INTRODUCTION

The phytoplankton spring bloom is a major, recurrent phenomena in temperate and cold planktonic systems, which fuels secondary production (Straile & Adrian 2000, Edwards & Richardson 2004, Wiltshire et al. 2008, Sommer & Lewandowska 2011). Global warming is known to result in an earlier onset of biological spring events, e.g. advancement of the phytoplankton spring bloom (Weyhenmeyer et al. 1999, Gerten & Adrian 2001, Weyhenmeyer 2001, Sommer et al. 2012, Winder et al. 2012). However, a delayed timing of spring bloom phenomena has also been reported, e.g. in the North Sea (Wiltshire & Manly 2004). Increased grazing by overwintering zooplankton at elevated temperatures has been suggested as an explanation for this delay. Only when light conditions improve during late spring does the growth rate of primary producers exceed the grazing losses of zooplankton (Rose & Caron 2007), thus opening a ‘loophole’ for the phytoplankton to form a bloom (Irigoien et al. 2005). This is related to the different...
temperature dependencies of photoautotrophic and heterotrophic processes. While light-limited photosynthesis is only weakly responsive to warming, the physiology and metabolism of heterotrophs is highly temperature-dependent (see the ‘metabolic theory of ecology [MTE]’ of Brown et al. [2004]). Such an unequal response of autotrophs and heterotrophs to warming is assumed to create shifts in the trophodynamic interactions of plankton (McCowan et al. 2003, Smol et al. 2005). Further, trophic pathways through protozooplankton (PZP) might be enhanced with warming due to short generation times and strong responses of PZP to elevated temperature conditions (Mueller & Geller 1993, Weisse & Montagnes 1998, Montagnes & Lessard 1999, Montagnes et al. 2003). In the light of a winter warming scenario for northern-central Europe (IPCC 2007), reduced time-lags between phytoplankton and PZP might cause imbalances between growth and the removal of phytoplankton, thus affecting phytoplankton bloom dynamics (Wiltshire et al. 2008) and composition (Keller et al. 1999, Sommer & Lengfellner 2008).

Traditionally, filter-feeding copepods have been considered to graze predominantly on large phytoplankton. This picture has changed during the last decades, and feeding modes of most copepods are now considered flexible. Omnivory is a common feature of copepods, and PZP (e.g. ciliates and heterotrophic flagellates) are known to contribute substantially to copepod diets (Jonsson & Tiselius 1990, Stoecker & Capuzzo 1990, Kleppel 1993, Calbet & Landry 1999, Sommer et al. 2005). In general, the feeding preferences of copepods are considered to depend strongly on the abundance and taxonomic composition of phytoplankton and PZP. Further, the role of PZP in improving the food quality for copepods by buffering nutritional imbalances at the interface between primary production and consumption (‘trophic upgrading’) has been stressed (Malzahn et al. 2010). On the other hand, PZP have the potential to modulate or even suppress phytoplankton blooms (Irigoien et al. 2005, Sherr & Sherr 2009, Löder et al. 2011). While ciliates feed mainly on nanoplankton (Jonsson 1986), heterotrophic dinoflagellates are able to suppress large, bloom-forming diatoms (Sherr & Sherr 2007).

In spring, overwintering copepod densities are usually low, and the copepods’ numerical response to increasing food sources is slow. In contrast, PZP grow instantaneously in response to increasing food sources (Johansson et al. 2004), often at rates comparable to phytoplankton growth rates, especially under warm conditions (Fenchel & Finlay 1983, Mueller & Geller 1993, Montagnes & Lessard 1999). In case of a tighter coupling between PZP and its algal food sources under elevated temperature conditions, enhanced feeding competition between micro- and mesozooplankton seems plausible. Such an enhanced trophic overlap could, in turn, enable PZP to function as a ‘trophic sink’ rather than a ‘trophic link’, thus reducing energy flow up the food web (Gifford 1991, Rollwagen-Bollens et al. 2011).

Experimental studies addressing the combined role of overwintering micro- and mesozooplankton on spring plankton dynamics are rare. The present study is unique in that it combines 2 driving factors affecting plankton dynamics during spring bloom formation: (1) temperature and (2) the density of overwintering copepods, in a fully factorial design.

