 kg, and total nitrogen 227 kg. Total mineralization of carbon to DIC was 1603 kg, while carbon leaving the sediment as DOC amounted to 541 kg (Fig. 10). The DIN efflux was 167 kg. This suggests that the total OC and nitrogen sedimentation was at least 3838 kg carbon and 444 kg nitrogen (Fig. 10). Total sedimentation at the fish farm and along the transect was underestimated by 37 to 40% relative to the accumulation and effluxes of carbon and nitrogen at the seabed. This demonstrates that care should be taken in extrapolating trap data from point sources to a larger area in heterogeneous environments.

Earlier studies of salmon and trout farms have found environmental carbon loss of 78 to 84% (Gowen & Bradbury 1987, Hall et al. 1990, Ackefors & Enell 1994, Blackburn et al. 1996) and nitrogen loss of 72 to 81% (Gowen & Bradbury 1987, Hall et al. 1992, Islam 2005); both are higher than the 67% carbon and 68% nitrogen loss in the present study. The difference is most likely a result of improved farming and feeding management during the last 2 decades, resulting in reduction of environmental impact per unit cultured fish (Brooks & Mahnken 2003).

While carbon input to the farm in the form of feed is accounted for in fish biomass (33%), in sedimentation (6%) and through respiration (from 52 to 70%), only 37% of nitrogen is accounted for by fish biomass and sedimentation (Fig. 10). Some nitrogen might have escaped the sediment as DON or through denitrification. This loss can be estimated (assuming the C:N ratio of trap material to be representative of particulate matter settling on the seabed), and this estimate may be multiplied by the sediment-derived accumulation of OC, thereby deriving the total sedimentation of nitrogen. Surprisingly this value aligns well with the sum of the benthic inventory and net exchange of nitrogen (leaving only 6 kg of nitrogen unaccounted for across the benthic footprint)—despite the fact that measurements directly below the farm indicated that DON and N₂ effluxes might be substantial. In sediments with conditions similar to those of the present study, i.e. high organic load, large ammonium efflux from the sediment, nitrate uptake, and presence of *Beggiatoa* spp., the N₂ efflux amounts to 2 to 6% of DIN efflux

(Christensen et al. 2000, 2003, Bissett et al. 2009), which agrees well with the small amount of 'missing' nitrogen in the present study.

The C:N ratio of DIC and DIN effluents from the sediment decreased considerably within a short distance from the farm (Fig. 9f). Below the farm it was 30 ± 17 , decreasing exponentially ($r^2 = 0.987$) to 7 ± 3 at a distance of 60 m in the first transect. Thus the relative area of highly impacted sediment directly below the farm in relation to total footprint area implies that in total, the DIN flux approximates total mineralization quite well. However, since the carbon budget is balanced while a large proportion of the nitrogen is unaccounted for, most of this nitrogen was probably lost directly to the water column in an inorganic form.

Most carbon input with feed ended up as inorganic carbon due to fish respiration, while only 6.1% of the carbon input fell to the seabed (Fig. 10). The fraction of OC input settling to the seabed was less than measured or predicted in most other studies, where values have ranged between 29 and 71% (Gowen & Bradbury 1987, Hall et al. 1990, Bergheim & Asgard 1996). The fraction is even lower than the calculated fraction of 8.8%, which assumed that 5% of the food input settles directly to the seabed (Brooks & Mahnken 2003). This suggests a small food loss at the farm, which is also supported by the high C:N ratio of settling particles. In other studies, the fraction of nitrogen input settling to the seabed is 11 to 28% (Gowen & Bradbury 1987, Hall et al. 1992, Bergheim & Asgard 1996), likewise higher than the 5% in the present study; however, we did not account for any DON or nitrogen efflux from the sediment. In addition we found higher remineralization efficiency than in comparable studies. In the present study, microbial processes in the sediment removed 56% of OC and 38% of nitrogen input to the sediment, while other studies found that microbial processes removed only from 3 to 25% of the OC input, and from 11 to 16% of the nitrogen input (Hall et al. 1990, 1992, Holmer et al. 2002, Carlsson et al. 2010). However, in those studies the input was measured by sedimentation traps moored close to the seabed, a factor which may have led to measurement of resuspended particles. Our mineralization efficiency is comparable to the controlled laboratory study by Valdemarsen et al. (2010), who found that 34 to 56% of OC added in the form of fish feed was released from the sediment as DIC and that 40 to 46% of nitrogen from fish feed was released as ammonium.