To elucidate the interactions between PZP and copepods and their consequences for plankton dynamics during spring, we addressed the following hypotheses on the combined effects of warming and copepod density:

1. High overwintering copepod densities will affect PZP biomass ($B_P$) leading to shifts in PZP community composition and/or size classes
2. Warming at high copepod densities leads to a stronger top-down control of PZP by copepods (PZP functions as a ‘trophic link’)
3. Warming at low copepod densities leads to an earlier occurrence of PZP and reduced time-lags between phytoplankton and PZP production (PZP functions as a ‘trophic sink’).

**MATERIALS AND METHODS**

**Experimental design**

A detailed description of the experimental set-up in 2009 is given by Sommer & Lewandowska (2011). Twelve mesocosms with a volume of 1400 l each and a depth of 1 m were placed in 4 temperature-controlled rooms. We used 3 zooplankton levels (1.5, 4 and 10 copepods l$^{-1}$) and 2 temperature levels, in the following called $\Delta 0^\circ C$ and $\Delta 6^\circ C$ ($0^\circ C$ and 6$^\circ C$ above the present-day temperature scenario for the Kiel Bight). Actual temperatures in the 12 mesocosms at ambient ($\Delta 0^\circ C$) and elevated ($\Delta 6^\circ C$) temperature regimes are shown in Fig. 1. Three zooplankton levels where established per room, and 2 of the rooms were run at the same temperature regimes (for each temperature level). All mesocosms were filled simultaneously with water from the Kiel Bight, Baltic Sea, assuring similar initial conditions. Damage through...
the pumping procedure was tolerable for phytoplankton and PZP, but mesozooplankton organisms did not survive this transfer. Mesozooplankton, comprising mainly copepods (70% *Acartia* sp., 17% *Oithona similis*, 7% *Pseudocalanus* sp., 3% *Temora longicornis* and 3% *Centropages hamatus*), was added from net-catches at density ranges similar to the ones found in the Kiel Bight in spring (Behrends 1996). The copepod composition remained relatively constant throughout the experiment, and no great changes were observed in any of the 12 mesocosms. Details on phytoplankton and mesozooplankton composition changes have been given in previous publications dealing with the same mesocosm study. Changes in phytoplankton composition and size ranges are found in the study by Sommer & Lewandowska (2011). Furthermore, the paper by Lewandowska et al. (2014) provides details on the zooplankton and stresses that composition did not change during the course of the mesocosm experiment.

The Δ0°C temperature treatment was based on the average seasonal pattern of surface temperatures for the Kiel Bight during the years 1993–2002. The elevated Δ6°C temperature treatment was intended to simulate future warming scenarios predicted by the IPCC (2007) for the end of the century. The natural seasonal temperature increase during spring was simulated starting with mean temperatures of 2.4°C (Δ0°C) and 8.4°C (Δ6°C), respectively. For details on the actual temperature regimes in the mesocosms see Fig. 1.

Light was supplied by computer-controlled light units (Profilux II, GHL Groß Hard- and Software Logistics) above each mesocosm, simulating daily triangular light curves, and the seasonal change in light climate was calculated using the astronomic model published by Brock (1981).

The starting date for the light and temperature programs was set to a theoretical start on 15 February (Day-of-the-year [DOY] 46) in order to allow comparisons with previous experiments using the same mesocosm set-up (Sommer et al. 2012). The actual starting date was 9 January 2009. The experiment was terminated after 38 d.

Details on phytoplankton and copepod sampling are given by Sommer & Lewandowska (2011).

### PZP sampling and enumeration

Weekly seawater samples from the mesocosms were taken for PZP enumeration and identification. The seawater was transferred to 250 ml bottles, fixed with acidic Lugol’s iodine (2% final concentration), settled in 100 ml sedimentation chambers and counted with a Zeiss Axiovert 135 at 200× magnification using an inverted microscope technique (Utermöhl 1958). Identification to the lowest possible taxonomic level (species or genus level) was made using Kahl (1932), Foissner et al. (1991–1995), Tomas (1996), Strüder-Kypke et al. (2002), Scott (2005). Geometric proxies following Hillebrand et al. (1999) where used.
for biovolume calculations and thereafter converted into carbon biomass using the conversion factors given in Putt & Stoecker (1989).

**Statistical analysis**

We fitted sigmoidal curves to the temporal evolution of cumulative PZP biomass; \( f(x) = \frac{a}{1 + \exp[-(x - x_0)/b]} \), where \( x \) is time. As sigmoidal curves are derivatives of bell-shaped curves, we consider the inflection point \( x_0 \) as the bloom timing and \( b \) as the bloom duration. We used ANCOVAs for further investigation on the effects of temperature (categorical factor) and log copepod densities (quantitative factor). Further interactions between the 2 factors on the bloom timing \( x_0 \), the bloom duration \( b \) and the bloom intensity \( B_p \) at \( D_{\text{max}} \) (the day when \( B_p \) maxima were reached) were analyzed. We excluded 2 \( x_0 \) values and 1 \( b \) value from the analysis, since they did not fit the sigmoidal curve and were hence omitted from the ANCOVA. Both cases were \( \Delta T \)C mesocosms, one with low and one with intermediate copepod densities. We used ANCOVAs to evaluate the influence of temperature (categorical factor) and log copepod densities (quantitative factor), as well as interactions between the 2 factors, on the relative biomass contribution of dinoflagellates >30 \( \mu \)m, dinoflagellates <30 \( \mu \)m, ciliates >30 \( \mu \)m and ciliates <30 \( \mu \)m at the biomass maximum, which turned out to always be the nearest sampling date to \( x_0 \) derived from the non-linear fitting described above. We applied a Bonferroni correction to the significance level \( \alpha \) (\( \alpha \)/number of tests). This was necessary due to the multiple tests performed on the same dataset. Hence, we set an \( \alpha \) level of 0.025 for timing (2 tests) and duration of the bloom and an \( \alpha \) level of 0.01 (5 tests) for the 5 biomass-related measures (total biomass, small ciliates, large ciliates, small dinoflagellates and large dinoflagellates). A list of variables, including abbreviations and units, is given in Table 1. We used SigmaPlot 10.0 for the non-linear curve fitting and Statistica 8.0 for the ANCOVAs.

**RESULTS**

In the warm mesocosms (\( \Delta T \)C) \( B_p \) showed a clear numerical response to increasing phytoplankton availability. Shortly after the peak of the phytoplankton bloom, PZP reached \( D_{\text{max}} \) on DOY 58 and decreased below the initial levels thereafter (Table 2, Fig. 2A–C). In the cold mesocosms, \( B_p \) responded more slowly to increases in phytoplankton biomass, and clear increases could only be observed in the mesocosms with low and moderate copepod densities (Fig. 2D,E). Here, \( D_{\text{max}} \) was reached at DOY 65 or 79 (Table 2). In the cold mesocosms with high copepod densities, \( B_p \) increased only slightly, and the biomass levels reached were much lower than those in the low and moderate copepod density treatments (Fig. 2F).

The bloom intensity (i.e. the \( B_p \) at \( D_{\text{max}} \)) was significantly affected by both temperature and copepod densities as well as their interaction (Table 3). At both temperatures, increasing copepod abundance decreased the bloom intensity; however, the decreasing effect was stronger in the \( \Delta T \)C treatment (Fig. 3), \( x_0 \) and \( b \) were not influenced by copepod biomass, but were significantly accelerated and shortened in the warm mesocosms (Table 3). Species-specific PZP biomass at \( D_{\text{max}} \) is shown in Fig. 4. In general, the small strobilidiid ciliate *Lohmaniella oviformis* dominated the spring PZP bloom community in all 12 mesocosms, reaching its highest \( B_p \) at \( D_{\text{max}} \) in

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**Table 1. List of variables.** PZP: protozooplankton; DOY: day-of-the-year.

<table>
<thead>
<tr>
<th>Name</th>
<th>Abbreviation</th>
<th>Unit</th>
</tr>
</thead>
<tbody>
<tr>
<td>PZP biomass</td>
<td>( B_p )</td>
<td>( \mu g \text{ C l}^{-1} )</td>
</tr>
<tr>
<td>Timing of the ( B_p ) maximum</td>
<td>( D_{\text{max}} )</td>
<td>DOY</td>
</tr>
<tr>
<td>Bloom timing</td>
<td>( x_0 )</td>
<td>DOY</td>
</tr>
<tr>
<td>Bloom duration</td>
<td>( b )</td>
<td>d</td>
</tr>
<tr>
<td>Bloom intensity</td>
<td>( B_p ) at ( D_{\text{max}} )</td>
<td>( \mu g \text{ C l}^{-1} )</td>
</tr>
</tbody>
</table>

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**Table 2.** Protozooplankton bloom timing and duration (see Table 1) at mesocosm temperatures and copepod densities. \( \Delta T \): temperature elevation; copepod density: initial copepod density in the mesocosm; ns: not significant.