Assuming that the mass balance is valid for all the farmed fish in the fjord, the DIN released from these fish during the season with sufficient light for primary production (from April to September 2006 and from April to May 2007) amounted to 98 700 kg. At the same time, the DIN input into the euphotic layer from other

sources (upwelling and runoff from land) is estimated to be ~606 000 kg (Gaard et al. 2011). Thus DIN release from farming activity represented ~14 % of total DIN input into the euphotic layer. Estimating the impact from the nutrient release by farmed fish on primary production is subject to large uncertainties, since a variable fraction of DIN input into the euphotic layer is exported without being incorporated into phytoplankton within the fjord (Gaard et al. 2011). Assuming that all available DIN was assimilated according to the Redfield ratio would give a new primary production figure (Dugdale & Goering 1967) of 740 gC m⁻². However, Gaard et al. (2011) found the new primary production to be ~250 gC m⁻². Thus, on average, approximately one-third of the DIN input into the euphotic layer was assimilated, while the remaining two-thirds was exported during the productive seasons. Due to variable nutrient concentrations in the euphotic layer, however, and to variable nutrient export rates, the influence on primary production is highly variable with time. When vertical mixing is high, additional DIN input (e.g. from fish farming) is expected to have minor influence on local primary production, and most of the nutrients are exported with the upper layer outflow. However, during extended calm periods, when nutrients in the upper layer often are depleted (Fig. 4c), most of the DIN release from farmed fish is probably assimilated within the fjord. Under these circumstances, the relative assimilation of anthropogenic nutrients is expected to be substantially higher than the long-time average of one-third of the DIN input, and this could potentially increase primary production substantially.

Recovery

During the 39 d pause in farming at this site (October to November 2006) substantial improvements in sediment conditions were observed (Figs. 6 & 7).

The OC content in the top 6 cm of the sediment decreased significantly during the recovery period (Fig. 6), from 83.3 to 65.4 mol m⁻². During the same period the sedimentation of POC was 0.8 mol m⁻², and thus the total removal of OC during the 39 d can be estimated as 18.7 mol m⁻², amounting to 22.4 % of OC content at the beginning of the recovery period. It is generally anticipated that benthic mineralization rates decrease exponentially after a pulse-like carbon enrichment (Kelly & Nixon 1984, Hansen & Blackburn 1991, 1992, van Duyl et al. 1992, Kristensen & Holmer 2001). At the farming site the benthic DIC release rate prior to the recovery period was 290 mmol m⁻² d⁻¹, a value that declined to 22 mmol m⁻² d⁻¹ after 39 d. Assuming that the carbon mineralization rate de-

creased exponentially during the period, and accounting for the DOC efflux, the total benthic mineralization equaled 6.1 mol m⁻², corresponding to only 33 % of OC removal as calculated from the sediment OC inventory above. Assuming a linear rather than an exponential decline in mineralization activity, the corresponding value for OC removal is 50 %. Thus both assumptions suggest the involvement of other processes in carbon removal during the recovery period. Potentially, methane release from the sediment could account for some of the 'missing' carbon, but more likely local resuspension of surface sediment caused by currents (and possibly by wild fish) played a bigger role. Resuspension would spread the carbon enrichment to the entire footprint area, and carbon mineralization and benthic solute exchange during resuspension events would be markedly enhanced. The latter aspect is not accounted for in our assessment of the benthic activity via standard core incubations. Indeed near-bottom trap data indicate some small degree of resuspension at all times, with resuspension actually increasing during the recovery period. At this time 51 ± 3.4 % and 63 ± 3.5 % of the flux at 40 and 45 m depth, respectively, was due to resuspended material (data not shown). Indications of resuspension may also be observed in the depth distribution of OC in the sediment (Fig. 6b), showing that the settling farming material with high OC content (27 ± 2.8 % of dry wt) is not merely deposited at the sediment surface, but that it is mixed deeper into the sediment with otherwise lower OC content. Since macrofauna was absent or scarce, this intense downward mixing can be attributed to relocation of sediment due to resuspension. It is likely that local resuspension was a major factor for the observed fast recovery below the farm.

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