<table>
<thead>
<tr>
<th>( \Delta T ) (°C)</th>
<th>Copepod density (DOY)</th>
<th>( x_0 )</th>
<th>95% CI</th>
<th>( b ) (d)</th>
<th>95% CI</th>
<th>( D_{\text{max}} ) (DOY)</th>
</tr>
</thead>
<tbody>
<tr>
<td>0</td>
<td>1.5</td>
<td>ns</td>
<td>95.6</td>
<td>47.7, 143.4</td>
<td>7.16</td>
<td>5.09, 9.23, 79</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>4</td>
<td>61.7, 62.8</td>
<td>4.22</td>
<td>3.79, 4.65, 65</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>10</td>
<td>60.0, 66.1</td>
<td>7.28</td>
<td>5.36, 9.19, 65</td>
</tr>
<tr>
<td>1</td>
<td>5.1</td>
<td>ns</td>
<td>62.6</td>
<td>57.3, 67.8</td>
<td>8.94</td>
<td>5.85, 12.03, 65</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>6</td>
<td>54.1, 60.1</td>
<td>2.53</td>
<td>2.21, 2.84, 58</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>54.9</td>
<td>54.2, 55.6</td>
<td>1.62</td>
<td>1.29, 1.96, 58</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>56.0</td>
<td>55.6, 56.4</td>
<td>2.37</td>
<td>2.08, 2.65, 58</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>57.8</td>
<td>57.5, 58.0</td>
<td>2.19</td>
<td>1.81, 2.57, 58</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>56.7</td>
<td>56.1, 57.2</td>
<td>2.66</td>
<td>2.14, 3.18, 58</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>56.8</td>
<td>55.8, 57.9</td>
<td>3.08</td>
<td>2.18, 3.99, 58</td>
</tr>
</tbody>
</table>
Variable & $F_{\text{Temp}}$ & $P_{\text{Temp}}$ & $F_{\text{Copepod density}}$ & $P_{\text{Copepod density}}$ & $F_{\text{Interaction}}$ & $P_{\text{Interaction}}$ & $F_{\text{Model}}$ & $P_{\text{Model}}$ \\ 
Bloom timing ($x_0$) & 11.73 & $<0.025$ & 4.05 & 0.09 & 4.45 & 0.08 & 229.61 & $<0.025$ \\ 
Bloom duration ($b$) & 9.41 & $<0.025$ & 0.55 & 0.48 & 0.01 & 0.97 & 28.81 & $<0.025$ \\ 
Bloom intensity ($B_P$ at $D_{\text{max}}$) & 13.02 & $<0.01$ & 42.26 & $<0.01$ & 11.55 & $<0.01$ & 97.08 & $<0.01$

Table 3. ANCOVA results, calculated with temperature as the categorical factor, log zooplankton initial abundance as the quantitative factor and the interaction term, as well as bloom timing, duration and intensity as dependent variables.
the Δ0°C treatments with low copepod densities (Fig. 4A). Overall, increasing copepod densities led to reduced $B_p$ at $D_{max}$ in both temperature regimes (Fig. 4A,B).

Proportions of the different PZP size classes (dinoflagellates >30 µm, dinoflagellates <30 µm, ciliates >30 µm and ciliates <30 µm) are given in Fig. 5. ANCOVAs on the different PZP groups and size classes showed that the factor copepod density significantly affected the response of ciliate biomass ($p < 0.01$) and ciliates <30 µm ($p < 0.01$) at $D_{max}$, while the factor ‘temperature’ and the interaction term did not have any significant impact (Table 4). Dinoflagellates and all other PZP size classes did not show any significant effects in relation to copepod density and/or temperature.

### DISCUSSION

While reduced time lags between phytoplankton and PZP occurrence under elevated temperature conditions are considered typical phenomena (Aberle et al. 2007, 2012, Winder et al. 2012), the interactions between warming and the top-down effects of mesozooplankton are not clear yet. One of the key questions on the responses of plankton communities to warming is whether grazing pressure can be strong enough to prevent a phytoplankton bloom in early spring, as has been hypothesized for delayed spring blooms during warmer late winter/early spring periods in situ (Wiltshire & Manly 2004). Here, we observed clear phytoplankton bloom formation during early spring under all combinations of temperature and copepod density irrespective of an advanced peak timing of PZP in the warmer mesocosms and overall grazer densities. This suggests the complex nature of spring bloom phenomena, which are influenced by a multitude of factors (e.g. PZP-mesozooplankton interactions, selective feeding of zooplankton).

#### Effects of overwintering copepod density

While copepod densities did not affect the timing and duration of the PZP bloom, the bloom intensity was negatively affected by increasing copepod density. Thus, Hypothesis 1, stating that overwintering copepod densities will affect PZP biomass was partially confirmed. To date, it is accepted knowledge that copepods suppress PZP efficiently and that omnivory plays a major role in copepod feeding behaviour (Jonsson & Tiselius 1990, Stoecker & Capuzzo 1990, Kleppel 1993, Calbet & Landry 1999, Sommer et al. 2005). Using PZP as an additional food item has 3 advantages for copepods: (1) broadening the food spectrum and thereby increasing the amount of available food, (2) suppressing resource competitors sensu ‘eating your competitor strategy’ (Thingstad et al. 1996) and (3) ‘trophic upgrading’ which results by e.g. conditioning the lipid content and lipid species of the food (Klein Breteler et al. 1999, Tang & Taal 2005) or buffering nutritional imbalances by heterotrophic protists which function as trophic intermediaries between phytoplankton and secondary consumers (Stoecker & Egloff 1987, Malzahn et al. 2010). Especially in nutrient-depleted situations, phytoplankton show a high carbon to nutrient ratio (Goldman et al. 1979). This creates an imbalance between the carbon and nutrient supply by the phytoplankton and the demand of consumers. This imbalance is usually reflected by the reduced growth and reproductive rates of the consumer (Elser et al. 2000, Malzahn & Boersma 2012). From a copepods’ point of view, the ability to feed on an interme-
Fig. 5. Proportion of the different protozooplankton (PZP) size classes at specific times during the experiment (DOY 58, 65 and 79) at 2 different temperature regimes and 3 initial copepod densities. PZP size classes: heterotrophic dinoflagellates >30 µm, heterotrophic dinoflagellates <30 µm, ciliates >30 µm and ciliates <30 µm. Orange background indicates $D_{\text{max}}$, the peak timing of PZP biomass ($\Delta 6^\circ\text{C}$ treatments: $D_{\text{max}}$ on DOY 58; $\Delta 0^\circ\text{C}$ treatments: $D_{\text{max}}$ on DOY 65 or 79)

Table 4. ANCOVA on various variables (biomass) at $D_{\text{max}}$, with temperature as the categorical factor, log zooplankton initial abundance as the quantitative factor and the interaction term

<table>
<thead>
<tr>
<th>Variable</th>
<th>$F_{\text{Temperature}}$</th>
<th>$P_{\text{Temperature}}$</th>
<th>$F_{\text{Copepod density}}$</th>
<th>$P_{\text{Copepod density}}$</th>
<th>$F_{\text{Interaction}}$</th>
<th>$P_{\text{Interaction}}$</th>
<th>$F_{\text{Model}}$</th>
<th>$P_{\text{Model}}$</th>
</tr>
</thead>
<tbody>
<tr>
<td>Dinoflagellates &gt;30 µm</td>
<td>0.87</td>
<td>0.38</td>
<td>0.75</td>
<td>0.41</td>
<td>0.44</td>
<td>0.53</td>
<td>5.33</td>
<td>&lt;0.01</td>
</tr>
<tr>
<td>Dinoflagellates &lt;30 µm</td>
<td>5.79</td>
<td>0.04</td>
<td>2.88</td>
<td>0.13</td>
<td>1.23</td>
<td>0.30</td>
<td>14.70</td>
<td>&lt;0.01</td>
</tr>
<tr>
<td>Ciliates &gt;30 µm</td>
<td>1.17</td>
<td>0.31</td>
<td>4.49</td>
<td>0.07</td>
<td>1.07</td>
<td>0.33</td>
<td>13.69</td>
<td>&lt;0.01</td>
</tr>
<tr>
<td>Ciliates &lt;30 µm</td>
<td>6.21</td>
<td>0.04</td>
<td>13.31</td>
<td>&lt;0.01</td>
<td>3.99</td>
<td>0.08</td>
<td>34.69</td>
<td>&lt;0.01</td>
</tr>
</tbody>
</table>
diary consumer might indeed mean lower food availability. But, feeding on organisms which already buffer nutritional imbalances to a large degree will, in turn, benefit the secondary consumer (Malzahn et al. 2010). As shown by Löder et al. (2011), the feeding preferences of copepods are rather flexible, and the question whether they prefer phytoplankton instead of PZP strongly depends on the community composition of the food sources, as well as on the nutrient status in the water column, which affects the quality of food items preyed upon by consumers.

Further, we hypothesized that alterations in overwintering copepod densities will not only affect PZP biomass per se, but also PZP community composition and/or size classes due to differential grazing of copepods. In our study, we found a strong negative effect of increasing copepod densities on ciliate biomass, and, more specifically, a strong suppression of small-sized ciliates (ciliates <30 µm). In general, a selectivity of omnivorous copepods for ciliates has often been reported (Nejstgaard et al. 2001, Vincent & Hartmann 2001, Jakobsen et al. 2005, Schnetzer & Caron 2005). Such a feeding preference has been linked to the favorable nutritional quality of ciliates as opposed to phytoplankton prey, thus stressing the relevance of ciliates in the diets of copepods such as Acartia sp., the dominant copepod species in our mesocosms (Stoecker & Egloff 1987, Gifford & Dagg 1988, Kleppel et al. 1991, Kleppel 1993). This is especially true during the decay of a phytoplankton bloom, when nutrients become limited and phytoplankton turns into a poor-quality food item (Löder et al. 2011). The small strobilidiid ciliate Lohmaniella oviformis acted as the main bloom-forming species during this mesocosm experiment and contributed almost exclusively to the ciliate size class <30 µm. L. oviformis is considered a typical ciliate species during the winter−spring transition (Johansson et al. 2004, Aberle et al. 2007, 2012), and, due to its bloom formation, it can be speculated that it served as an abundant and well-edible prey item for copepods in our mesocosms. So far, many omnivorous copepods have been shown to feed selectively on specific ciliate species (Stoecker & Capuzzo 1990, Nishibe et al. 2010, Wu et al. 2010) and, among other selection criteria (e.g. swimming behaviour, nutritional status of the prey), size seems to play only a minor role with regard to the feeding behaviour of copepods (Jons son & Tiselius 1990, Broglio et al. 2001, Gismervik 2006). Therefore, it seems more likely that the preference for ciliates <30 µm in our mesocosms resulted from selective grazing of copepods (e.g. the calanoid copepods Acartia sp.) on the ciliate L. oviformis rather than from size-specific copepod predation. This is in line with studies on the feeding behaviour of Acartia sp., showing high clearance rates and a high capture efficiency when preying upon L. oviformis (Tiselius 1989, Gismervik 2006).

In contrast to the strong suppression of ciliates by copepods, heterotrophic dinoflagellates remained unaffected by copepod densities in our study. Although both thecate and athecate dinoflagellates are known to be consumed by copepods (Sommer et al. 2005), a preference for ciliates over dinoflagellates is documented in the literature in general (Vincent & Hartmann 2001, Schnetzer & Caron 2005) and for Acartia sp. more specifically (Jakobsen et al. 2005). Here, we consider the thecate dinoflagellate Protoperidinium bipes, which occurred at moderate abundances in our mesocosms, as a less preferred food item in our mesocosm, which resulted in a release from grazing pressure by copepods.

Combined effects of warming and copepod density

Hypotheses 2 and 3 aimed at analysing trends in energy transfer at the interface between microbial and classical food chains in the light of future warming conditions. We hypothesized that warming at high copepod densities will lead to a strong top-down control by copepods, thus supporting the PZPs’ role as a ‘trophic link’ by channelling carbon consumed by PZP through the ‘classic’ food chain. Warming at low copepod densities, however, should hypothetically lead to an earlier timing of PZP, thus functioning as a ‘trophic sink’ due to reduced time-lags between phytoplankton and PZP production.

With regard to Hypothesis 2, we found a strong effect of temperature and copepod density on the bloom intensity of PZB. At high temperatures, PZP was suppressed substantially throughout all copepod treatments, with a slightly stronger impact on PZP bloom intensity at high copepod density. At low temperature conditions and the lowest copepod density, however, PZP fully escaped top-down control by copepods. Thus, our hypothesis that the combined effects of warming and overwintering copepod densities will lead to stronger top-down control in the plankton was confirmed. Our findings are supported by the results of other studies showing that enhanced grazing by copepods as a result of higher metabolic activities is expected in relation to warming (Isla et al. 2008, O’Connor et al. 2009), and, consequently, higher copepod densities at elevated temperatures are likely to result in a stronger suppression of PZP.
At low temperatures and low overwintering copepod density, however, PZP growth rates seemed to exceed grazing losses, thus leading to a release from copepod predation and enabling PZP to form intense blooms in the cold mesocosms.

Further, we found a strong effect of temperature on PZP bloom timing and bloom duration at elevated temperatures, leading to an advanced peak timing, but shorter bloom duration, irrespective of copepod densities. This leads to a rejection of Hypothesis 3, assuming that warming in combination with copepod densities will result in an earlier timing and reduced time-lags between phytoplankton and PZP. Simulations conducted during a previous study using the same mesocosm set-up showed that the reduction of time lags between primary producers and consumers increases with increasing $Q_{10}$ values of PZP (Aberle et al. 2012). Consequently, warming alone strengthens the match between PZP and its algal prey which, in turn, promotes the phytoplankton–PZP link. In general, PZP is known to suppress phytoplankton standing stocks efficiently (Sherr & Sherr 2007, 2009, Löder et al. 2011). Based on an analysis from the same mesocosm study, Sommer & Lewandowska (2011) reported a shift in phytoplankton composition from larger diatoms to smaller, flagellated phytoplankton under elevated temperature conditions. Such a shift in phytoplankton community composition might have created a positive feedback loop, where PZP growth rates were boosted not only by elevated temperatures, but also by a phytoplankton community that contained optimal prey for PZP. As an example, the rapid growth and decline in *L. oviformis* biomass might be related to an overexploitation of algae between 2 and 15 µm, the preferred food of this ciliate species (Jonsson 1986, Christaki et al. 1998). In the given case, enhanced grazing by the ciliate *L. oviformis* on small-sized phytoplankton might channel more energy away from the direct phytoplankton–copepod link. Since *L. oviformis* is considered to serve as a preferred food item for copepods (Tiselius 1989, Gismervik 2006) and its growth will rather benefit from warming conditions, it seems more likely that *L. oviformis* will act increasingly as a ‘trophic link’ to copepods (and copepod-feeding fish). However, a completely different picture might evolve if other PZP groups such as heterotrophic dinoflagellates are promoted by changes in environmental conditions. Since heterotrophic dinoflagellates are considered efficient consumers of blooming diatoms (Sherr & Sherr 2007, Calbet 2008), direct feeding competition with copepods would occur, thus creating a ‘trophic sink’ and reducing energy flow up the food web. Such a trend would support the predictions of Berglund et al. (2007) indicating a reduced energy transfer to higher trophic levels in relation to climate change when the direct pathway from phytoplankton to mesozooplankton is interrupted by an intermediary trophic level comprised of heterotrophic dinoflagellates.

**CONCLUSIONS**

The results of the present study provide evidence that warming in combination with increases in over-wintering copepod densities will have considerable impact on the bloom intensity of PZP and top-down control mechanisms in the plankton. While warming will support the phytoplankton–PZP link in general, the distinct trophic roles of ciliates as opposed to dinoflagellates must also be taken into consideration. While stimulated growth of ciliates might enhance energy transfer efficiency to higher trophic levels, dinoflagellates might act as more of a ‘trophic sink’ in future plankton communities by supporting microbial pathways and reducing energy transfer up the food web as well as reducing export production to the deep oceans.

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**LITERATURE CITED**

Püt M, Stoecker DR (1989) An experimentally determined
carbon:volume ratio for marine ‘oligotrichous’ ciliates from estuarine and coastal waters. Limnol Oceanogr 34: 1097–1103


Scott FJE (2005) Antarctic marine protists. ABRS, Canberra


Sommer U, Lewandowska A (2011) Climate change and the phytoplankton spring bloom: warming and overwintering zooplankton have similar effects on phytoplankton. Glob Change Biol 17:154–162


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